## Opening of the Central Atlantic Ocean: implications for geometric rifting and asymmetric initial seafloor spreading after continental breakup.

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#### Oceanic crust Zone of crusta thinning (Necking) cruci Upper crust Oceanic crust 7.25 6.25 -AC 7.25 6.5 Middle crust - 20 () Depth () - 30 Lower crust MIRROR-1 100 200 Distance (km) Unthinned continental Zone of crustal thinning (Necking) тос Oceanic crus Upper crust 10 -Middle crus PSM <u>کی</u> 20 SMART-1 200 300 Distance (km) uem lfremer INSTITUT UNIVERSITAIR EUROPEEN DE LA ME

#### <sup>.</sup> Introduction

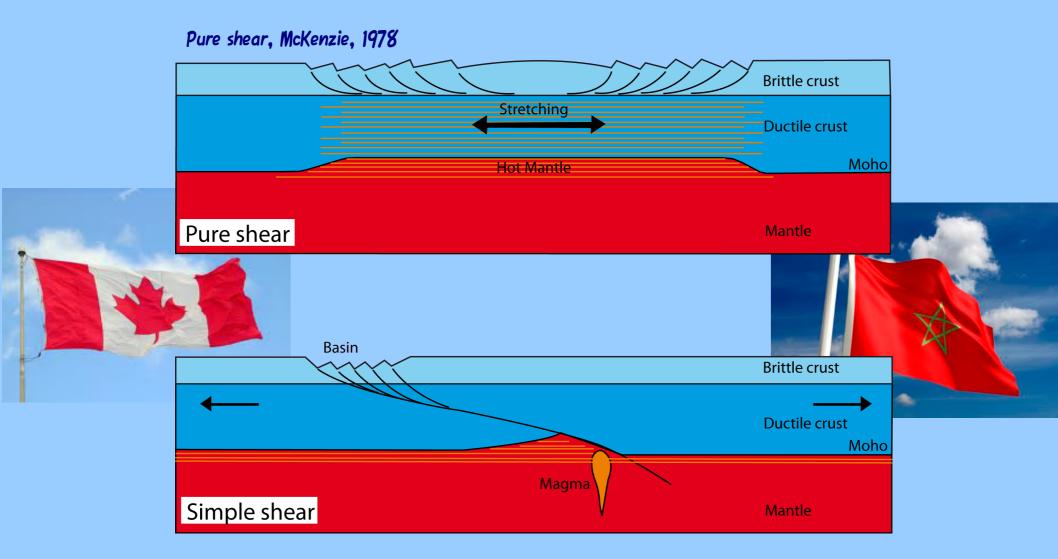
- <sup>•</sup> Deep structure of the Moroccan Margin
- <sup>.</sup> Comparison of the conjugate margins
- · Opening mechanism
- · Conclusions





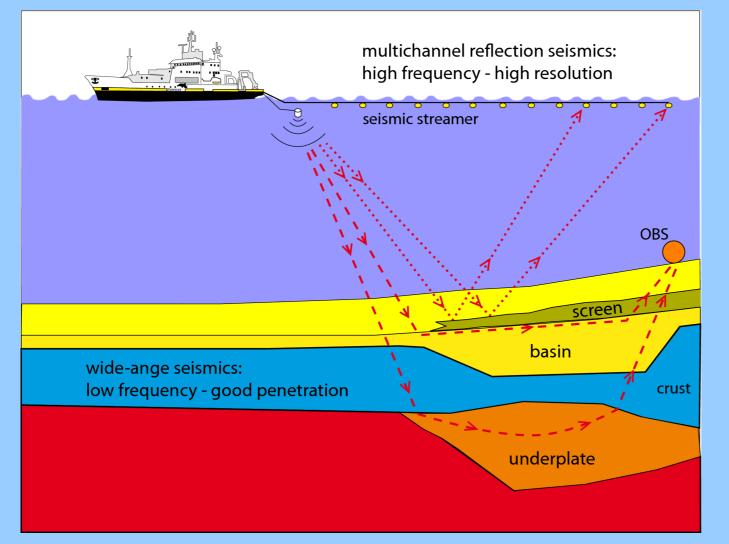


#### Why conjugate margins ?



Simple shear, Wernicke, 1985

## Why OBS

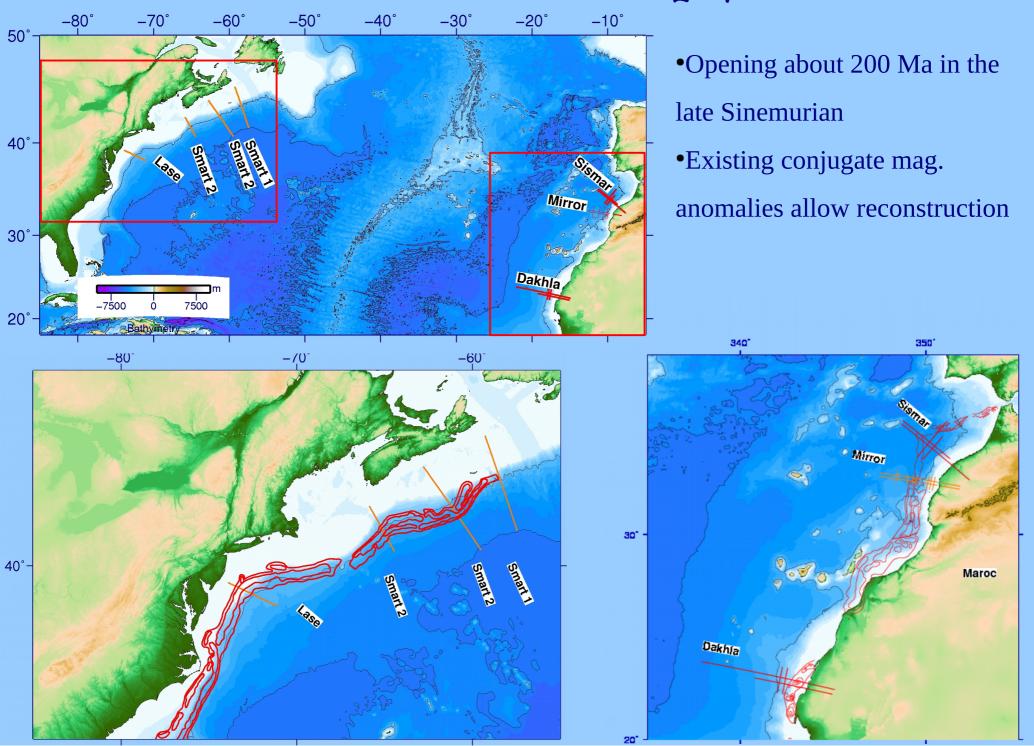


- Imaging below screen (basalt, salt)
- True geometry
- Lithology

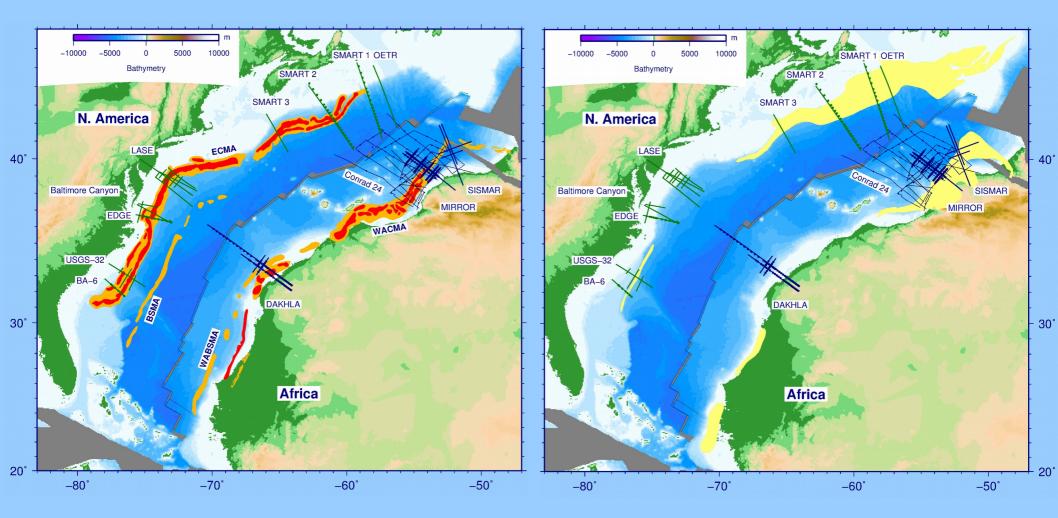
eg. underplate, crustal nature



#### The Morocco-Nova Scotia margin pair



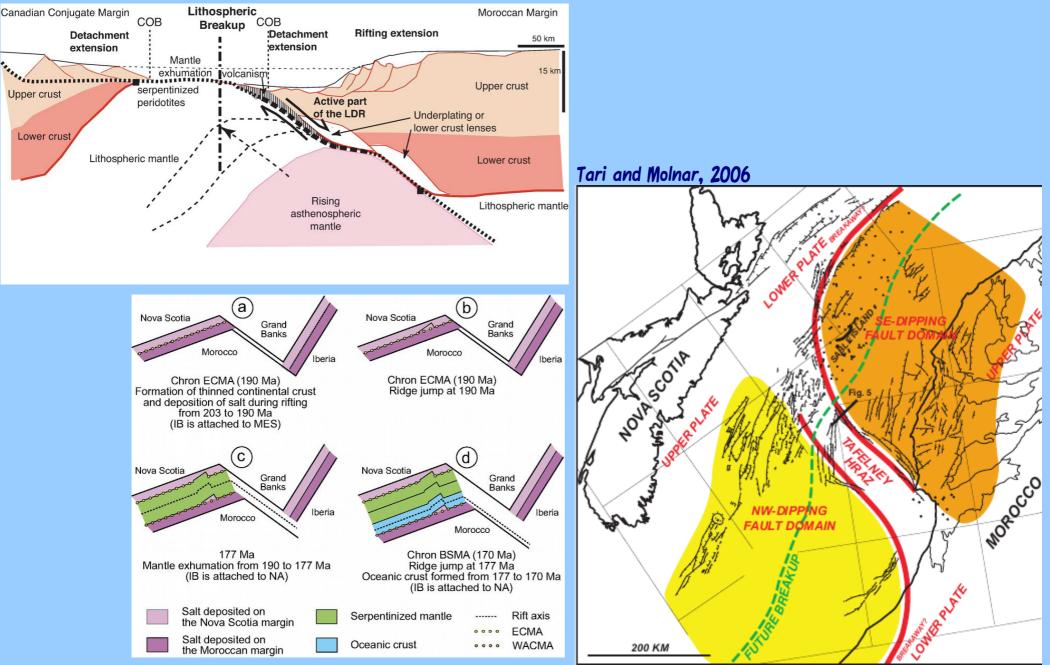
#### Plate kinematic reconstruction at M25



Reconstruction at M25 after Sahabi et al., 2005

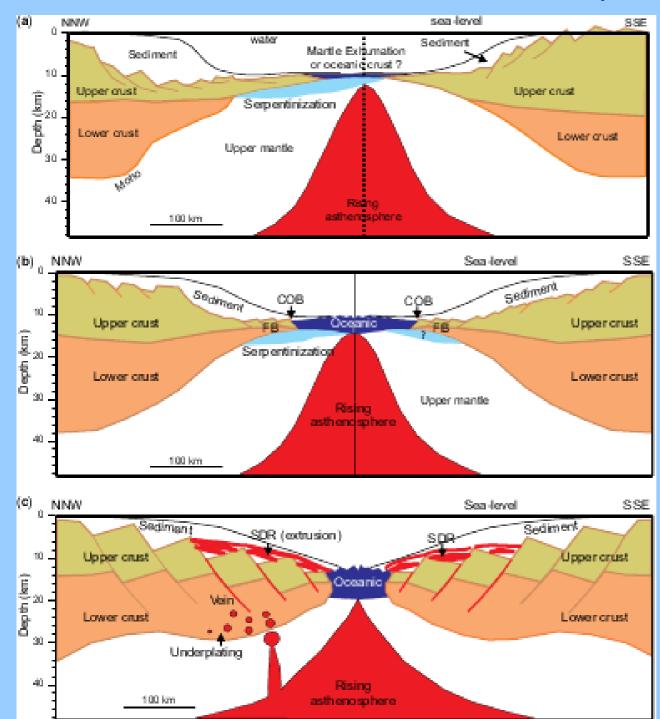
### Opening mechanism

#### Maillard et al., 2005



Sibuet et al., 2012

### Volcanism at opening



**Non-volcanic** Exhumation of upper mantle material

#### **Transition**

**Reduction of volcanism** 

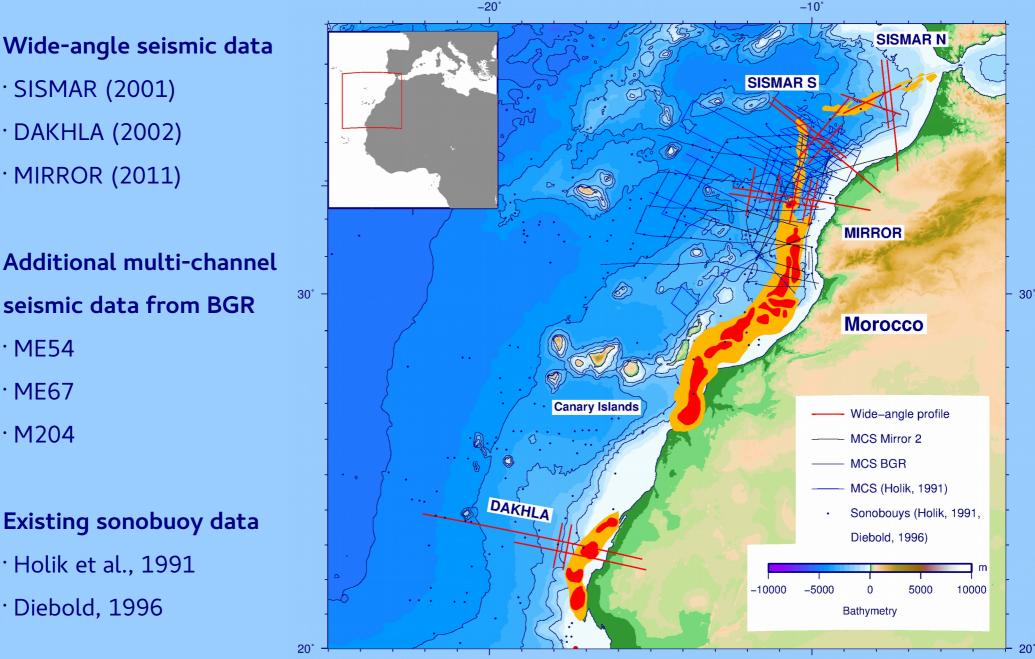
Volcanic

SDR, Underplating

Louden et al., 2012

# The NW-African continental margin

-20°



-10°

· SISMAR (2001)

· DAKHLA (2002)

• MIRROR (2011)

Additional multi-channel seismic data from BGR

• ME54

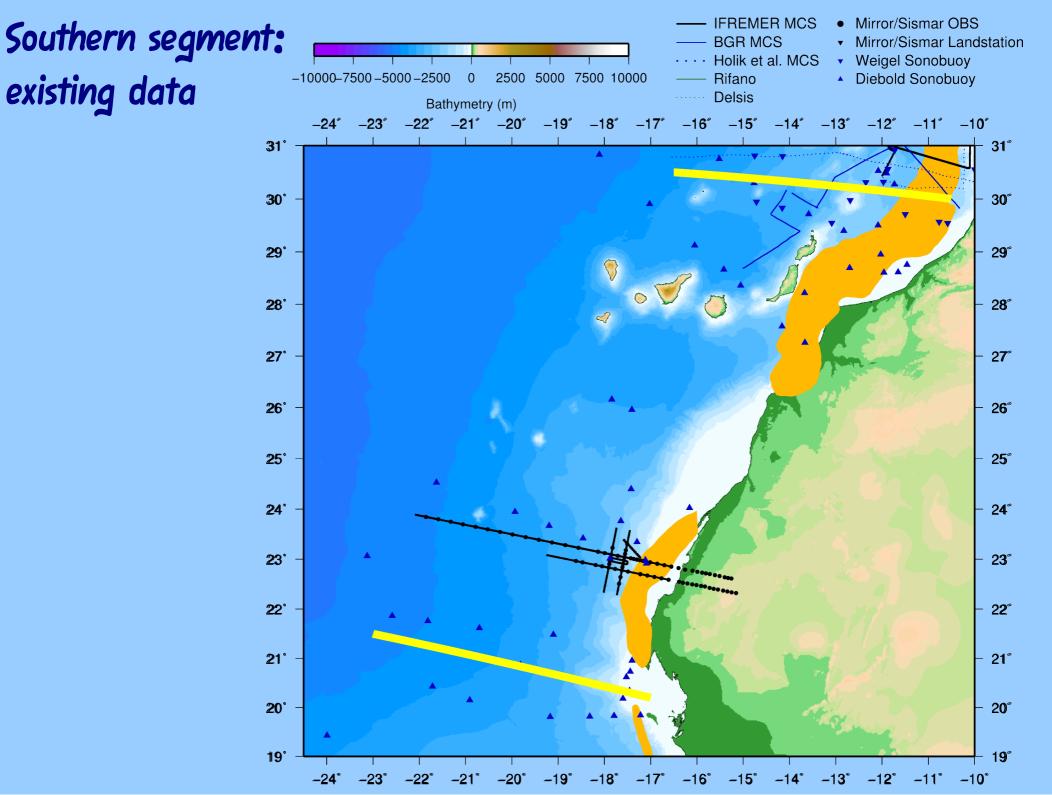
• ME67

· M204

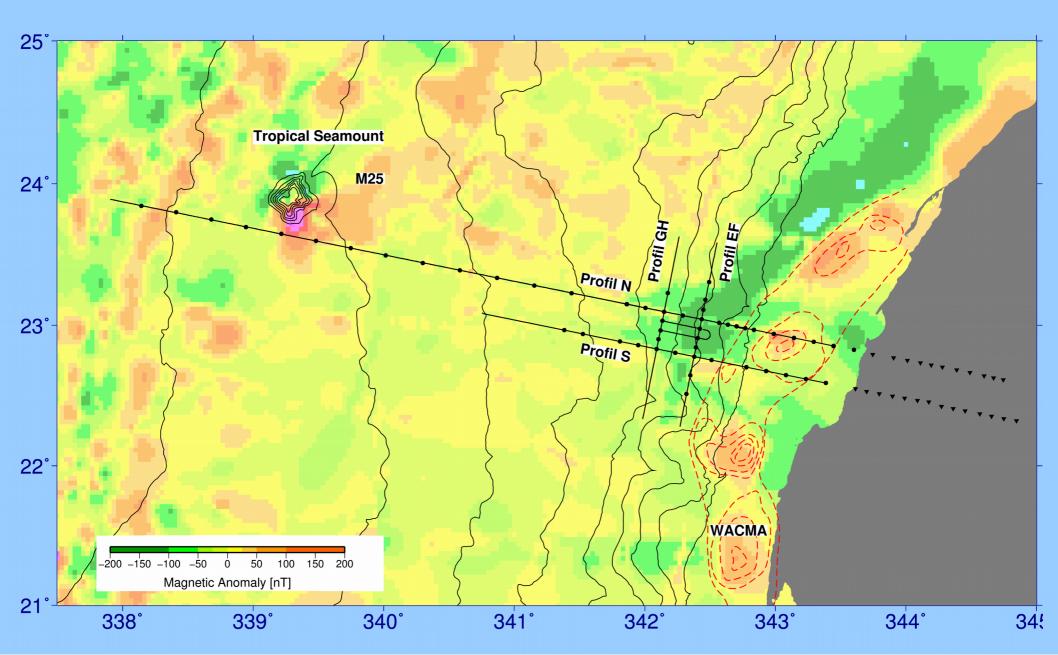
#### Existing sonobuoy data

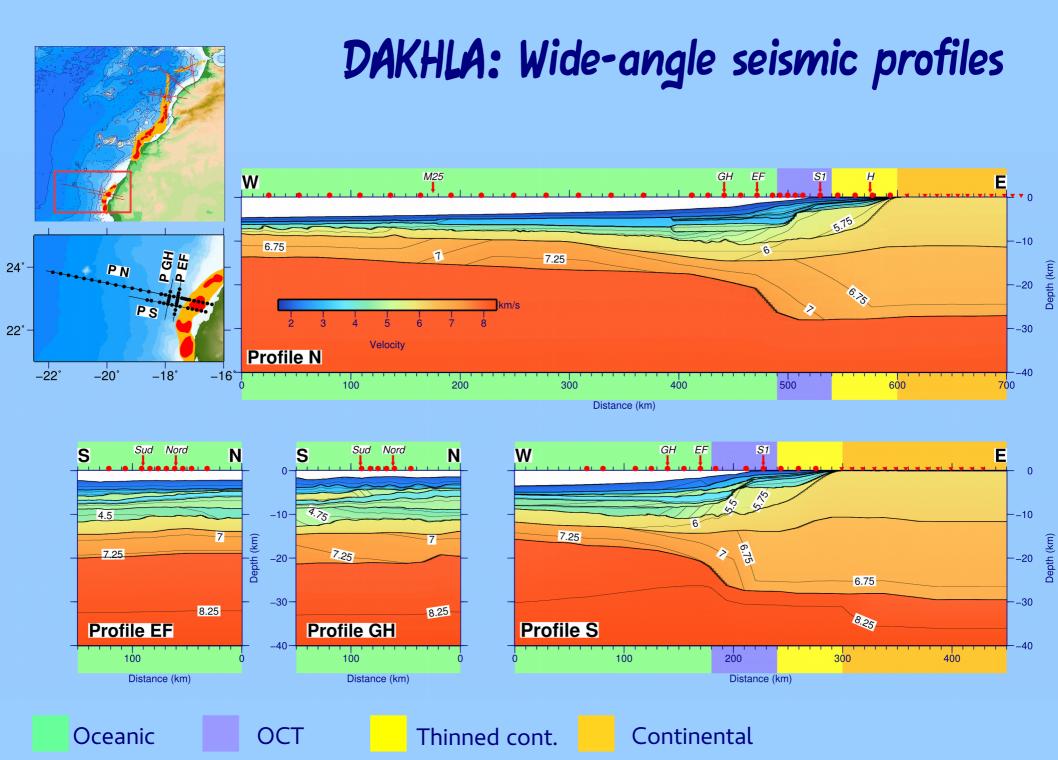
<sup>.</sup> Holik et al., 1991

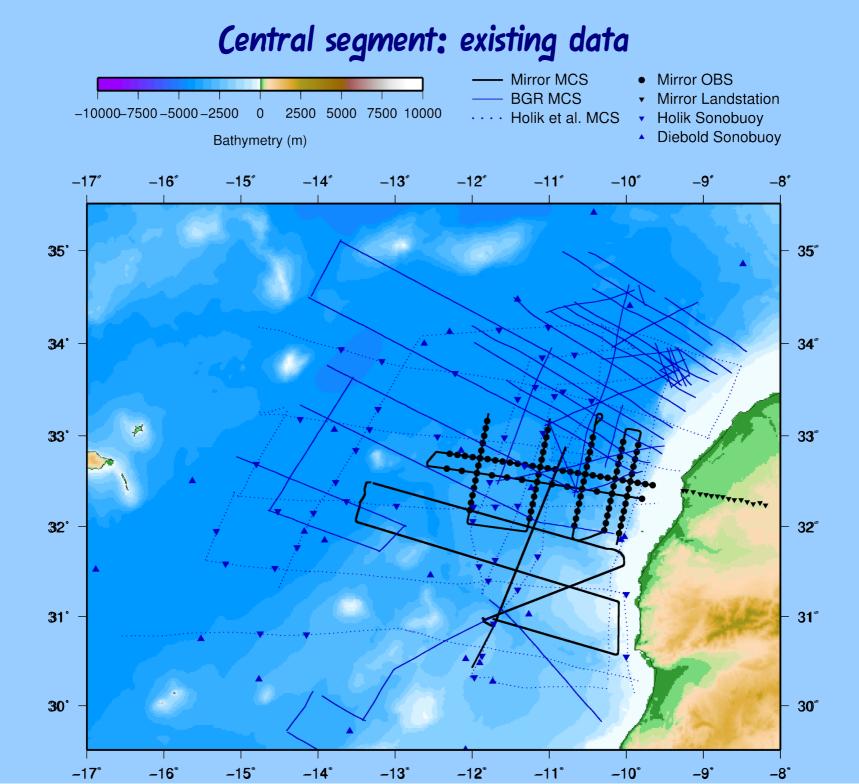
<sup>.</sup> Diebold, 1996

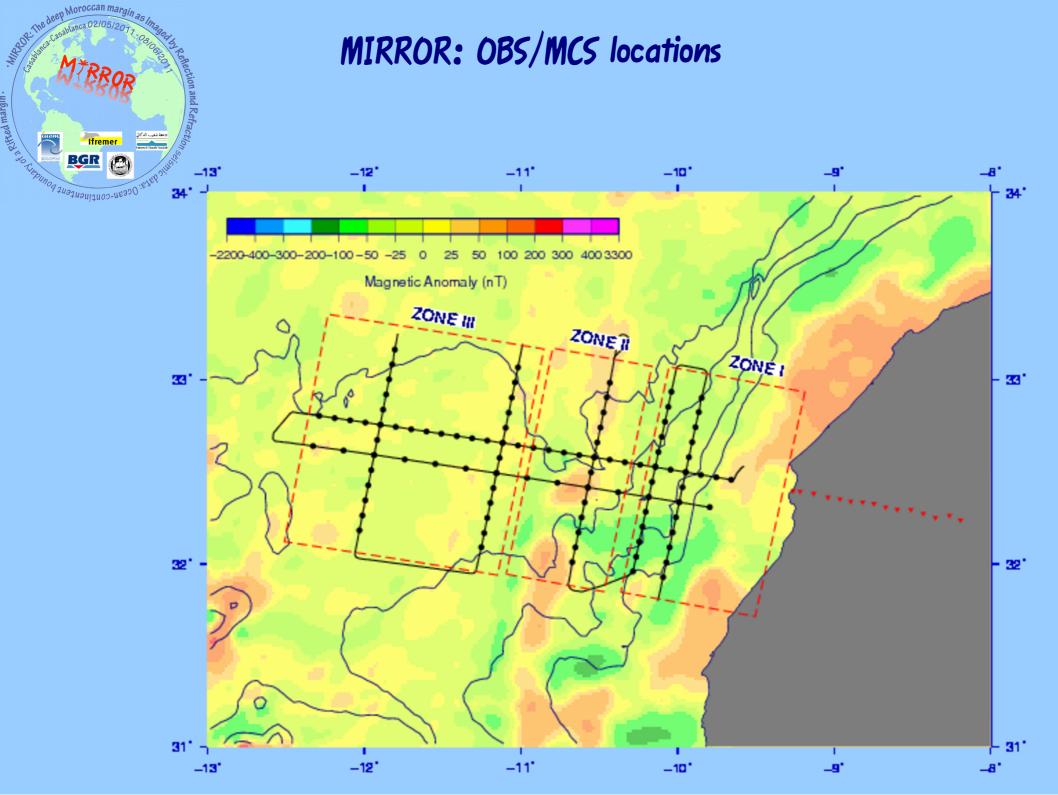


#### DAKHLA: magnetic anomaly

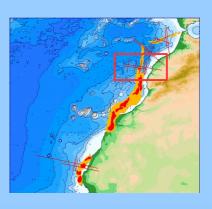




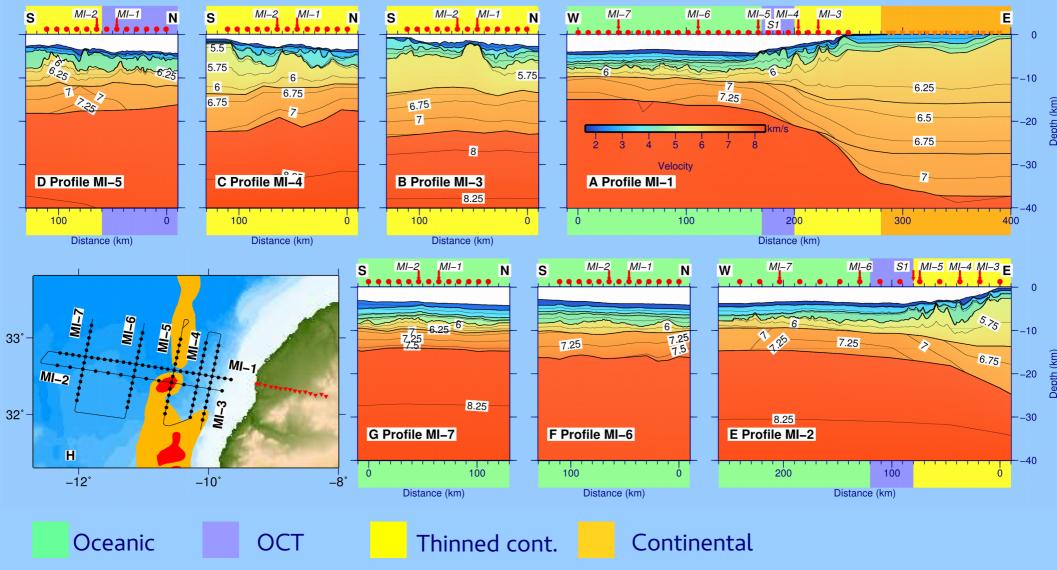


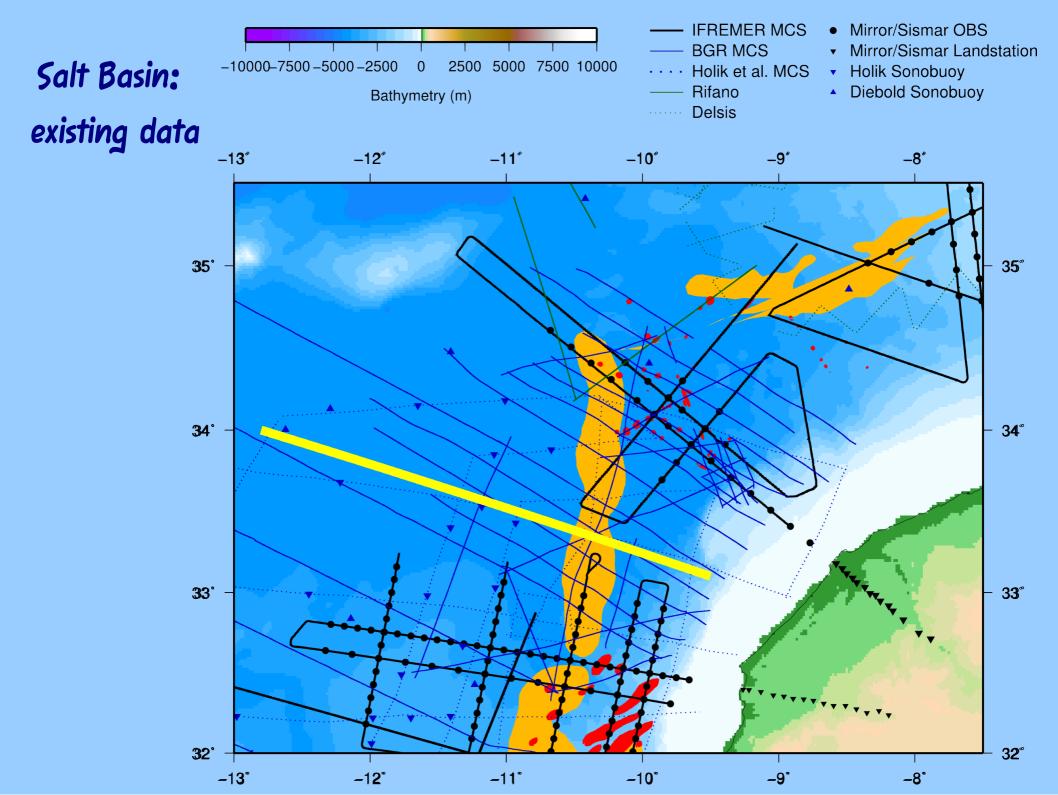


## MIRROR: Wide-angle seismic profiles

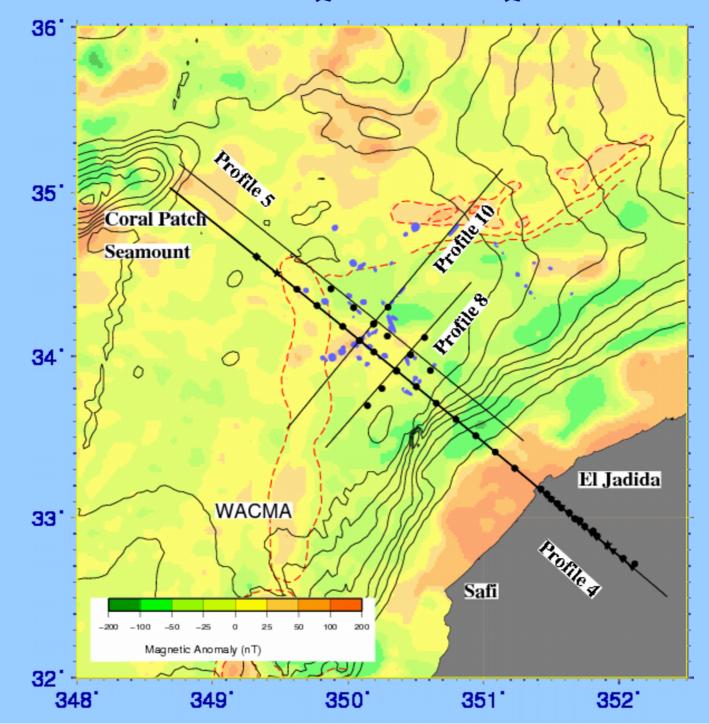


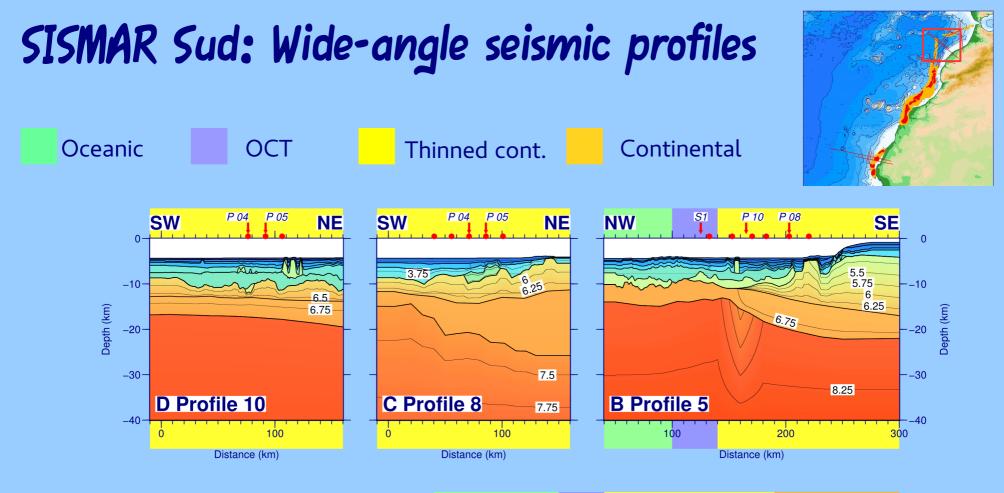


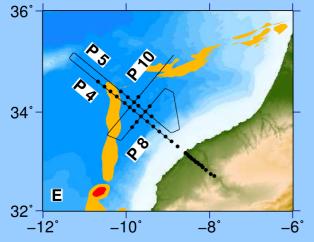


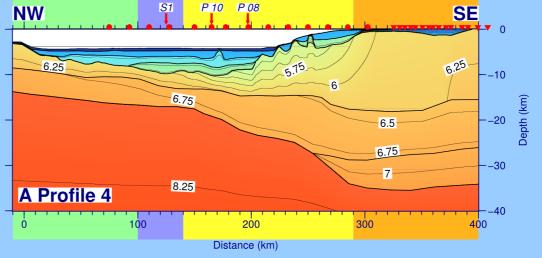


#### SISMAR: magnetic anomaly

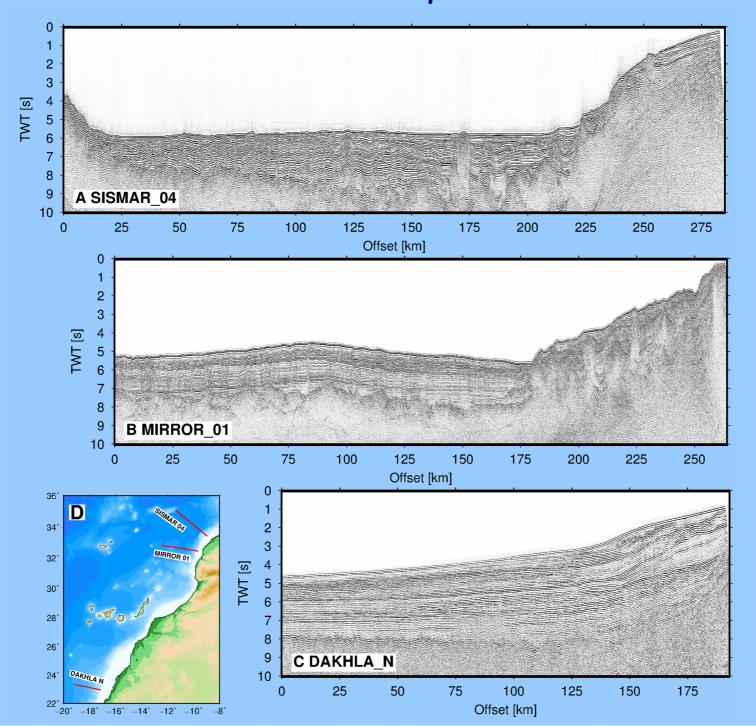




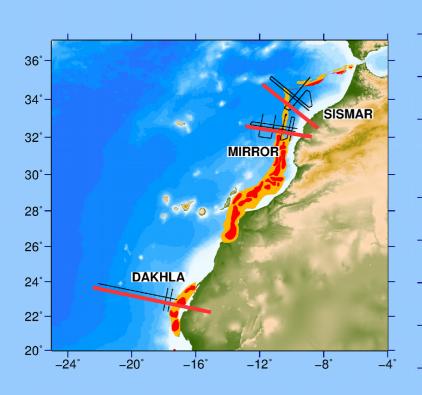


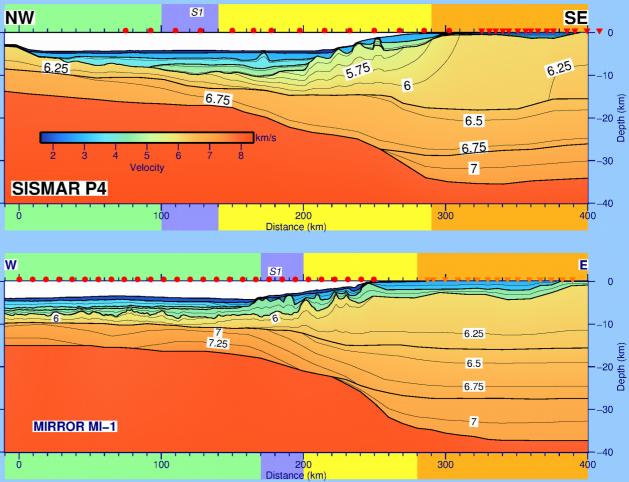


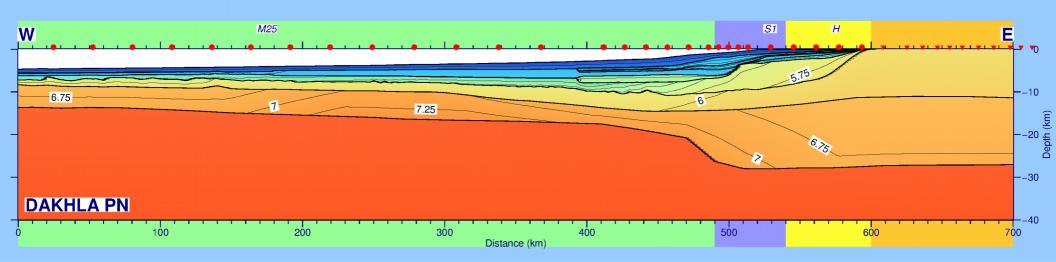
#### Dakhla, Mirror and Sismar MCS profiles



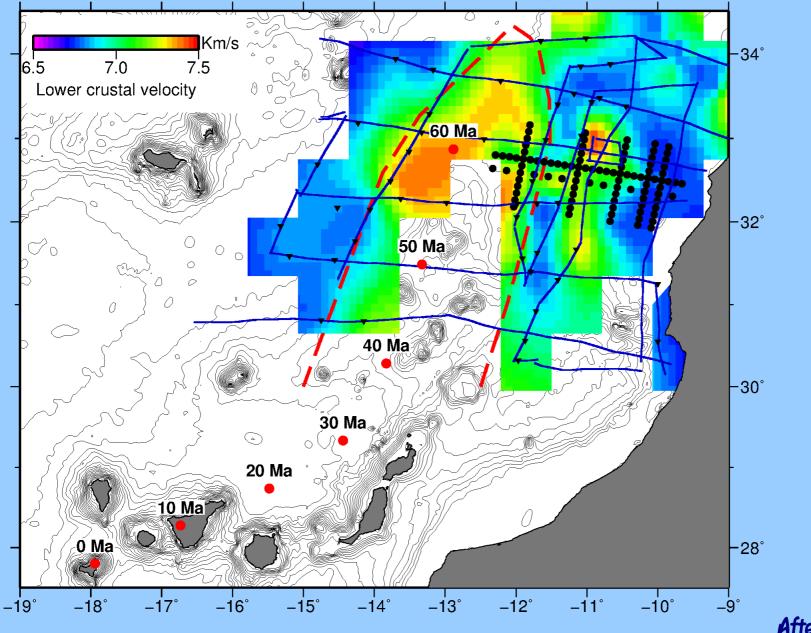
Comparison





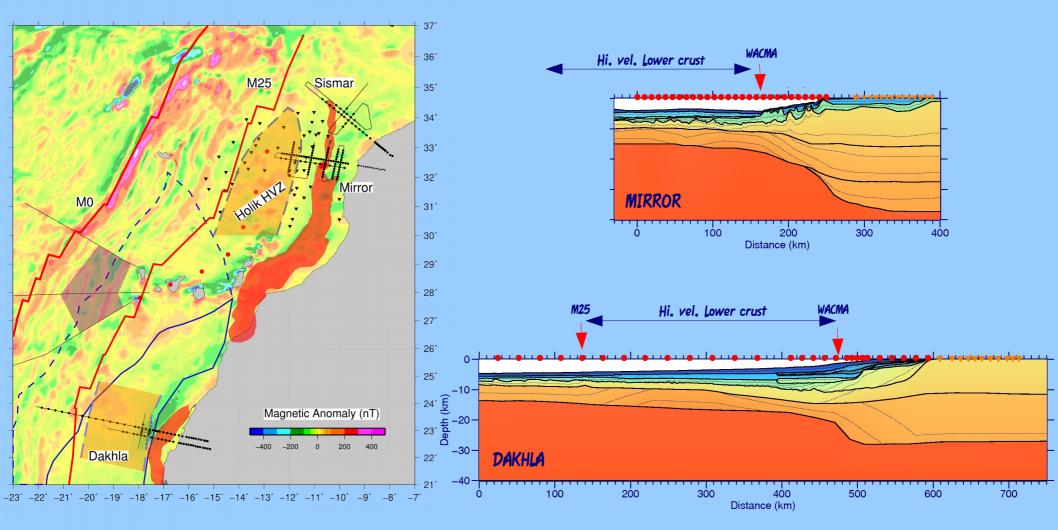


### Lower crustal velocity



After Holik, 1991

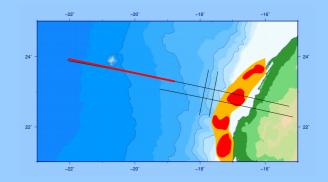
#### Magnetic anomalies

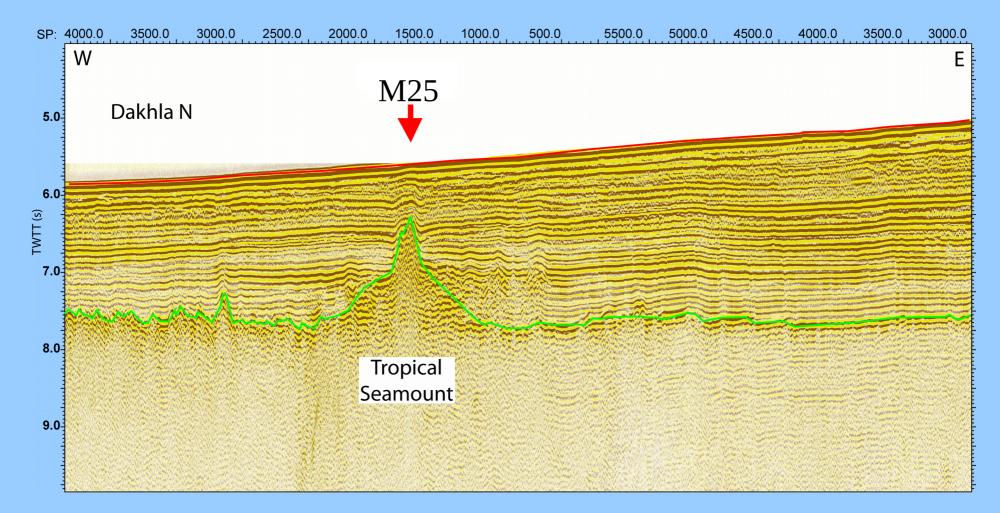


Verhoef et al., 1986

#### The southern Morrocan margin

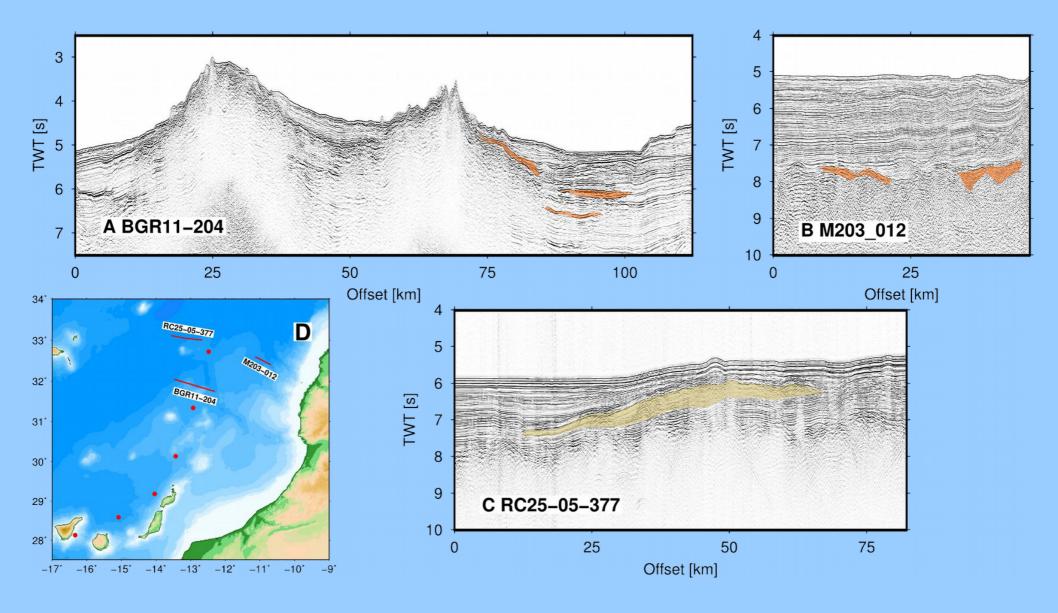
Seabed Morphology of the offshore Dakhla (off Western Sahara)



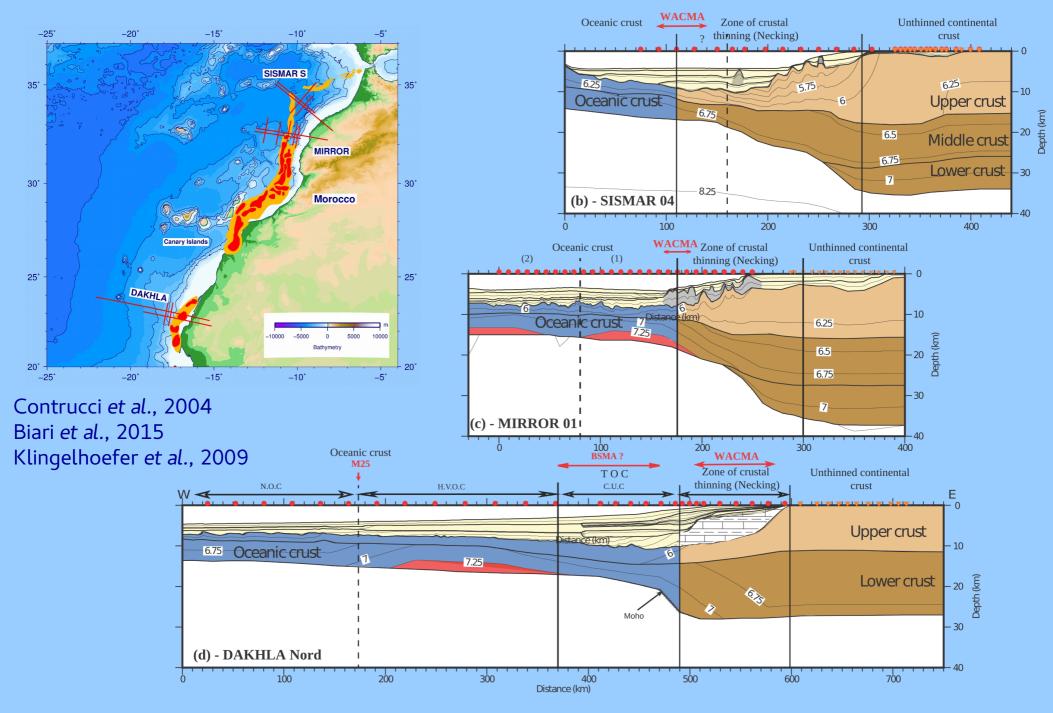


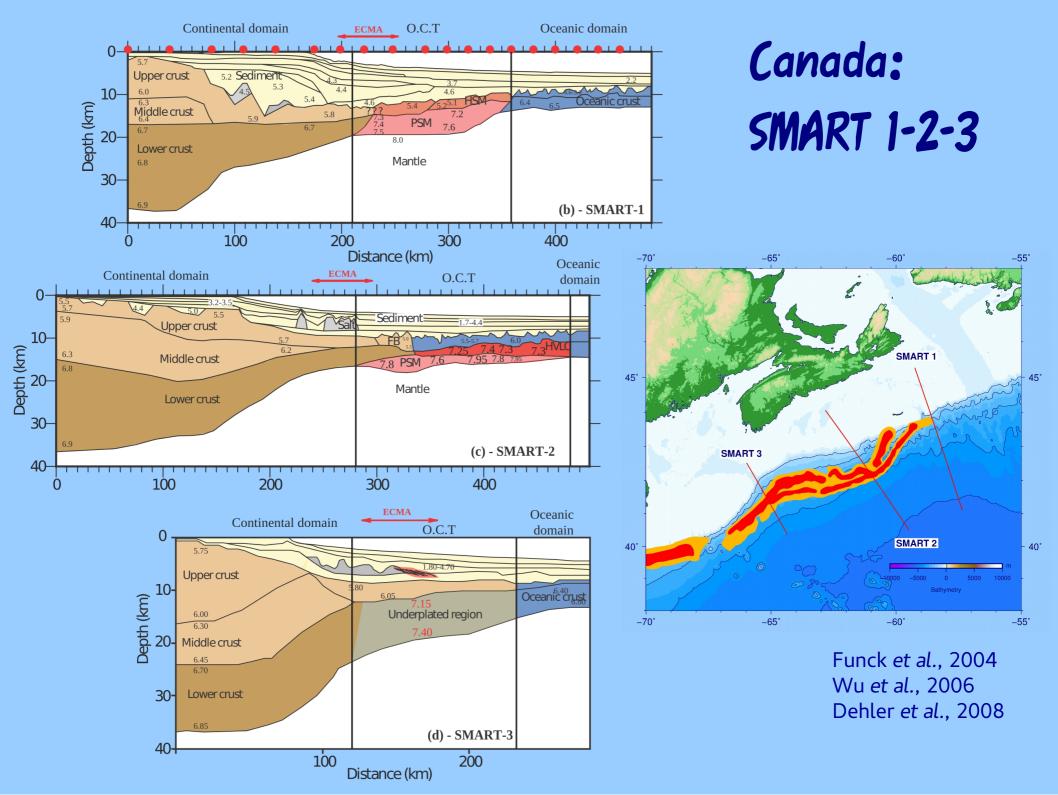
#### Volcanism along the margin

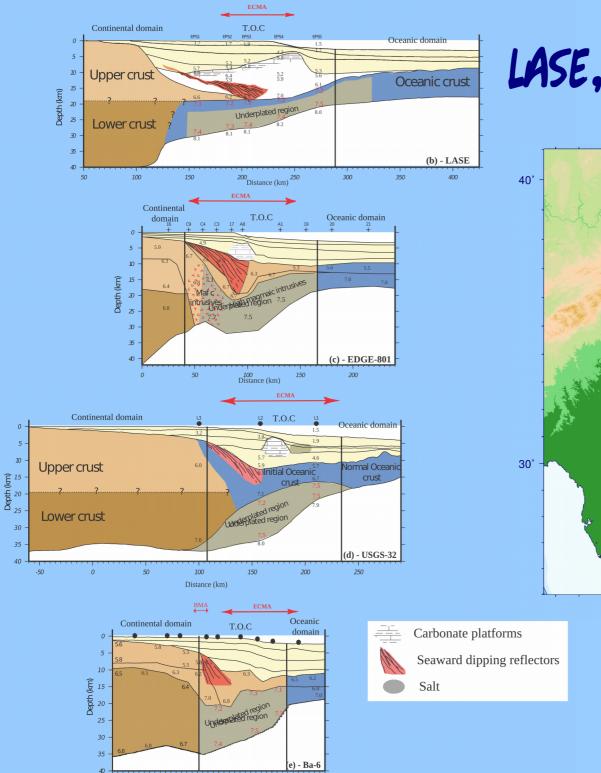
- Small seamounts and dykes in the southern part of the study region
- Mostly related to the Canary Hotspot



### Morocco: SISMAR, MIRROR, DAKHLA







50

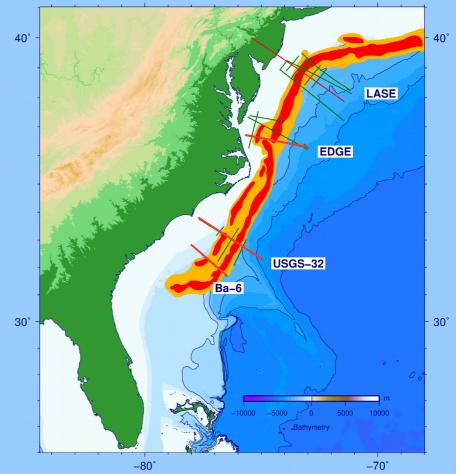
100

Distance (km)

150

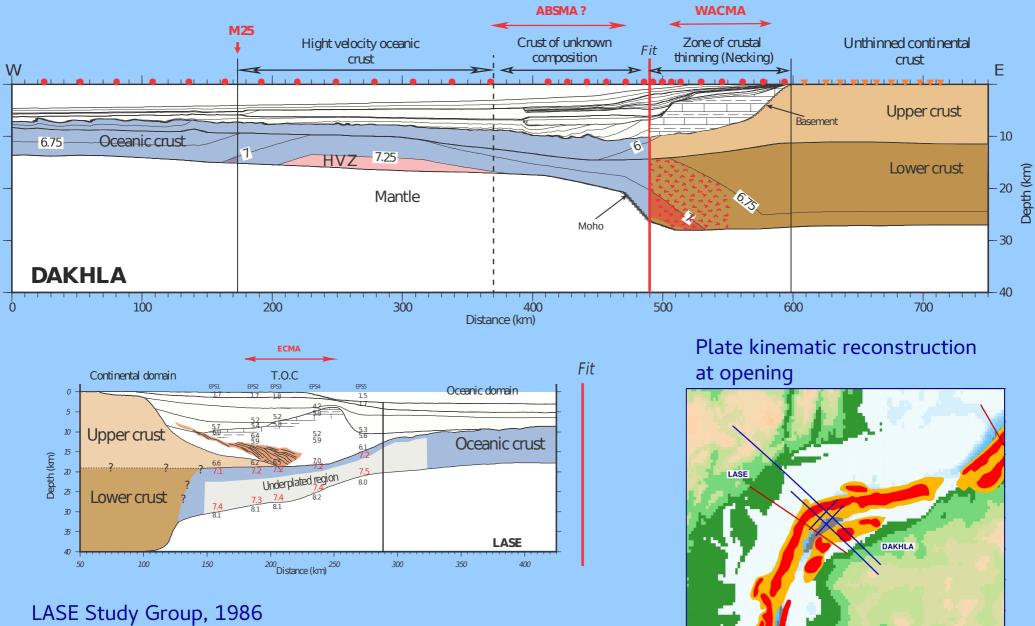
200

### LASE, EDGE, USGS-32, BA-6



LASE Study Group, 1986 Holbrook *et al.*, 1994 Tréhu *et al.*, 1989 Austin *et al.*, 1990

### Dakhla and Lase



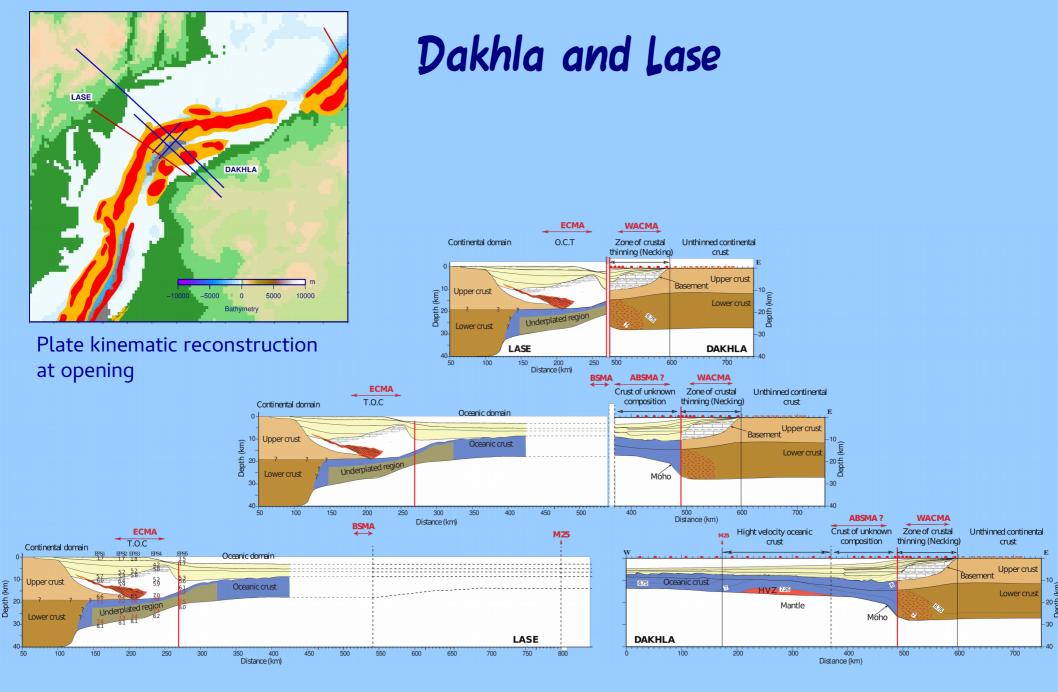
-5000 0

Bathymetry

5000

10000

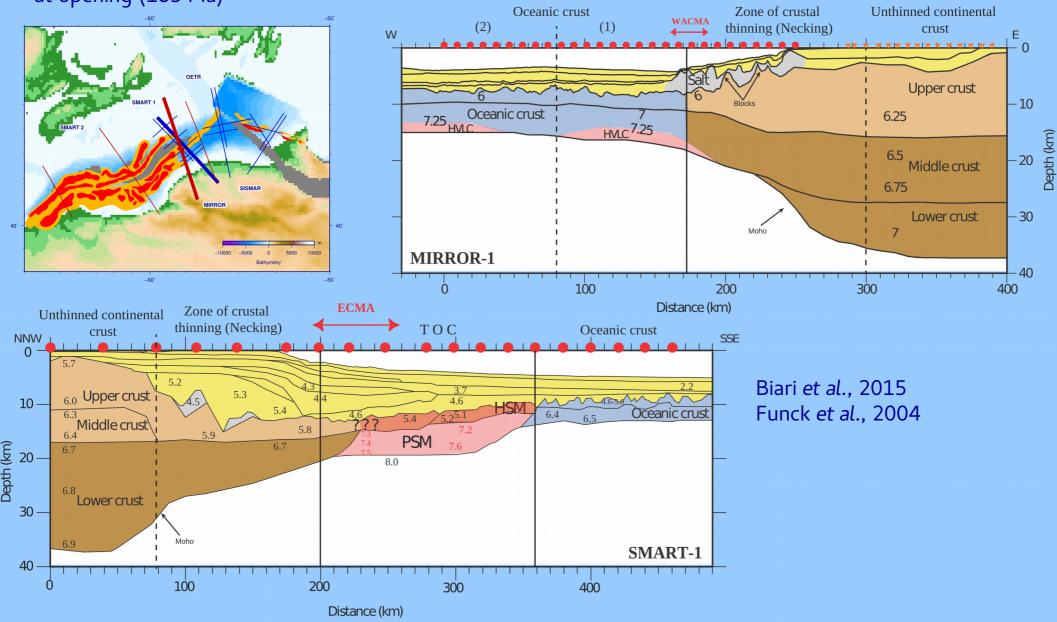
Klingelhoefer et al., 2009

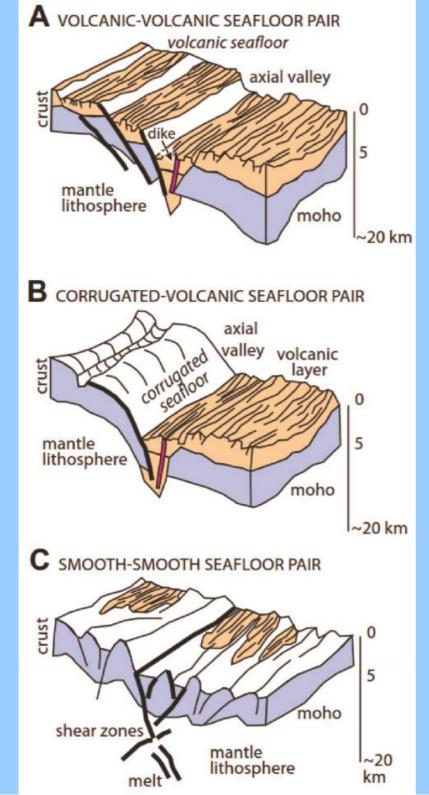


LASE Study Group, 1986 Klingelhoefer *et al.*, 2009

### MIRROR-1 and SMART-1

Plate kinematic reconstruction at opening (185 Ma)





#### Modes of seafloor generation at a melt-poor ultraslow-spreading ridge

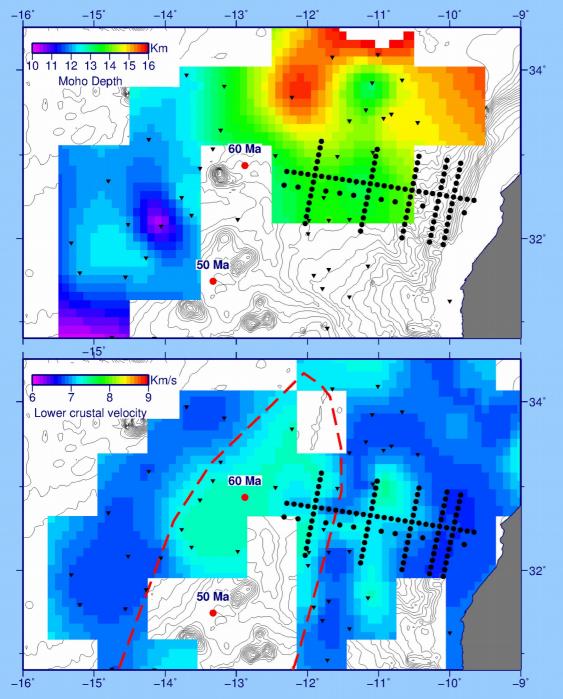
Mantle derived rocks can be continuously exhumed for at least 11 million years.

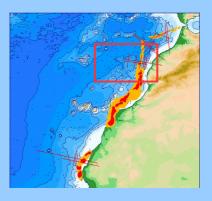
These mantle derived serpentinites form broad patches of smooth seafloor, several tens of kilometers across

The crust at Smart 1 and Mirror 1 profiles could have been generated by asymmetric spreading.

Cannat *et al.*, 2006

### Moho depth and lower crustal velocity





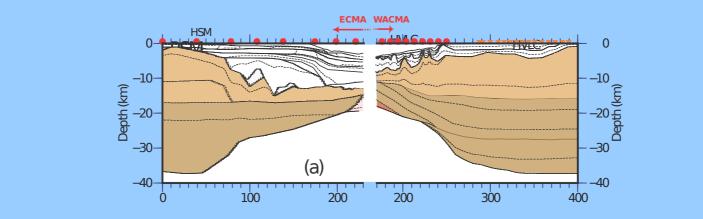
Crustal thickness is higher in the region of the passage of the Canary Hotspot

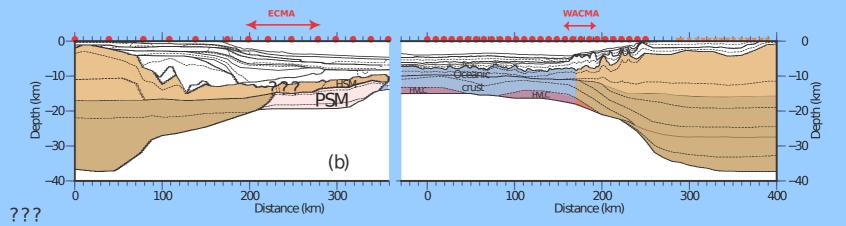
This region is located inside WACMA and the M25 magnetic anomalies

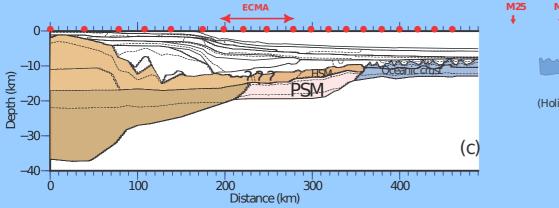
Seaward of M25 the crustal thickness decreases to 4-5 km.

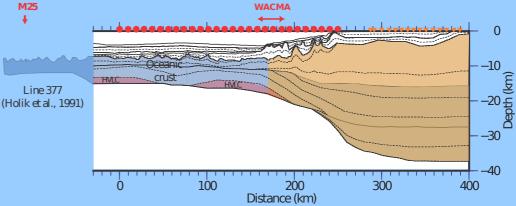
After Holik et al., 1991

### **Opening of the Central Atlantic**



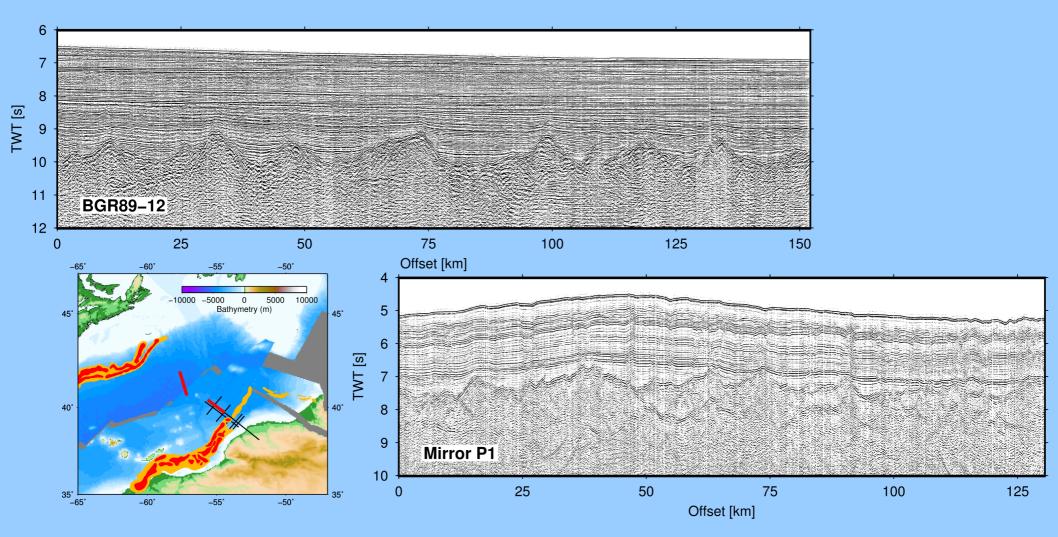


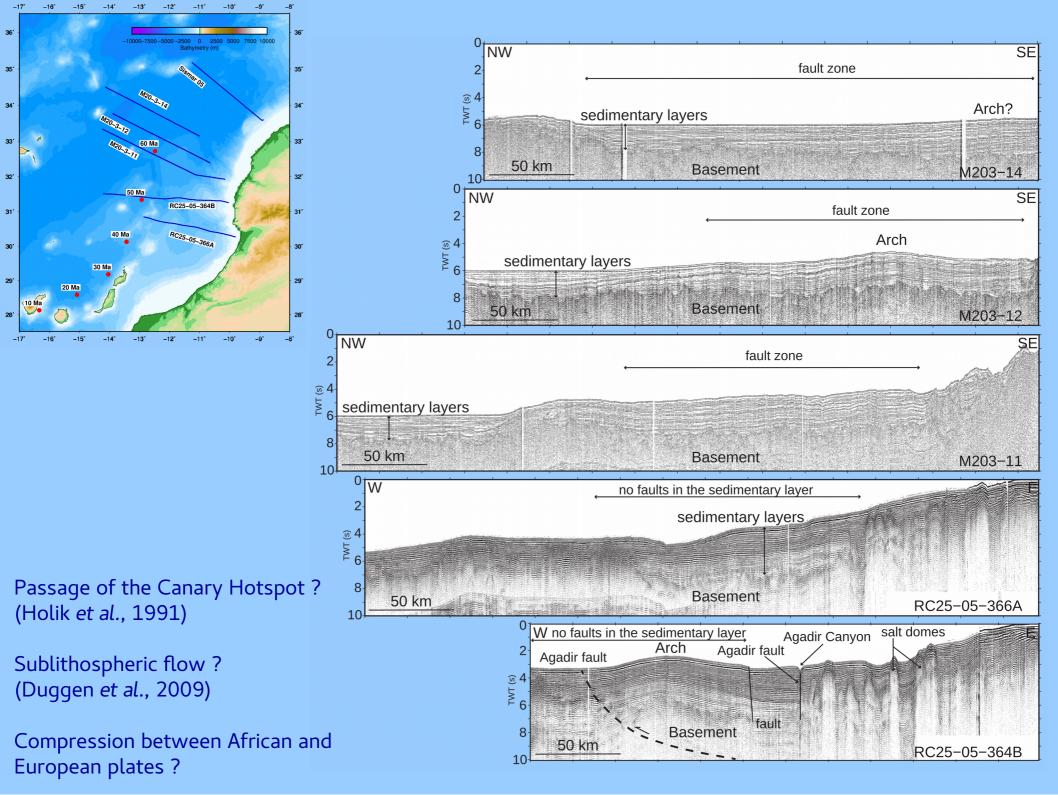




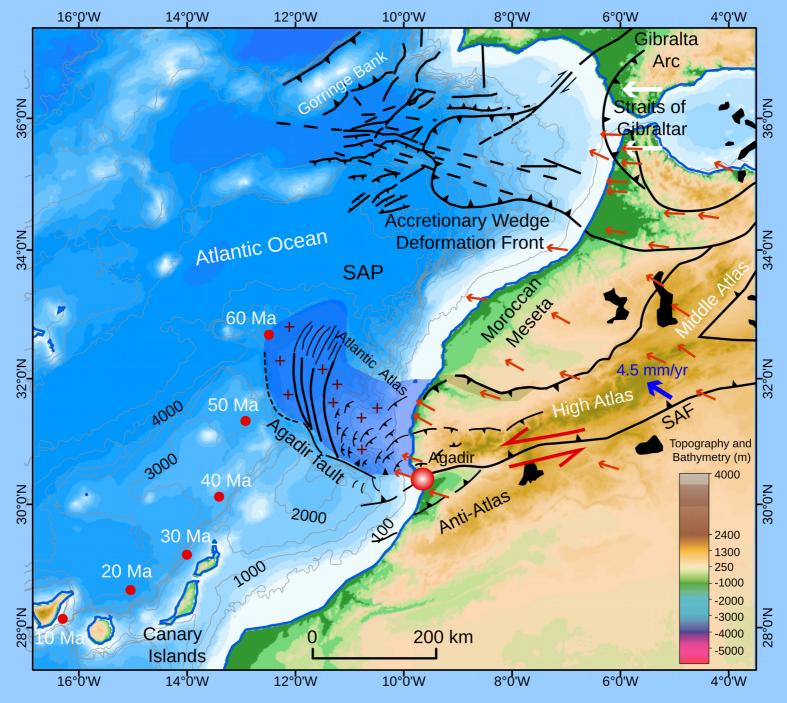
Biari et al., submitted

### Sedimentary basins





#### Atlantic Atlas

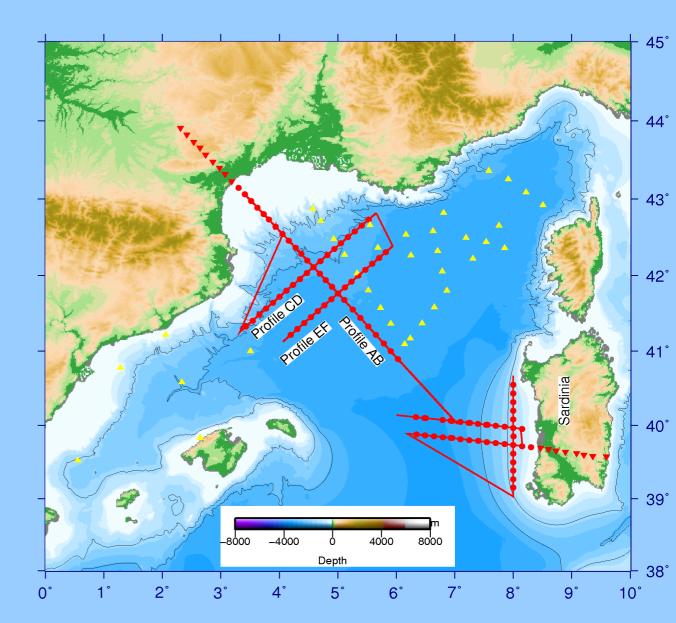


Benabdellouahed et al., submitted



- Young margin pair
- Opening as back-arc basin in Oligocene
- Second phase from rotation of the Corsica Sardinia Block.
- -> good constraints on reconstruction

#### Sardinia: the Liguro-Provencal Basin



### Reconstruction at opening

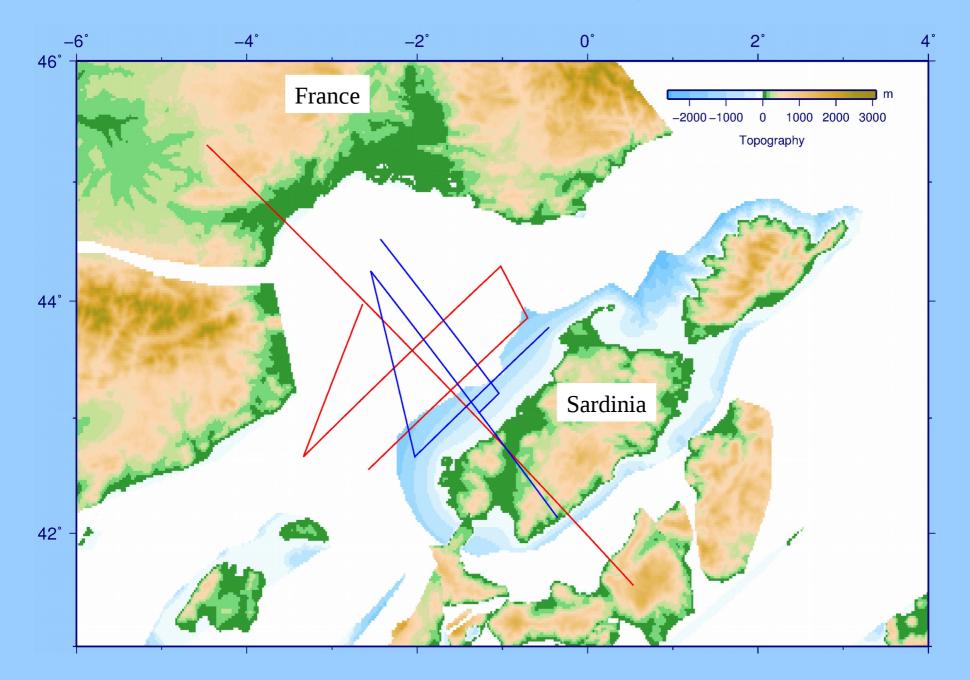
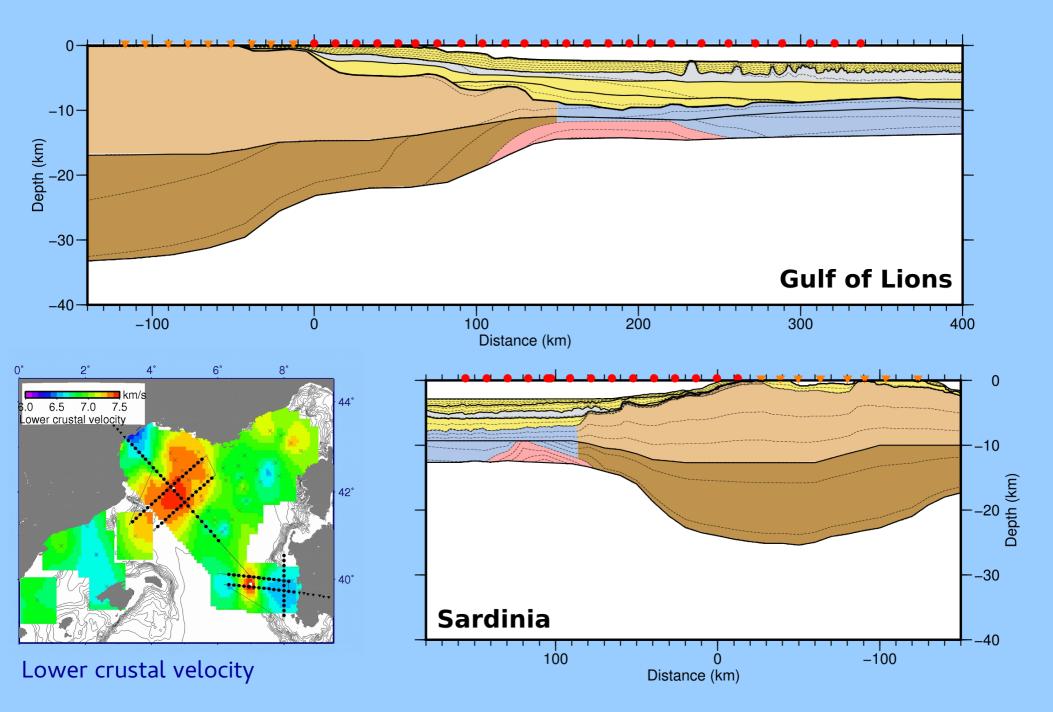
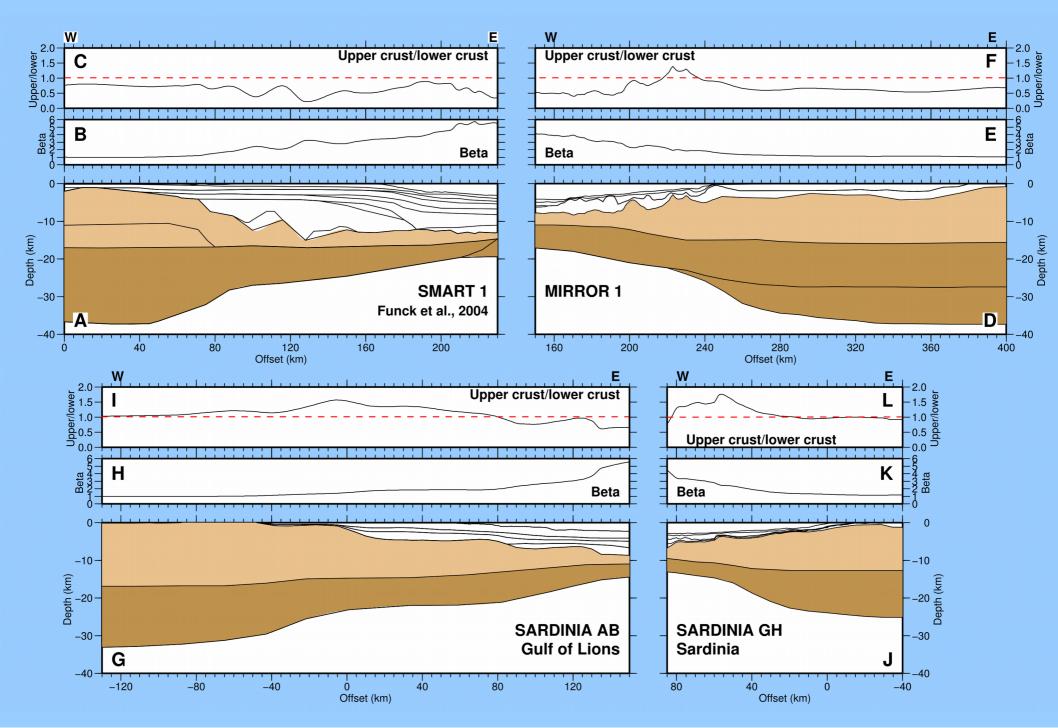


Plate cinematic reconstruction at 25 Ma, Poles from: Van Hinsbergen et al., 2014

### Sardinia margin pair



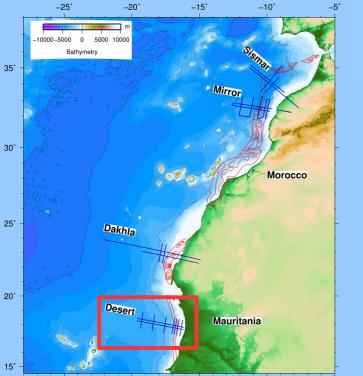
### Upper/lower crust



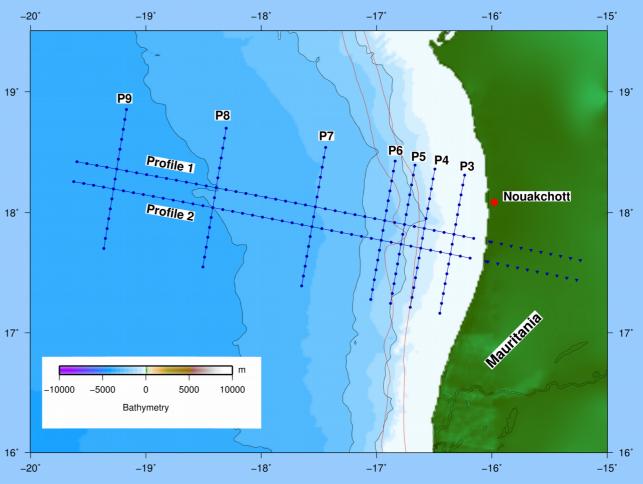
#### Conclusions

- The central Atlantic opened symmetrically, but first seafloor spreading produced mainly serpentinised upper mantle material on the Canadian and untypical oceanic crust on the Moroccan side.
- More volcanic products can be imaged along the western than on the eastern part of the margin, which might explain the higher amplitude of ECMA as compared to WACMA
- Sedimentation on the Canadian side was mainly undisturbed, however the compression between the African and European plates and/or the passage of the Canary Island Hotspot created fault structures and updoming of the seafloor at the Moroccan margin.
- Comparison with the Sardinia-Gulf of Lions margin pair shows a more asymmetric distribution of zones of high lower crustal velocities, which might be linked to the presence of a cold downgoing slab at the Sardinian side.



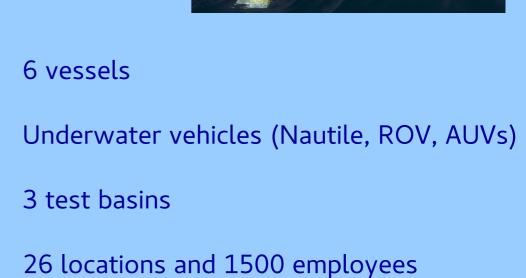


#### DEep SEismic Refraction Transect of Mauritania



The French Research Institute for Marine Studies Exhibition: Hall B 2412

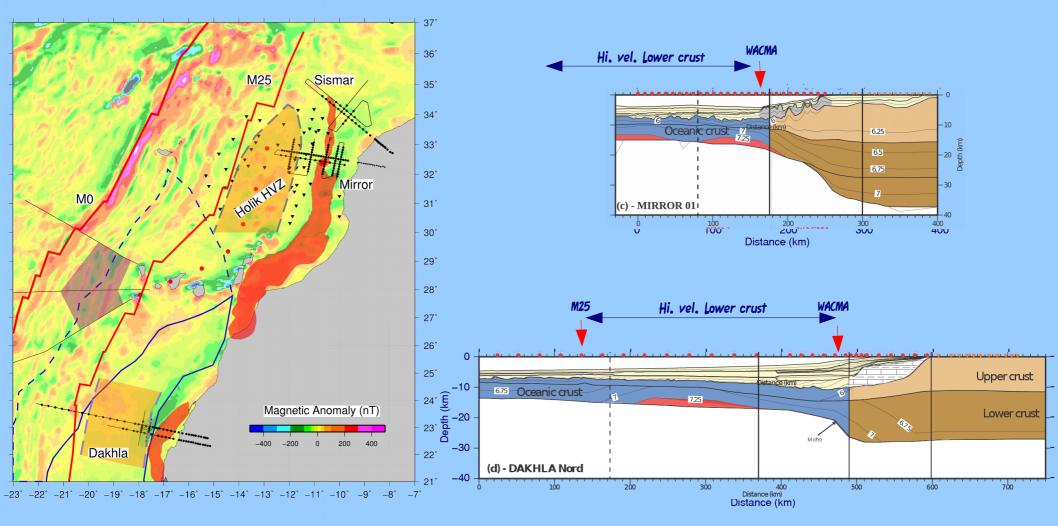
- Geophysical & geodynamic studies
- Geotechnical studies (oil-gas, geohazard)
- Geochemistry, metallogeny
- Biochemistry, biotechnology





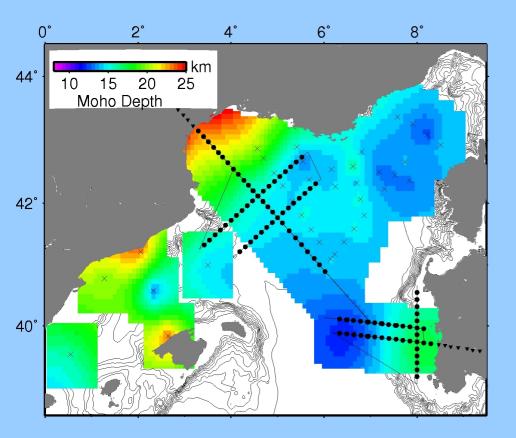


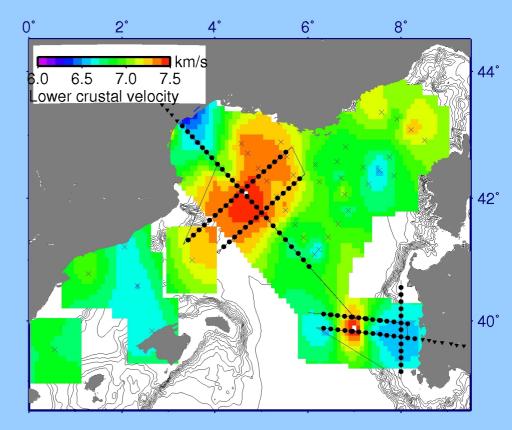
#### Magnetic anomalies



Verhoef et al., 1986

### Moho depth and lower crustal velocity





#### Moho depth

Lower crustal velocity