Full waveform inversion of OBS data for gas hydrate exploration in SW Taiwan

Kun-Sung Li and How-Wei Chen

Outline

- Introduction and Motivation
- Theoretical Background and Verification
- Preprocessing of Marine Seismic Data
- Application of Marine Seismic Survey
- Conclusions and Discussion

Pre-stack vs. Post-stack

Standard Processing	Proposed Pre-stack Processing
 Demultiplex Preprocessing Elevation Static Correction Common Mid-point Sorting Velocity Analysis Residual Statics Correction Normal Moveout Correction Dip Moveout Correction Stacking Post-stack Migration 	 Demultiplex Preprocessing Velocity Analysis Pre-stack Depth Migration Stacking

Literature Reviews

- Claebout (1976) suggested adjoint-state method for seismic data migration and imaging
- Lailly (1983) and Tarantola (1984)
 - Adjoint-state Method to Calculate Directional Gradient
 - Steepest Descent Method
- Tarantola (1984) Suggests Reverse-Time Imaging and Velocity Reconstruction of Waveform Residuals
- Mora (1987, 1988), Tarantola (1986, 1987), Sun and McMechan (1988), Pratt (1990), Pratt et al. (1998), Sirgue and Pratt (2004), Operto et al. (2006), Sheen et al. (2006), Shine and Ha (2008), Brossier et al. (2009), Virieux and Operto (2009), ...

Full Waveform Inversion

- FWI is a technique for seismic depth imaging, for velocity-model building, and for obtaining models of physical properties in the sub-surface at high spatial resolution.
- Prestack, Wide-aperture, Common- Shot/Receiver Seismic Data
- No Travel-time Picking and Phase Identification are Required
- Direct Wavefield Imaging of Velocity Distribution Using Either SH or Acoustic Responses
- Efficient FD Computation for Both Forward and Reverse-time Extrapolations
- No Explicit Process of Multiples, Diffractions, and Non-Physical Wavefields
- Imaging Conditions Based on Diffracted Wavefield by Adjoint-state Method

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Wavefield Extrapolation Approach 1

Pressure/Acoustic Wave Equation

$$\frac{1}{\rho^2(x,z)}\frac{\partial p(x,z)}{\partial t^2} - \rho(x,z)\nabla \cdot \left[\frac{1}{\rho(x,z)}\nabla p(x,t)\right] = s(t)$$

Where

- P : Pressure Wavefield
- v : Velocity Distribution
- *s* : Source Time Function
- ρ : density

Wavefield Extrapolation Approach 2

Velocity-Pressure Wave Equation

$$\begin{cases} \frac{\partial v_x}{\partial t} = b(x,z) \frac{\partial P(x,z)}{\partial x} + f_x \\ \frac{\partial v_z}{\partial t} = b(x,z) \frac{\partial P(x,z)}{\partial z} + f_z \\ \frac{\partial P(x,z)}{\partial t} = \kappa(x,z) \left(\frac{\partial v_x}{\partial x} + \frac{\partial v_z}{\partial z} \right) \\ \frac{\partial P(x,z)}{\partial t} = \kappa(x,z) \left(\frac{\partial v_x}{\partial x} + \frac{\partial v_z}{\partial z} \right) \end{cases}$$

Where

P(x, z) : Pressure

- **v** : Velocity
- **b** : Buoyancy
- k : Bulk modulus
- F_x , F_z : Hori. and Vert. Comp. Force

Generalized Nonlinear Inverse Problems: Solved Through Least Squares Criterion

d=Gm
$$\Delta d=d_{obs}-d_{syn}$$
 Misfit Function $S(m) = \frac{1}{2}\Delta d^{\dagger}\Delta d$
 $S(\mathbf{m}_0 + \Delta \mathbf{m}) = S(\mathbf{m}_0) + \sum_{j=1}^M \frac{\partial S(\mathbf{m}_0)}{\partial m_j} \Delta m_j + \frac{1}{2} \sum_{j=1}^M \sum_{k=1}^M \frac{\partial^2 C(\mathbf{m}_0)}{\partial m_j \partial m_k} \Delta m_j \Delta m_k + O(m^3)$
 $\Delta m = -\left[\frac{\partial^2 S(\mathbf{m}_0)}{\partial m^2}\right]^{-1} \frac{\partial S(m_0)}{\partial m} = H^{-1}(m)\gamma(m)$
 $H^{-1}(m) = \left[\frac{\partial^2 S(\mathbf{m}_0)}{\partial m^2}\right]^{-1} \qquad \gamma(m) = -\frac{\partial S(m_0)}{\partial m} \checkmark$ Frechét Derivative

$$\gamma(m) = \sum_{s} \int \dot{p}_{cal}(m) \dot{p}'_{res}(m) dt$$

Using Cross-correlation to Estimate Frechét Derivative

(Tarantola, 1984)

Optimization

Conjugate Gradient Method

 $\mathbf{m}^{(n)} = \mathbf{m}^{(n-1)} - \alpha^{(n)} \phi^{(n)}(\mathbf{m})$ $\phi^{(n)}(\mathbf{m}) = \gamma^{(n)}(\mathbf{m}) + \beta^{(n)} \phi^{(n-1)}(\mathbf{m})$

$$\alpha^{(n)} = \frac{[F_n \phi^{(n)}(\mathbf{m})]^t \Delta \mathbf{d}}{[F_n \phi^{(n)}(\mathbf{m})]^t [F_n \phi^{(n)}(\mathbf{m})]} \qquad F_n \phi^n(\mathbf{m}) = \lim_{\varepsilon \to 0} \frac{\phi(\mathbf{m} + \varepsilon \mathbf{m}) - \phi(\mathbf{m})}{\varepsilon}$$
$$\beta^{(n)} = \frac{[\gamma^{(n)}(\mathbf{m}) - \gamma^{(n-1)}(\mathbf{m})]^t [\gamma^{(n)}(\mathbf{m})]}{[\gamma^{(n-1)}(\mathbf{m})]^t [\gamma^{(n-1)}(\mathbf{m})]} \qquad \frac{\text{Restarting, Polak and Ribiere (1969)}}{\beta^{(n)} = \max\{\beta^{(n)}, 0\}}$$

Steepest Descent Method

$$\mathbf{m}^{(n)} = \mathbf{m}^{(n-1)} - \alpha^{(n)} \gamma^{(n)}(\mathbf{m})$$

Regularization and Precondition

Total Variation Regularization

$$V(y) = \sum_{i,j} \sqrt{|y_{i+1,j} - y_{i,j}|^2 + |y_{i,j+1} - y_{i,j}|^2}$$

- Precondition
 - $W(x,z) = (\rho(x,z)v^3(x,z))^{-1}$

 - Weighting for Correct Self-Adjoint

Multi-scale Consideration Through Frequency-domain FWI Approach



CPU-efficient Frequency-domain FWI



Fresnel Zone/FK Resolution Analysis

Wavenumber Illumination

$$k = \frac{2f}{c} \cos\left(\frac{\theta}{2}\right) n$$

Depends on Aperture Angle θ and Angular Frequency f



One frequency and one aperture in the data space map one wavenumber in the model space. Therefore, frequencies and apertures have a redundant control on the wavenumber coverage. 14



Ps

Problems in Cycle-skipping

Solid line: « Recorded » monochromatic seismogram.

Dash lines: « Computed » monochromatic seismograms.

Top: the error exceeds half a period. Computed cycle *n*+1 is matched with cycle *n* leading to cycle skipping.

Bottom: the error is smaller than half a period. The recorded and modeled cycles *n* are matched.



Inversion Scheme



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Frechét Derivative



Forward Extrapolation → Predicted Seismograms Backward Extrapolation → Time Reversed Propagation of Residual Seismograms

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Reciprocity Principal and Self-adjoint

Shot: E: Explosive Fx, Fz: Hori. and Vert. Forces Receiver: P: Pressure Vx, Vz: Hori. and Vert. Velocities

Illumination and Resolution Analyses For Marine Seismic Data Acquisition

Velocity Model + HVZ

Red Arrow: HVZ, 500 m/s Anomaly

Sensitivity Kernel of MGL0905-08

Sensitivity Kernel of MCS881-

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Pre-processing for FWI

- Trace Editing
- Resample
- Frequency Analysis
- Band-pass Filter on Frequency Domain
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- Spiking Decon.
- Mute Direct Wave

Noise Filtering and Suppression

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Stacking Velocity Analysis: Semblance

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Bottom Simulating Reflector (BSR)

Two Initial Velocity Models

FC: Formosa Canyon; DF: Deformation Front; PC: Penghu Canyon; KC: Kaoping Canyon

Dominant Frequency: 7 Hz

2009 3D OBS Survey, MCS881-42

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MCS881-42 FWI and Interpretation

Velocity of BSR: 1.7 km/s; FG:1.4 km/s Red: seafloor; Black: unconformity; Blue: BSR; Green: Fault

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Computation Cost for FWI

- MS881-42 Velocity Model
 - Dimension: 20 km * 7 km(561 grid * 1601 grid)
- **TDFWI**
 - Number of time samples: 5000; dt=0.002 sec
 - 8 OBSs, Need Total Memory of 126.4 GB
 - 8 Cores with 15.8 GB/core
 - Need Disk Storage of 4.9 TB
 - 3.77 hr/iter
- **FDFWI**
 - 10 Frequencies (3~12 Hz)
 - 8 OBSs, Need Total Memory of 37.57 GB
 - 16 Cores with 2.35 GB/core
 - Need Disk Storage of 10 GB
 - 0.67 hr/iter

Conclusions and Discussions

- Theoretical Development, Validation and Implementation for Field Data are Presented
- FWI is successfully applied to synthetic and practical examples.
- The Inversion is Well Suited for Parallel Inversion over Cluster Computing Environment.
- The image of Frechét kernels provide a unique opportunity in understanding the characteristics and limitations of reversetime inversion and imaging strategy.
- The illumination analysis in terms of resolution for the targeted zone becomes important even during the survey planning stage.
- We apply spatial reciprocity relationship between vertical particle velocity wavefields generated by explosions (vertical geophone of an OBS) and pressure wavefields generated by a vertical force (air gun) for inversion strategy.

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Conclusions and Discussions

- Multi-frequency or Multi-scale Strategy can Provide Better Resolution and Result with FWI
- FWI: Mitigation of Secondary Minima through Data Manipulation
 - Frequency hierarchy: from low to high frequencies (decimation)
 - Time windowing: from first-arrival phases to later phases
 - Space windowing: from low wavenumbers to high wavenumbers
- Practical Application of FWI to Seismic Line mcs881-42
 - Velocity of BSR: 1700 m/s
 - Depth of BSR: 200-500 m
 - Velocity of FG: 1400 m/s
 - Resolution: 58-333 m(Velocity Range:1400-2000 m/s)
- Geology and Tectonic Interpretation of mcs881-42
 - Depth of BSR, Formation Boundary and Unconformity
 - Low/High Velocity Zones for Gas-hydrate
 - Identification of Faults and Sediments
- Substitution of BSR and FG
 Boundary States of States and FG
- Resolution of FWI Depends on the Source Wavelet, Dominant Frequency, Bandwidth and Shot/Receiver Interval

Split-step Plane Wave Waveform Inversion

Kun-Sung Li, How-Wei Chen and Paul Stoffa

Marmousi Velocity Model

nx=384, nz=122, dx=dz=0.02 km, dt=0.004, nt=800

Initial Velocity Model

nx=384, nz=122, dx=dz=0.02 km, dt=0.004, nt=800

Marmousi (5 Ray Parameters) and Initial Models' Seismograms

Marmousi Model vs Updated Model

Velocity Difference

Thank You for Your Attention.

The Future Development of FWI

- Survey planning and Instrumentation
- Non-linear Algorithm and Optimization Methods
- Multi-componment, Multi-attribute and Multiparameter Full Wavefield Inversion

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