

Integrating onshore and offshore paleoseismic data: a Bayesian model

**Chris Goldfinger¹, Steve Galer¹, Jeffrey Beeson¹, Ann E. Morey¹,
Bran Black¹, C. Hans Nelson²**

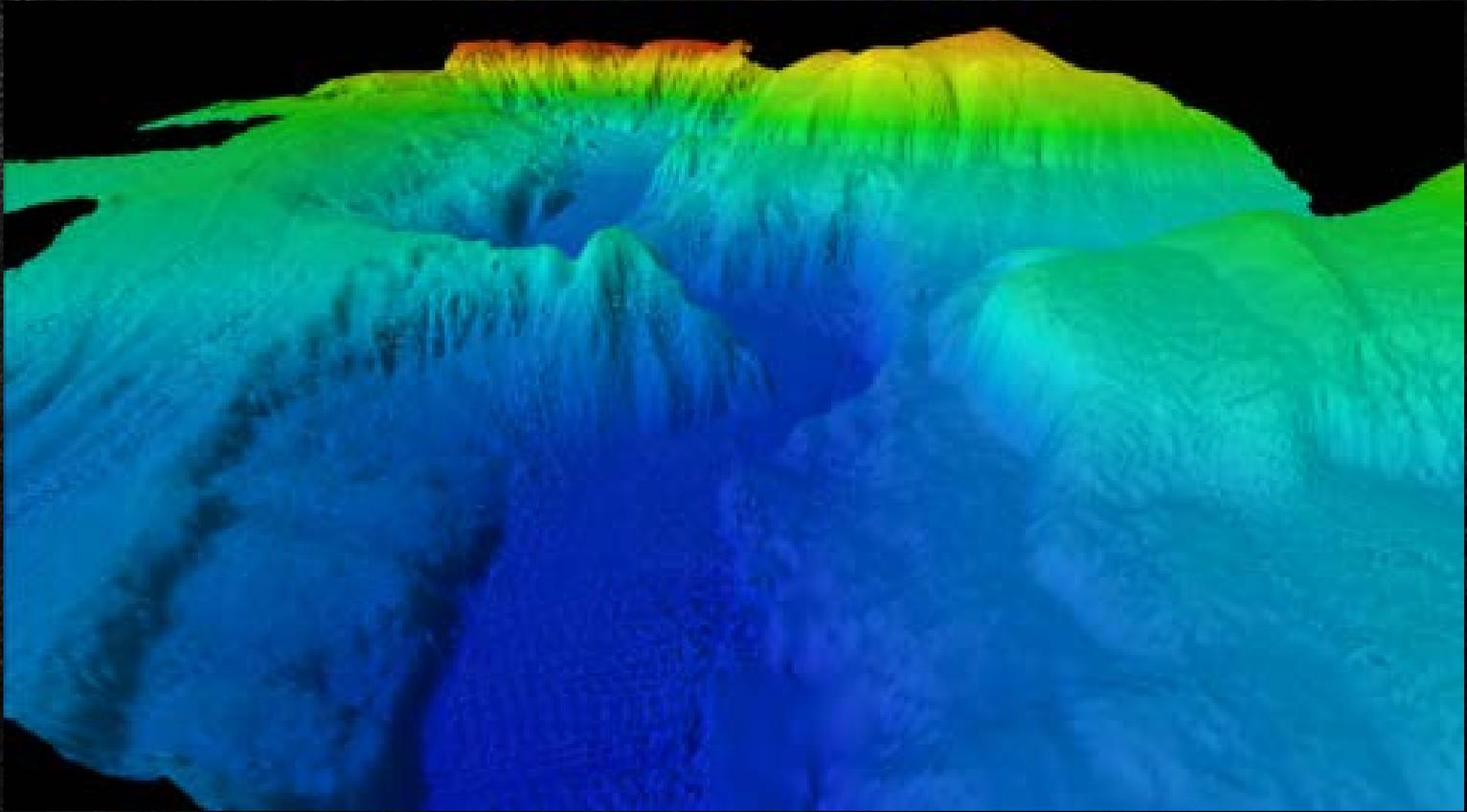
¹*Oregon State University, College of Earth, Ocean and Atmospheric Sciences, 104 Ocean Admin. Bldg., Corvallis OR 97331, USA.*
gold@coas.oregonstate.edu

²*Instituto Andaluz de Ciencias de la Tierra (IACT) CSIC-Univ. de Granada Campus de Fuentenueva s/n 18002 Granada, Spain*
odp@ugr.es





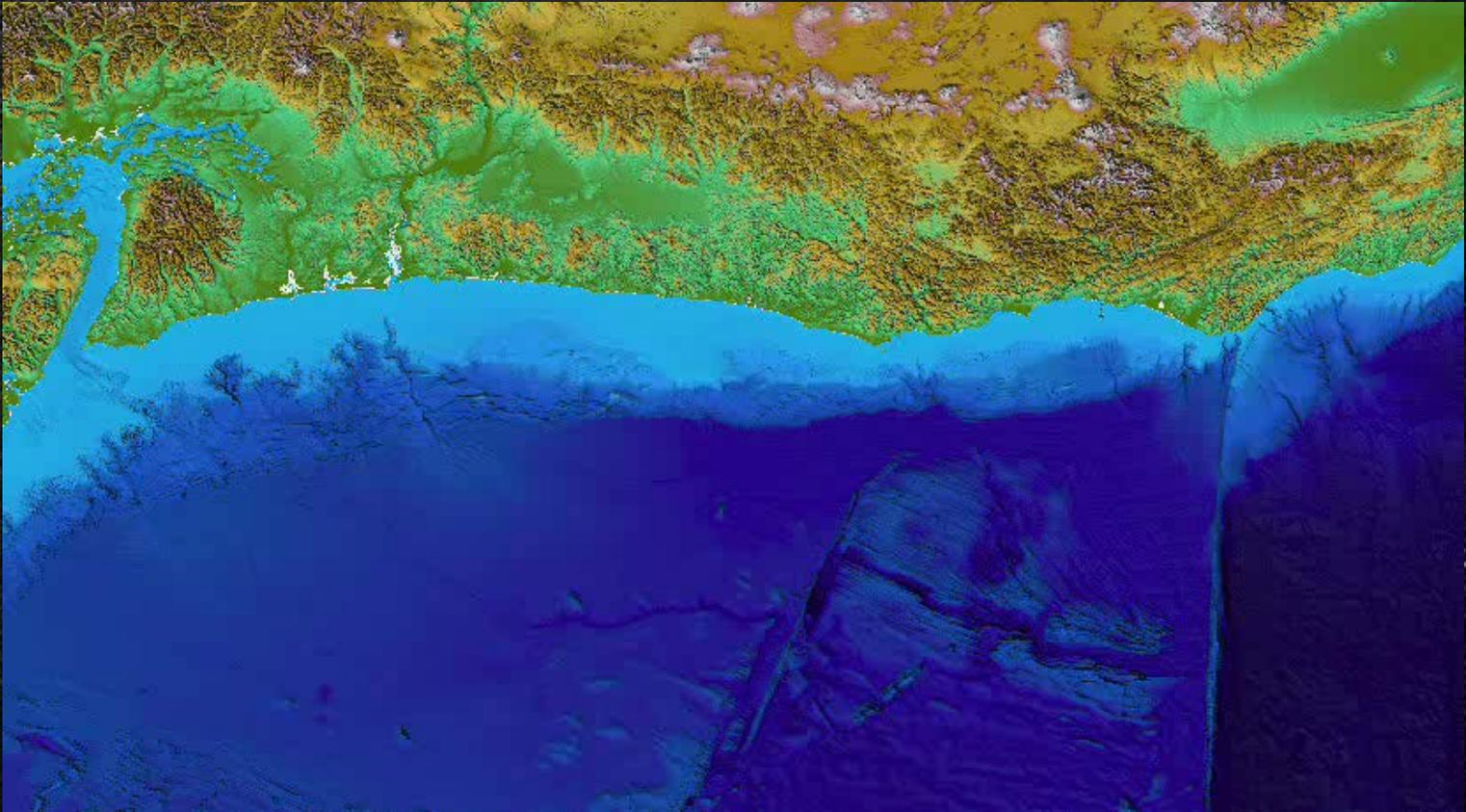
What actually happens during the earthquake?





What actually happens during the earthquake?

Synchronous turbidity currents are triggered within a few minutes of each other along the length of the margin



Turbidity Current Triggering Mechanisms

Tests of aerial extent can eliminate many possible triggering mechanisms.

Lack of land external sediment sources on the Cascadia and Sumatran outer slope and absence of cyclones near the equator eliminates storms (almost)

Extreme rarity eliminates impacts

- storm or tsunami wave loading/liquefaction
- sediment loading (trigger still needed)
- storm (hyperpycnal) discharge
- bolide impact
- great earthquakes
- crustal earthquakes
- tectonic over-steepening (trigger still needed)
- gas hydrate destabilization (trigger still needed)

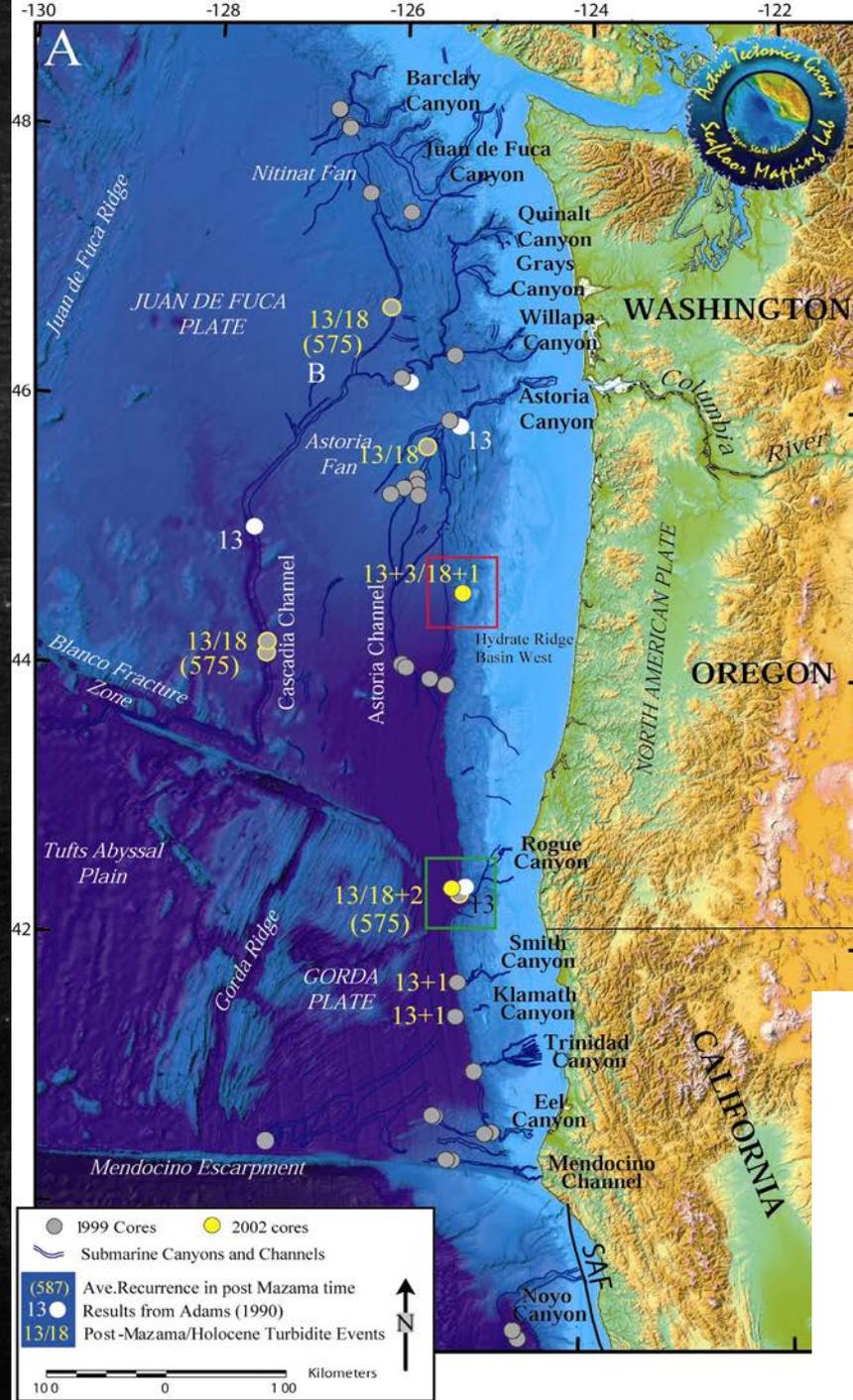
So our primary criteria for distinguishing earthquakes are

- 1) Aerial extent
- 2) Synchronicity, and
- 3) Sedimentology.

Synchronous means within a few minutes to hours at most...

^{14}C dating gets us only to within a few decades.... So how do we constrain relative timing to within a few minutes?

Cheat!



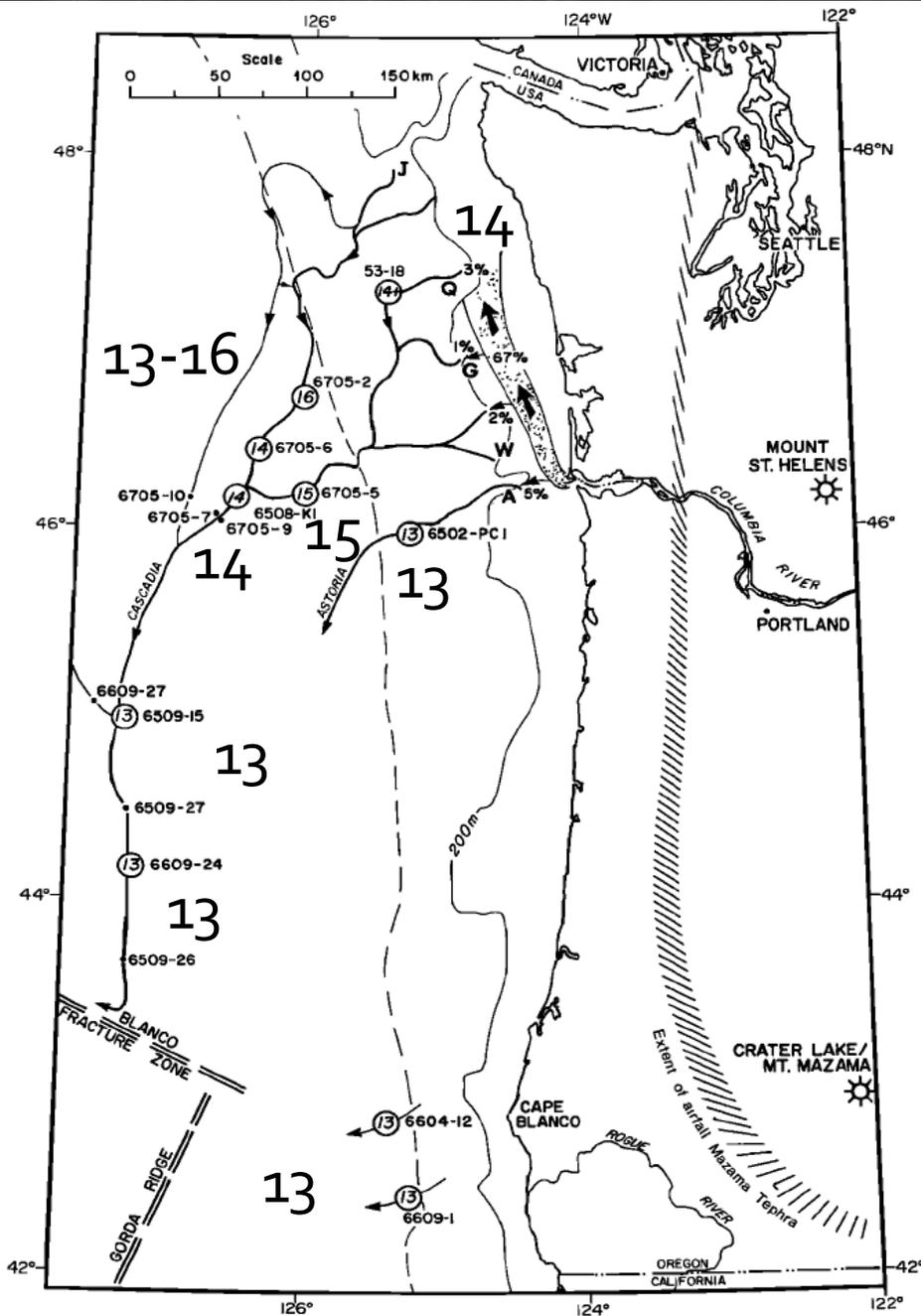
Turbidite
Paleoseismology:
Extending the
earthquake record

Cascadia Core Sites:

1999 = gray, 2002 = yellow

Older existing cores = white

Washington Channels defined by 12 days of multibeam survey, now **un-classified!**



Early tests for earthquake origin. Hats off to John Adams 1984-1990

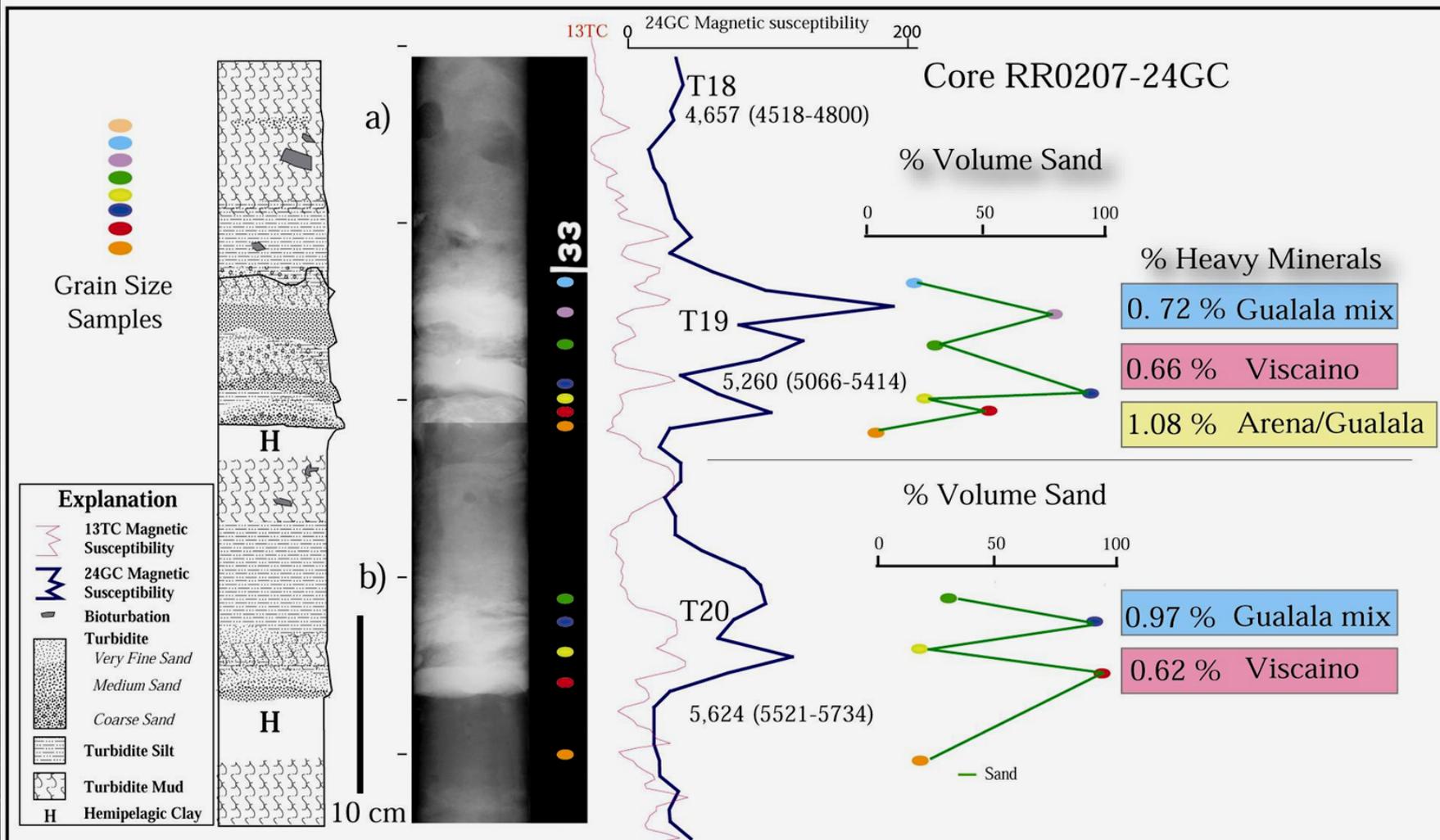
While brilliant, and this strategy has been successful, much more is needed for confidence in definition of channel systems and stratigraphy!

These cores were collected by OSU and UW investigators and students using Loran (~ .25 nm uncertainty) with PDR and occasional TRANSIT fixes.

Detailed correlations are constructed from high-resolution physical property data collected from the cores, including magnetic susceptibility (high and low), gamma density, P-wave velocity, resistivity, and CT imagery.

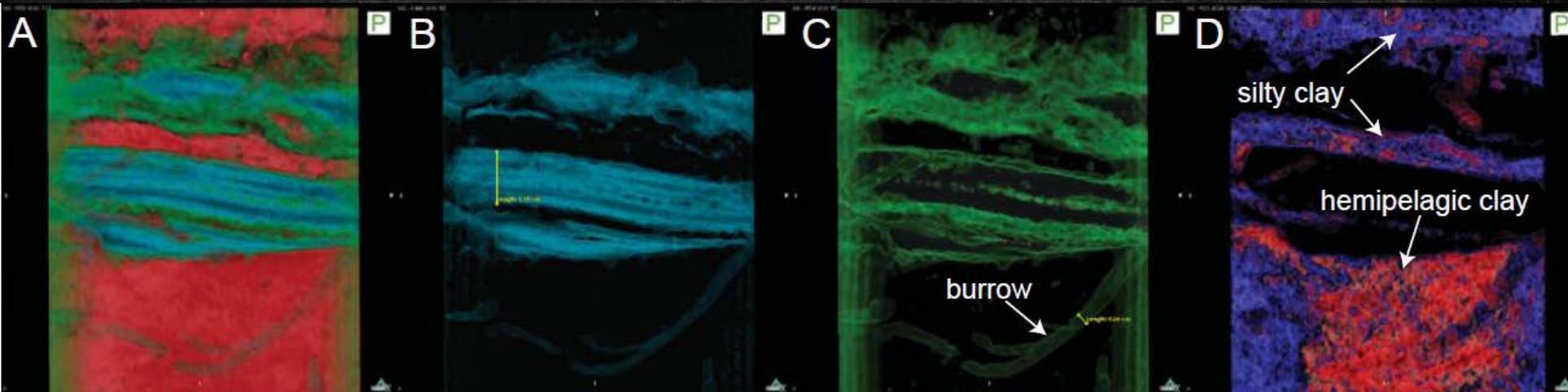


Francois Charlet



In addition to the confluence test, we correlate turbidites between remote sites to establish continuity, and test for synchronous triggering.

Correlations are made on the basis of grain-size/physical property “fingerprints” within a ^{14}C age framework



CT imagery is invaluable for understanding turbidite structure and defining stratigraphic boundaries in detail. This image breaks out the sand fraction, the silt fraction, and the hemipelagic clay by their respective CT density values.

The CT can reveal such subtle features as a worm burrow which is apparently lined with material slightly more dense than its surroundings (biogenic clay)

CT movie...

M9907-54KC

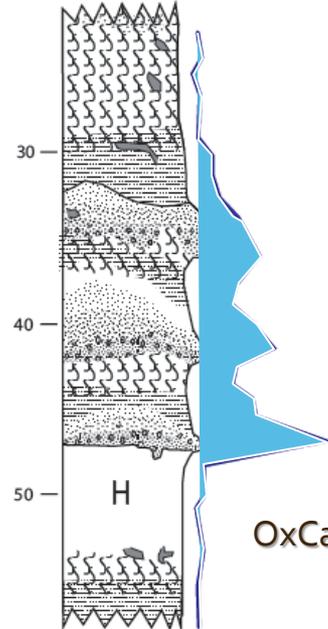
M9907-25PC



Lithology	
	Hemipelagic clay
	Turbidite silty mud
	Silt
	Very fine sand
	Sand
	Mottled clay
	Burrows

Event T3

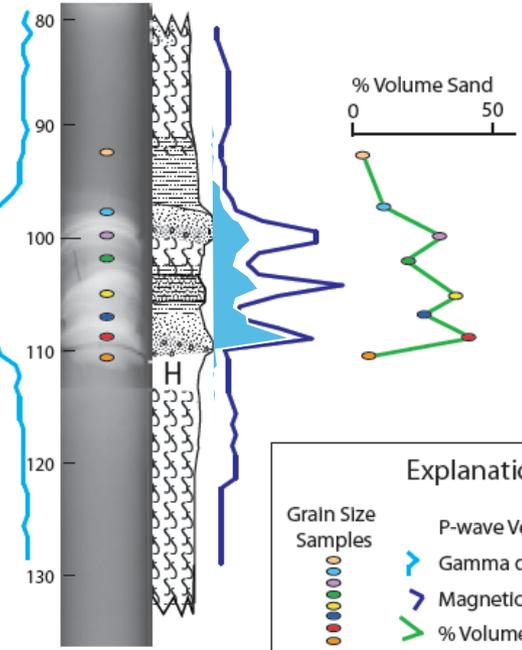
750
(680-820 cal BP)



Event T3

830
(740-920 cal BP)

OxCal modeled age: 800
(760-840 cal BP)



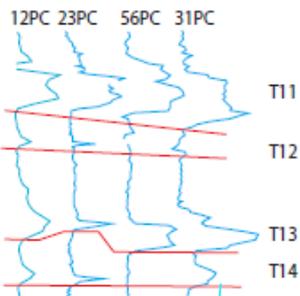
Explanation	
	Grain Size Samples
	P-wave Velocity (m/s)
	Gamma density (g/cm ³)
	Magnetic Susceptibility (SI)
% Volume Sand symbol: green line with 's' markers"/>	% Volume Sand

Looking closely, the main structure of these turbidites is a series of fining upward "pulses" (Bouma A-C) capped by a fining upward tail. The pulsing structure is commonly maintained through channel confluences, and between isolated sites as shown by this example from two cores 300 km apart, with source areas 420-500 km apart. These channels never meet.

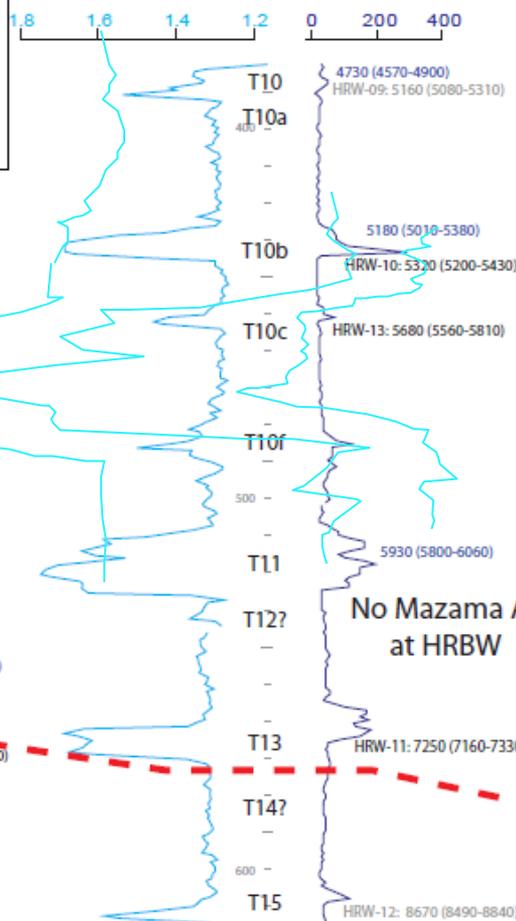
Explanation

-  Magnetic Susceptibility (SI units), high-resolution point sensor
-  Density (gm/cc)
-  First Occurrence of MA > 1.0%
-  Mazama Ash (%)
-  Radiocarbon Age, in cal BP & 2σ range, purple if erosion corrected, gray if reversed
-  Calculated hemipelagic age in cal BP & 2σ range

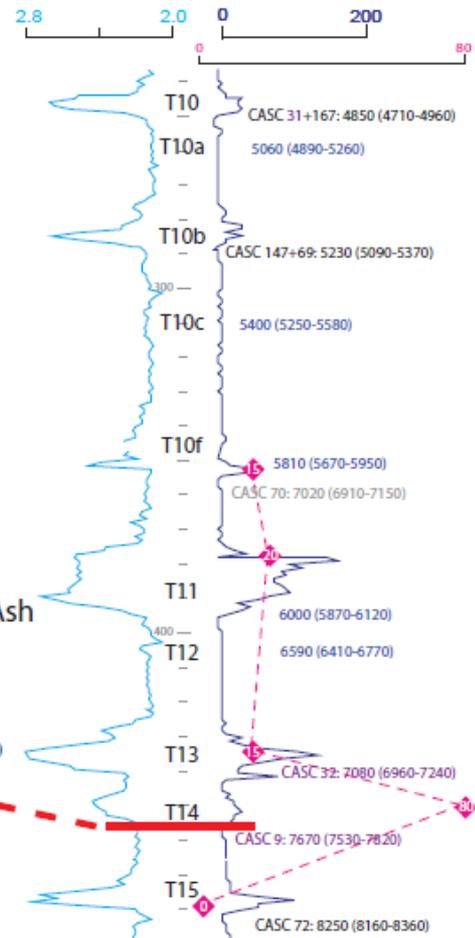
Correlation Summary



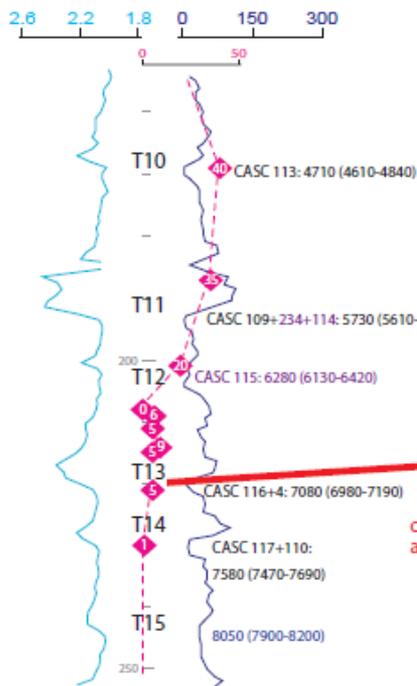
Hydrate Ridge Basin RR0207 56PC



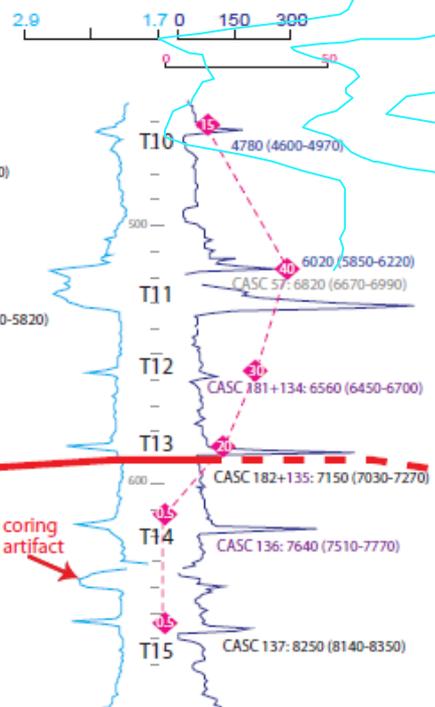
Rogue Canyon M9907-31 PC



Juan de Fuca Canyon M9907-12PC



Cascadia Channel M9907-23PC

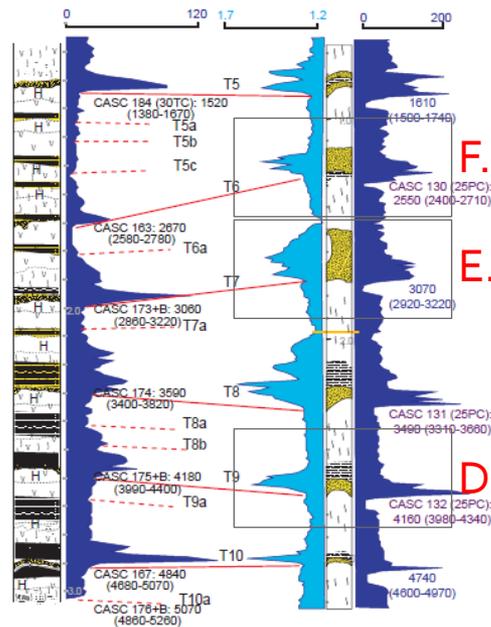


coring artifact

A. T5-T10

Rogue Channel Cascadia Channel

RR0207-55KC M9907-22PC



F.

E.

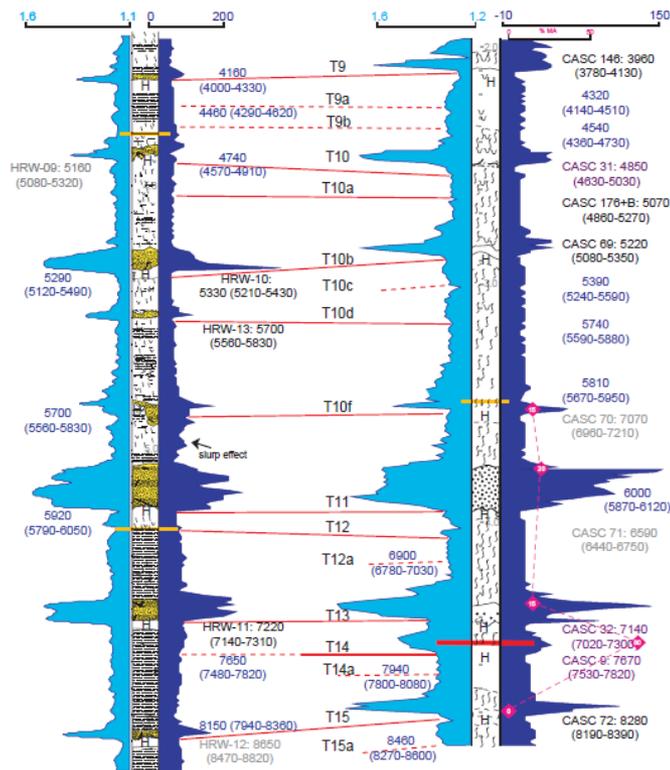
D.

B. T9-T15

Hydrate Ridge

Rogue Channel

RR0207-56PC M9907-31 PC

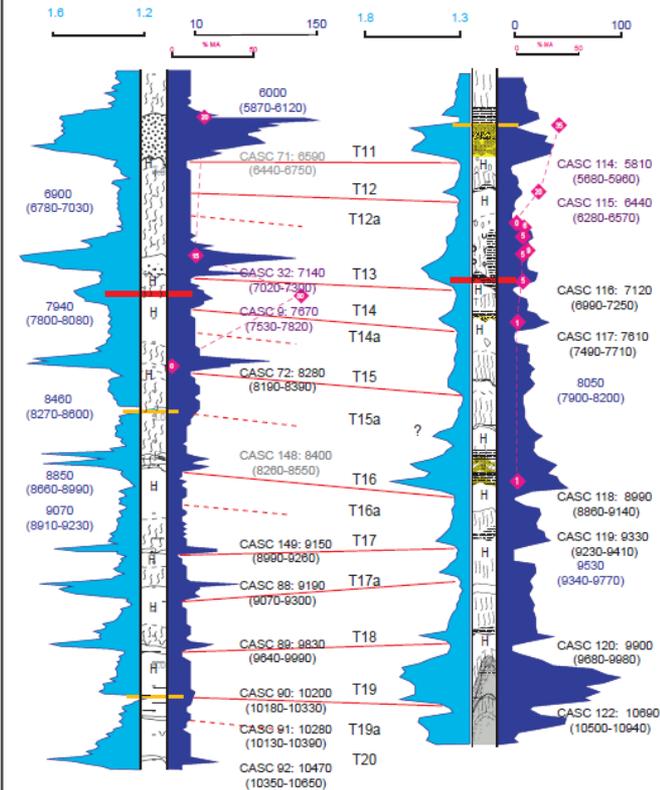


C. T11-T20

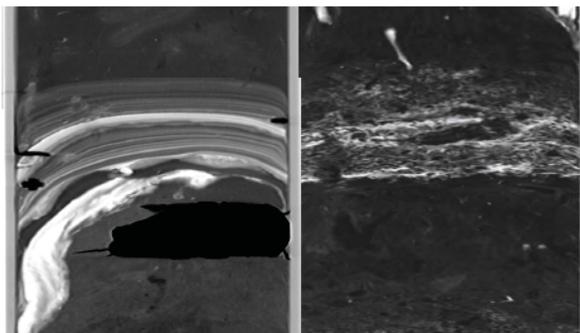
Rogue Channel

Juan de Fuca Channel

M9907-31 PC M9907-12PC

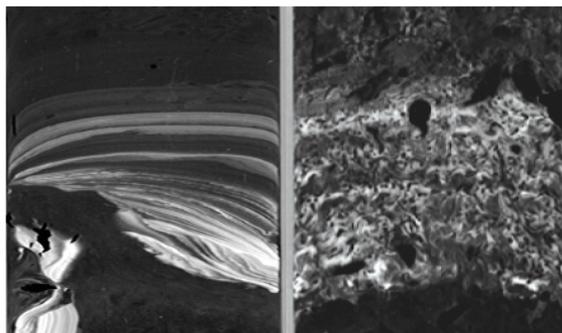


D.



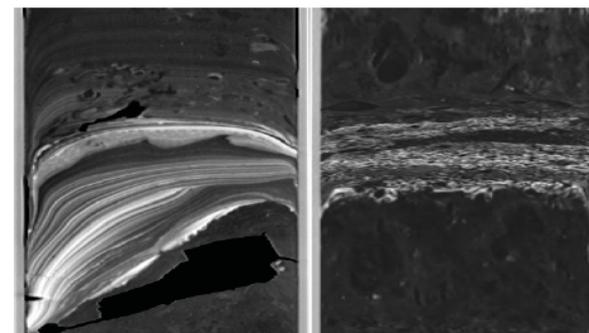
T9 Cascadia Ch. T9 Juan de Fuca Ch.

E.



T7 Cascadia Ch. T7 Juan de Fuca Ch.

F.



T6 Cascadia Ch. T6 Juan de Fuca Ch.

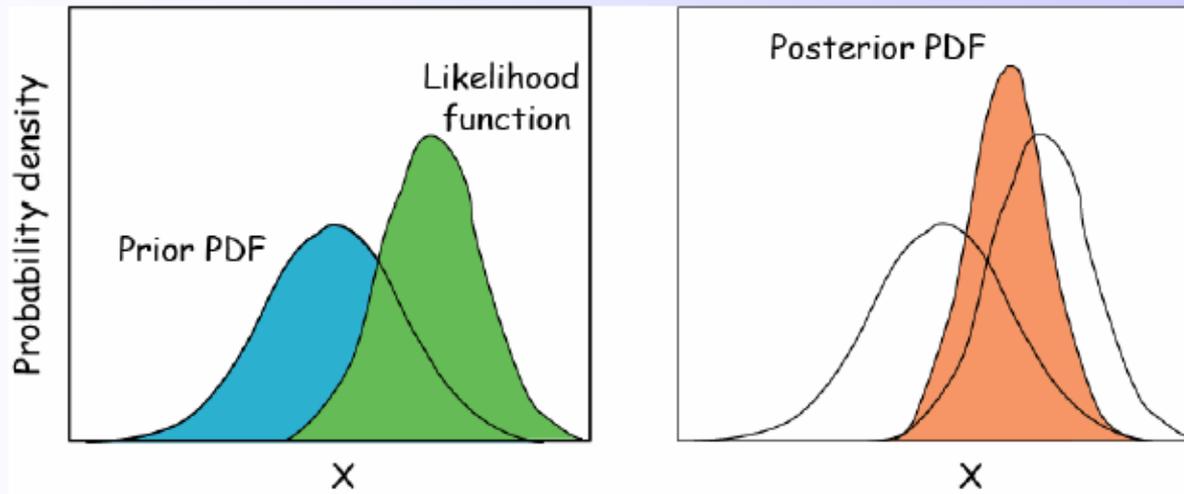
Now we'll develop the Bayesian model to assess the tests for turbidite correlation. We'll use the Washington margin as our example

But first... A Short Time-Out for Statistics!



Bayes' theorem

All information is expressed in terms of probability density functions



Bayes' rule (1763)

$$p(m|d, I) \propto p(d|m, I) \times p(m|I)$$

Posterior probability density \propto Likelihood \times Prior probability density

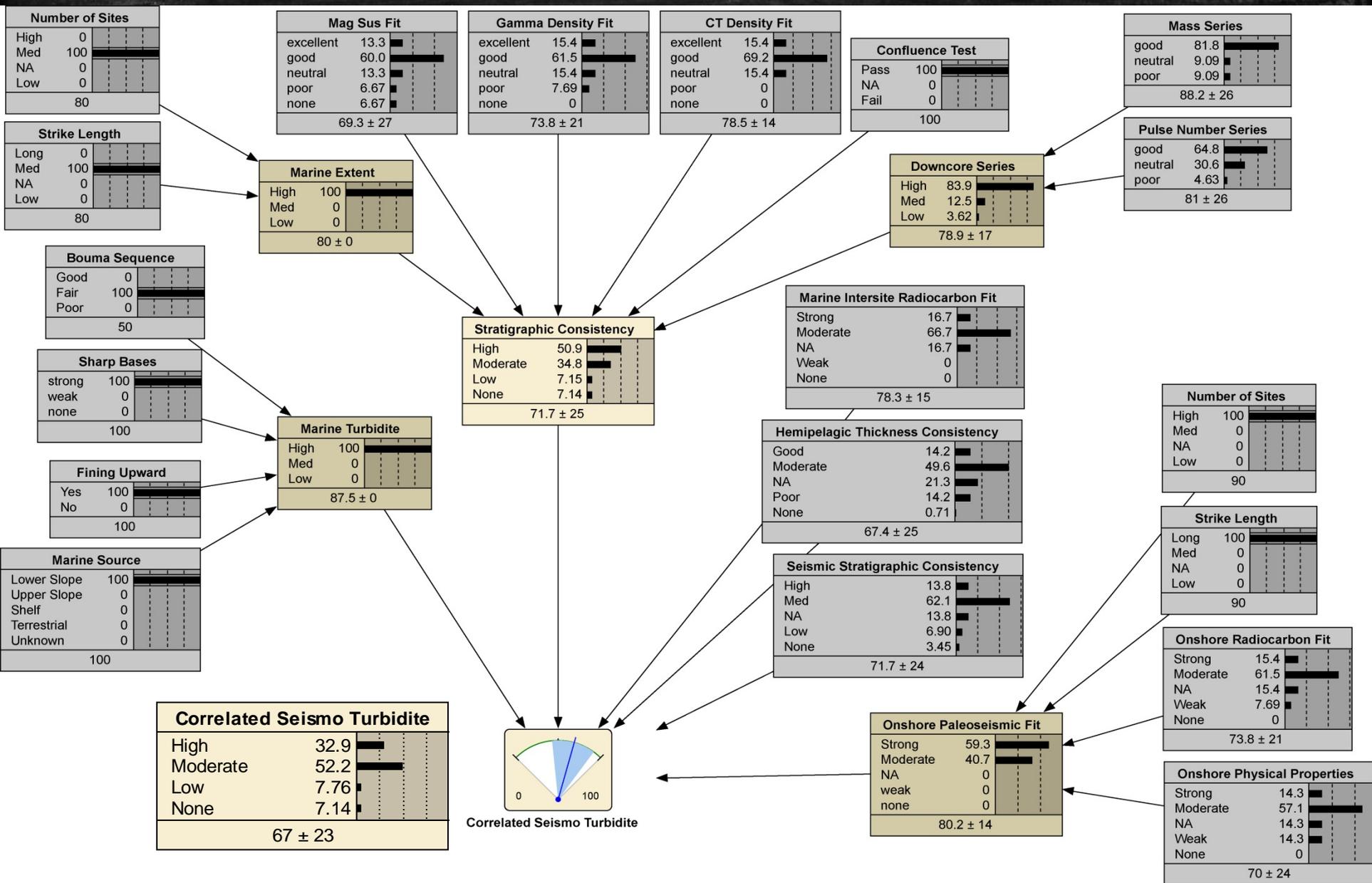
*What is known after
the data are collected*

*Measuring fit
to data*

*What is known before
the data are collected*

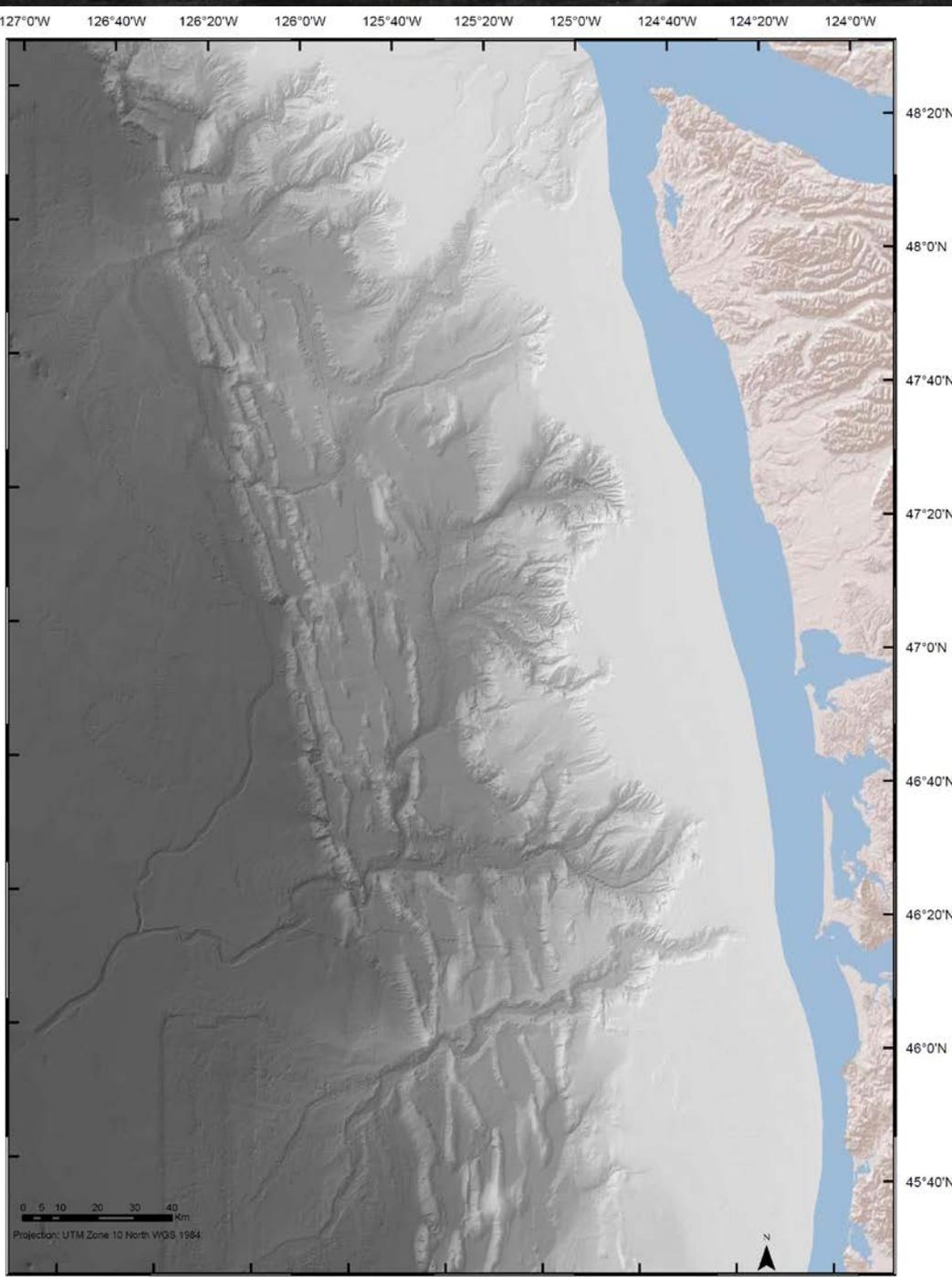


1702-1761



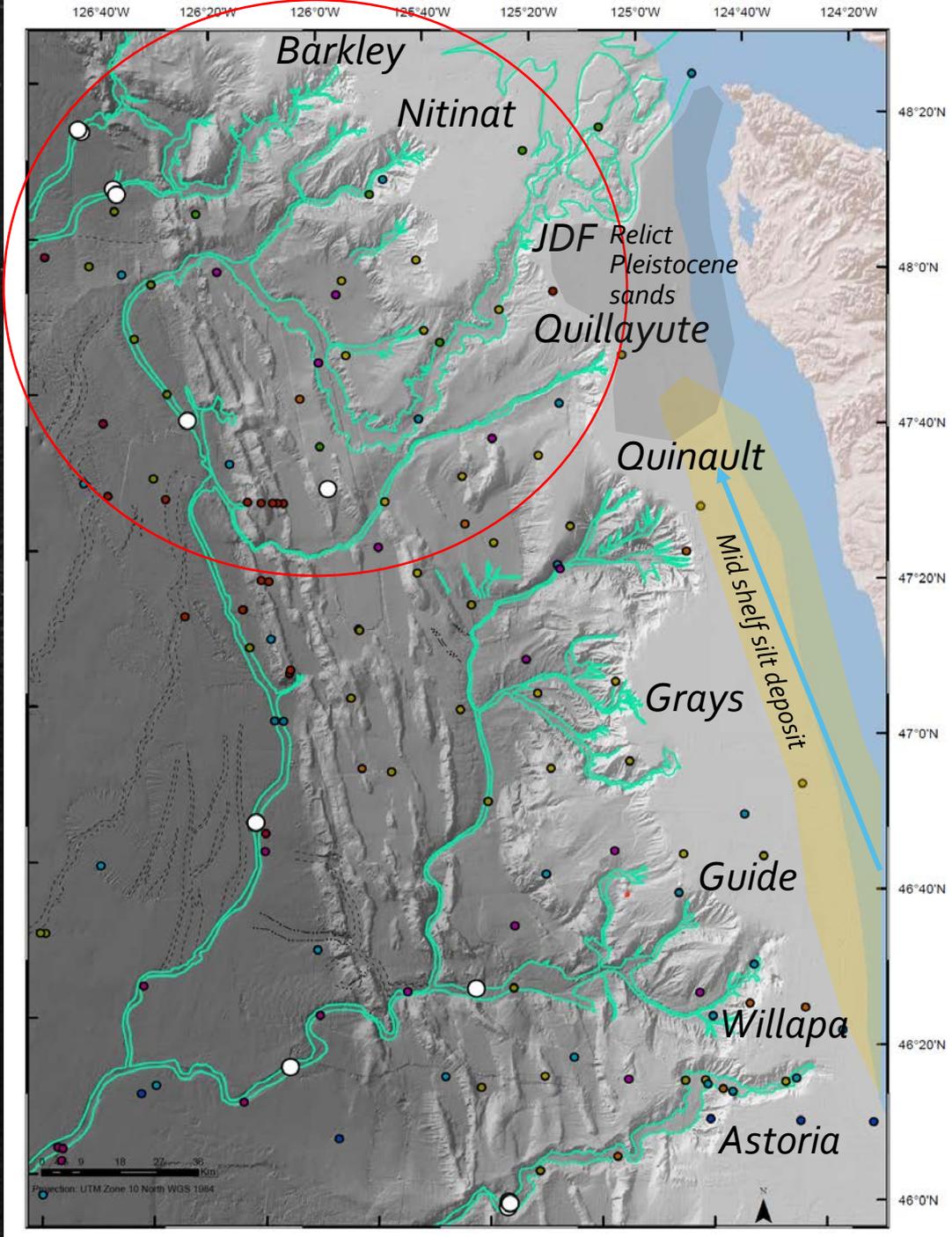
Bayesian Probabilities of earthquake origin under uncertainty

Every model follows the same rule: garbage in, garbage out. So we'll spend most of the talk looking at the depositional details and the data...



Northern Cascadia

Nearly complete multibeam (various resolutions) now exists. OSU cruises 1992-2012 cover ~ 80% of the continental slope of Washington. Other sources cover major portions of the southern Canadian slope.



Northern Cascadia

Let's examine some of the complexities in the Cascadia system, and how they can be used to advantage, (or not!)

All existing cores are shown, our primary cores (Professional Paper 1661-F) in white.

All of these systems are more or less relict since the late deglacial/early Holocene.

127°0'W

126°40'W

126°20'W

126°0'W

125°40'W

125°20'W

48°20'N

48°0'N

47°40'N

M9907-08PC
M9907-09PC

M9907-07PC

M9907-06PC

TT039 010

NV952 2

NV912 12

NV951 6

NV967 1

NV967 2

TT029 023

TT039 009

BB291 004

TT063 020

TT063 019

W8306C 16

NV967 5

TT039 026

TT039 027

TT063 018

NV982-14

NV967 7

NV982 11

NV967 6

NV952 1

NV982 13

NV912 10

TT039 004

BB312 014

NV951 1

TT063 016

BB291 008

TT029 025

M9907-05PC

NV912 11

NV952 4

NV951 3

NV967 9

NV967 8

NV912 08

0 1.5 3 6 9 12 Km

Projection: UTM Zone 10 North WGS 1984

TT039 007

BB291 004

M9907-10PC

NV967 10

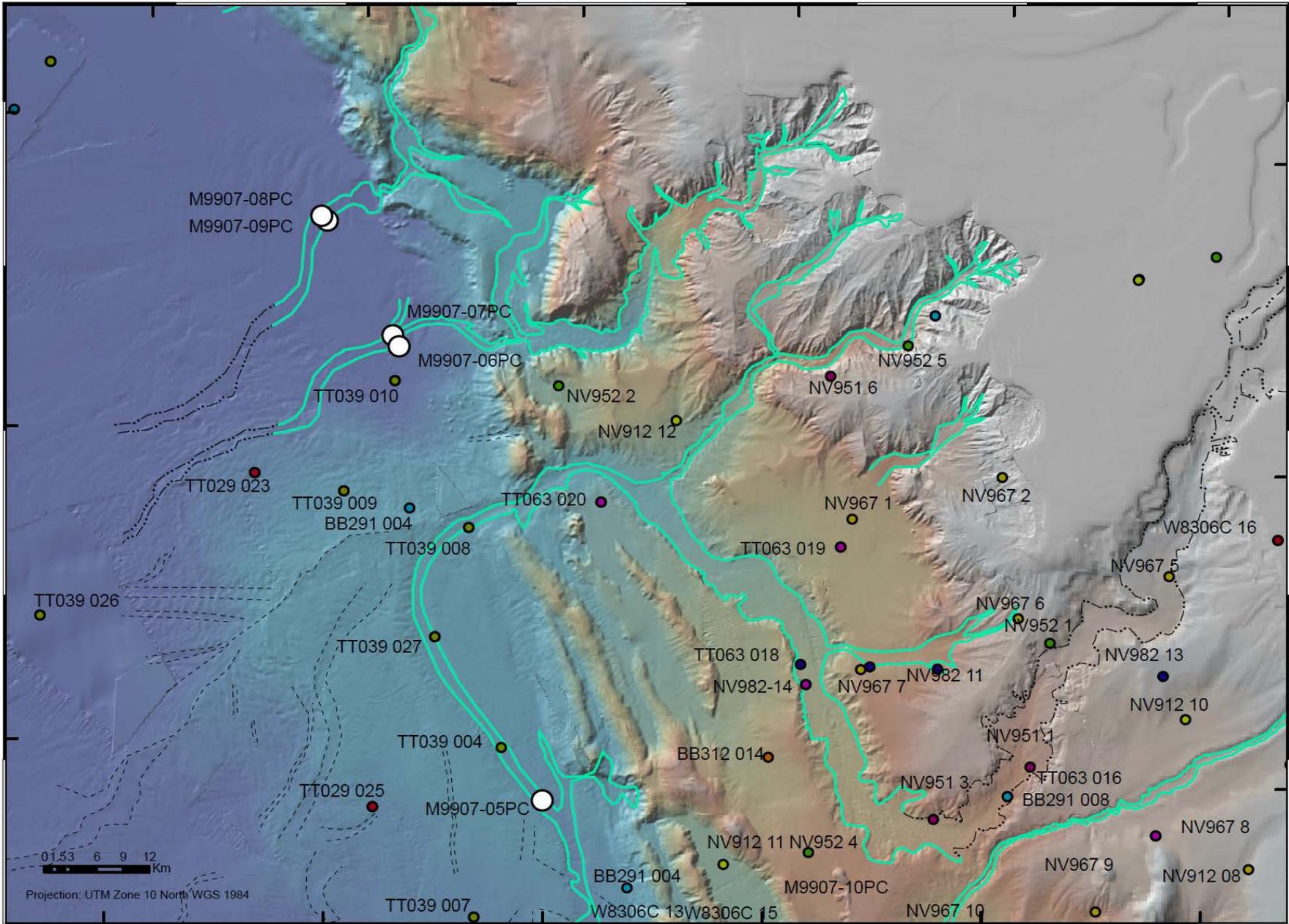
W8306C 13

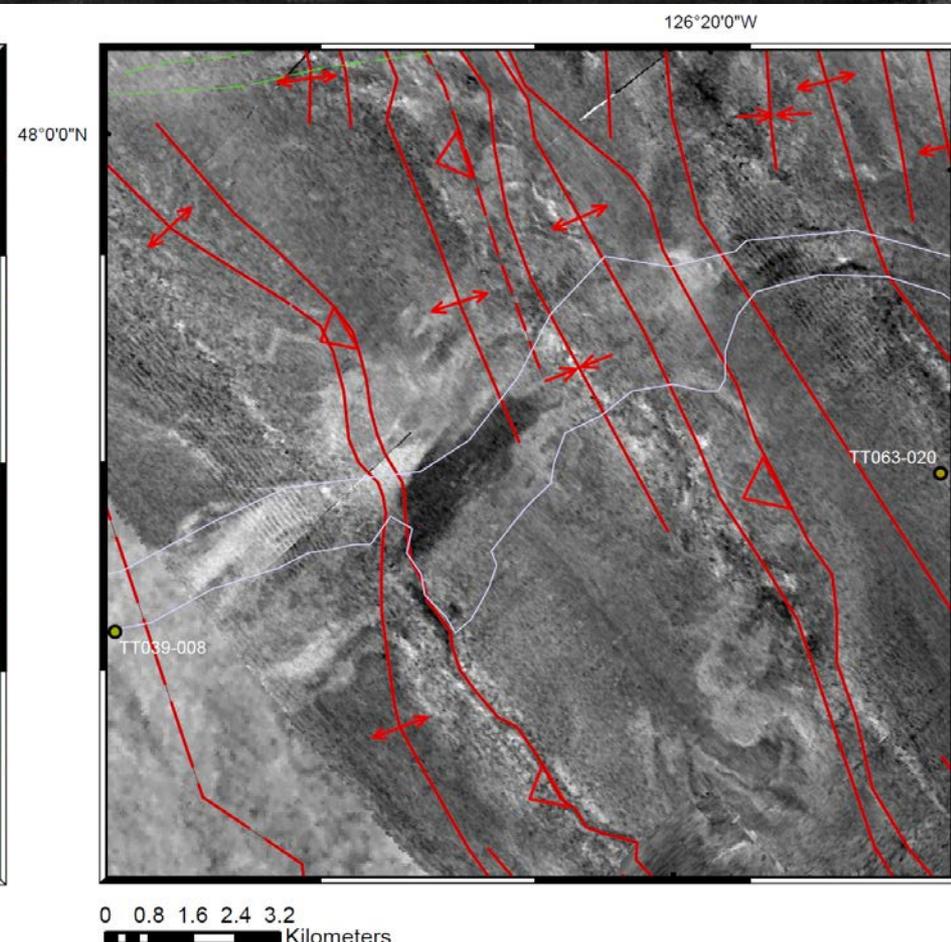
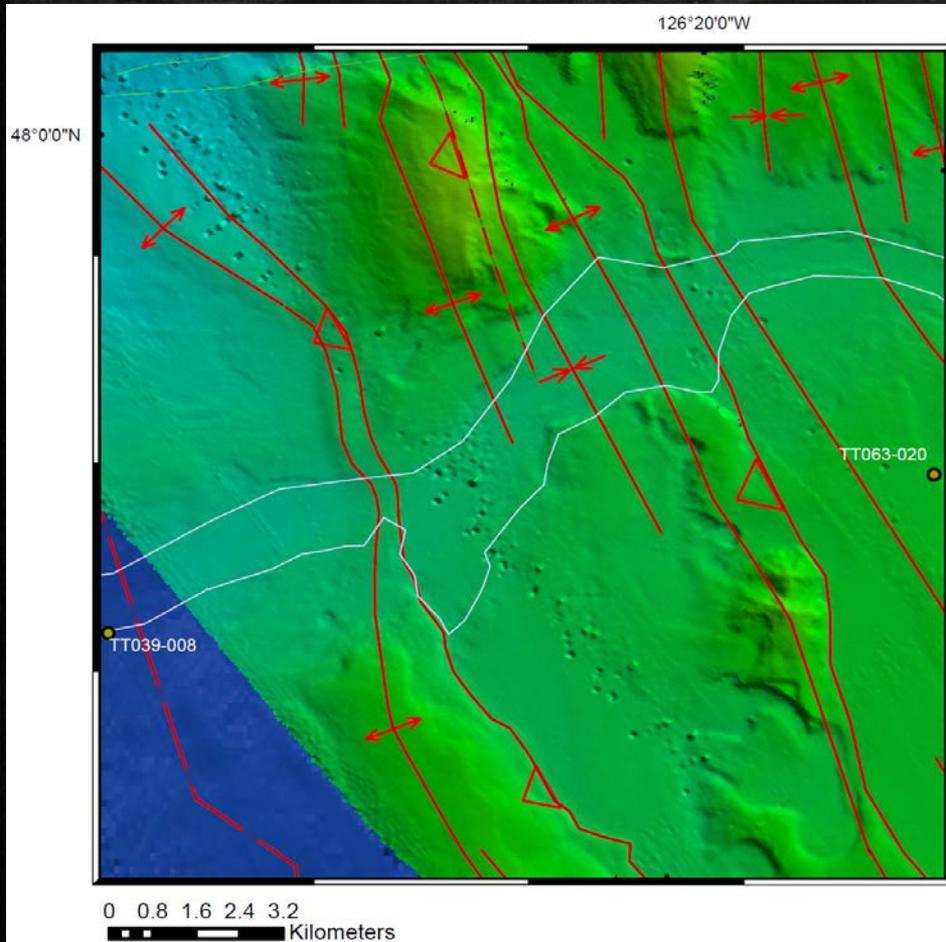
W8306C 15

NV967 10

NV967 9

NV912 08





Partial blockage and Holocene turbidity current response, frontal thrust, JDF outlet, Nitinat Fan Apex.

Visible in the backscatter are the fine grained pool behind the landward vergent thrust fault blockage, crevasse-splay spillover to the north, and sand tongue continuing down JDF Channel.

127°0'W

126°40'W

126°20'W

126°0'W

125°40'W

125°20'W

explanation
 234/**21** Holocene
 thickness/turbs.

 Inactive fan
 Active fan

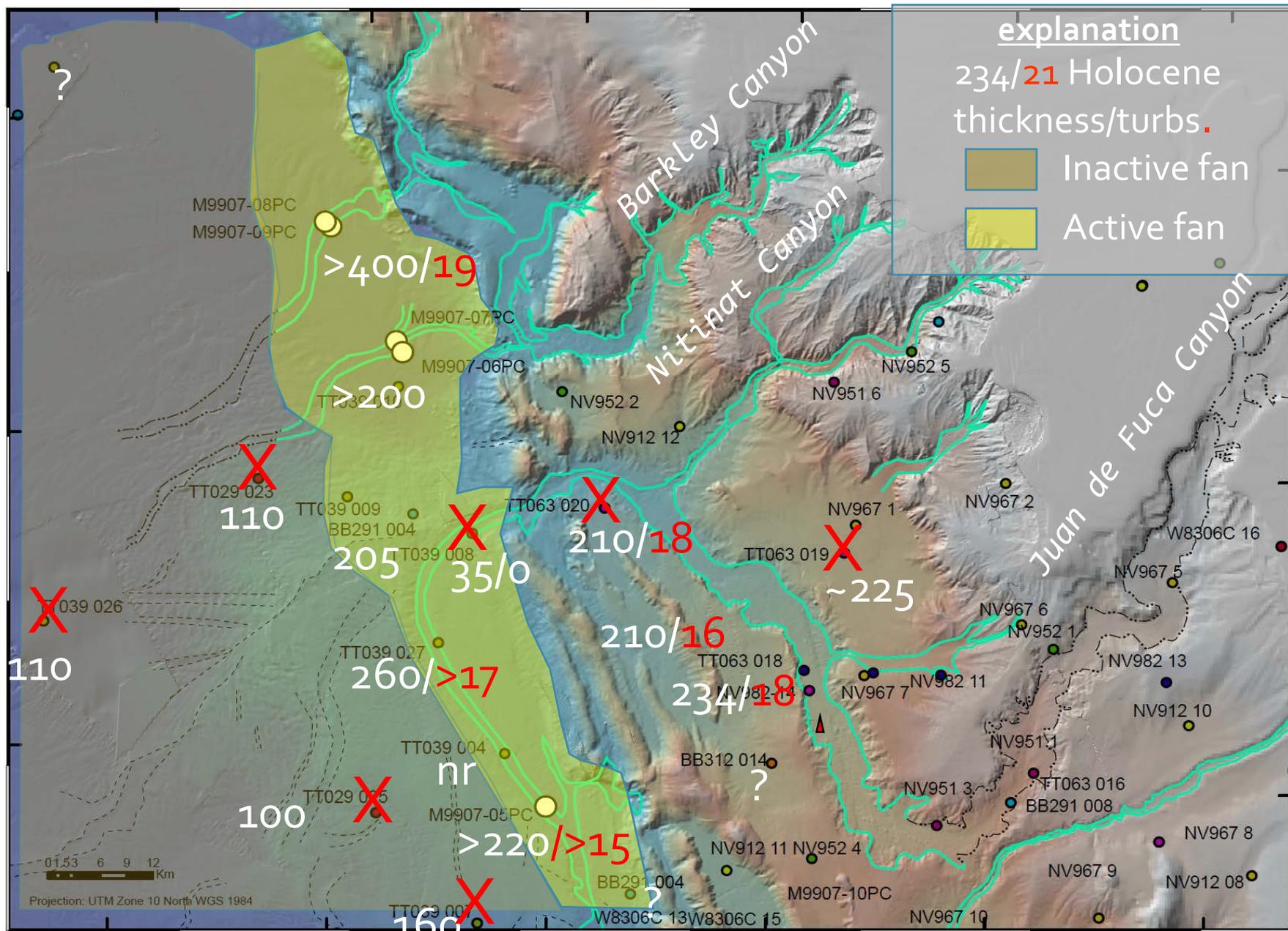
48°20'N

48°0'N

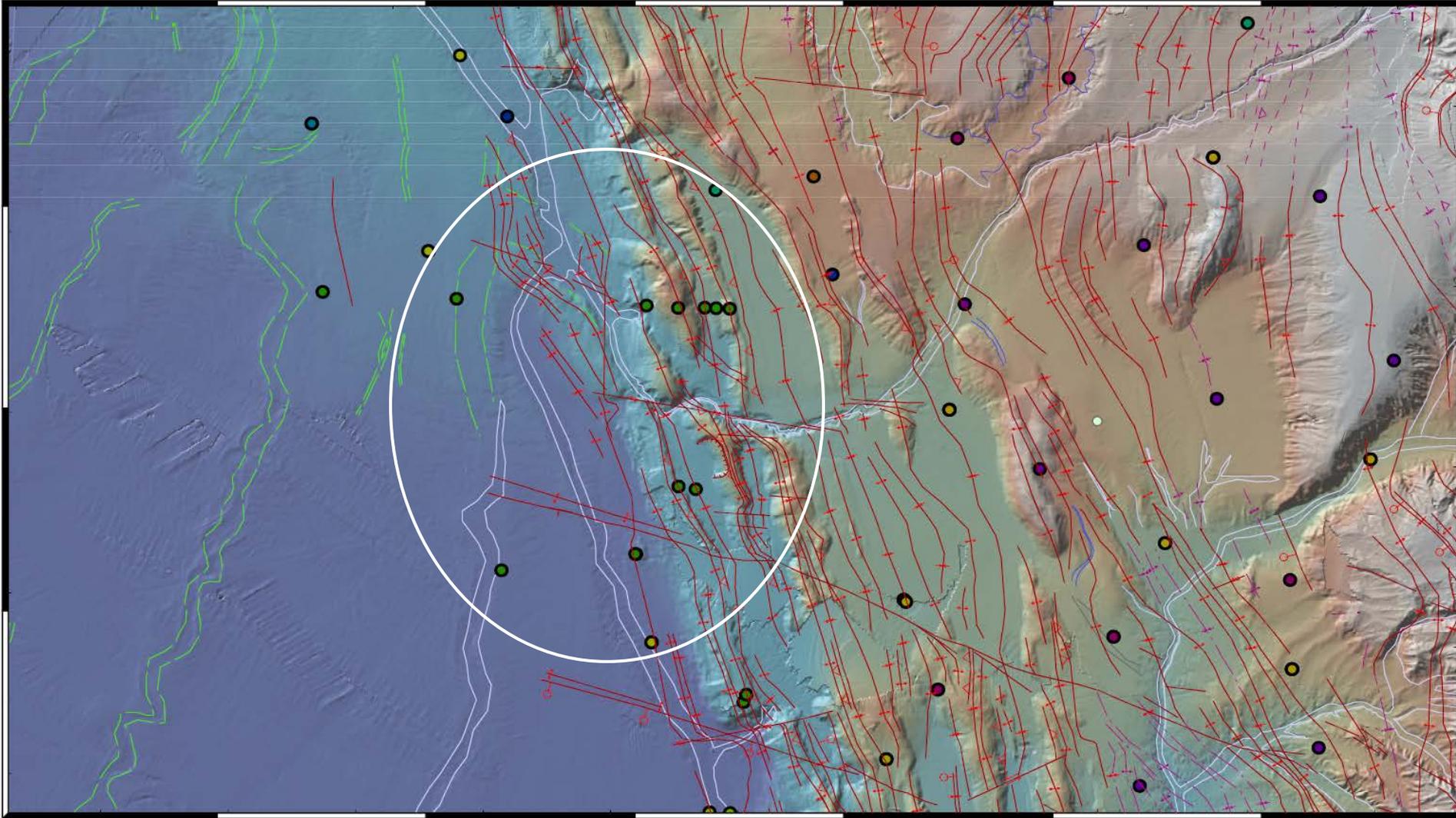
47°40'N

Barkley Canyon
 Nitinat Canyon

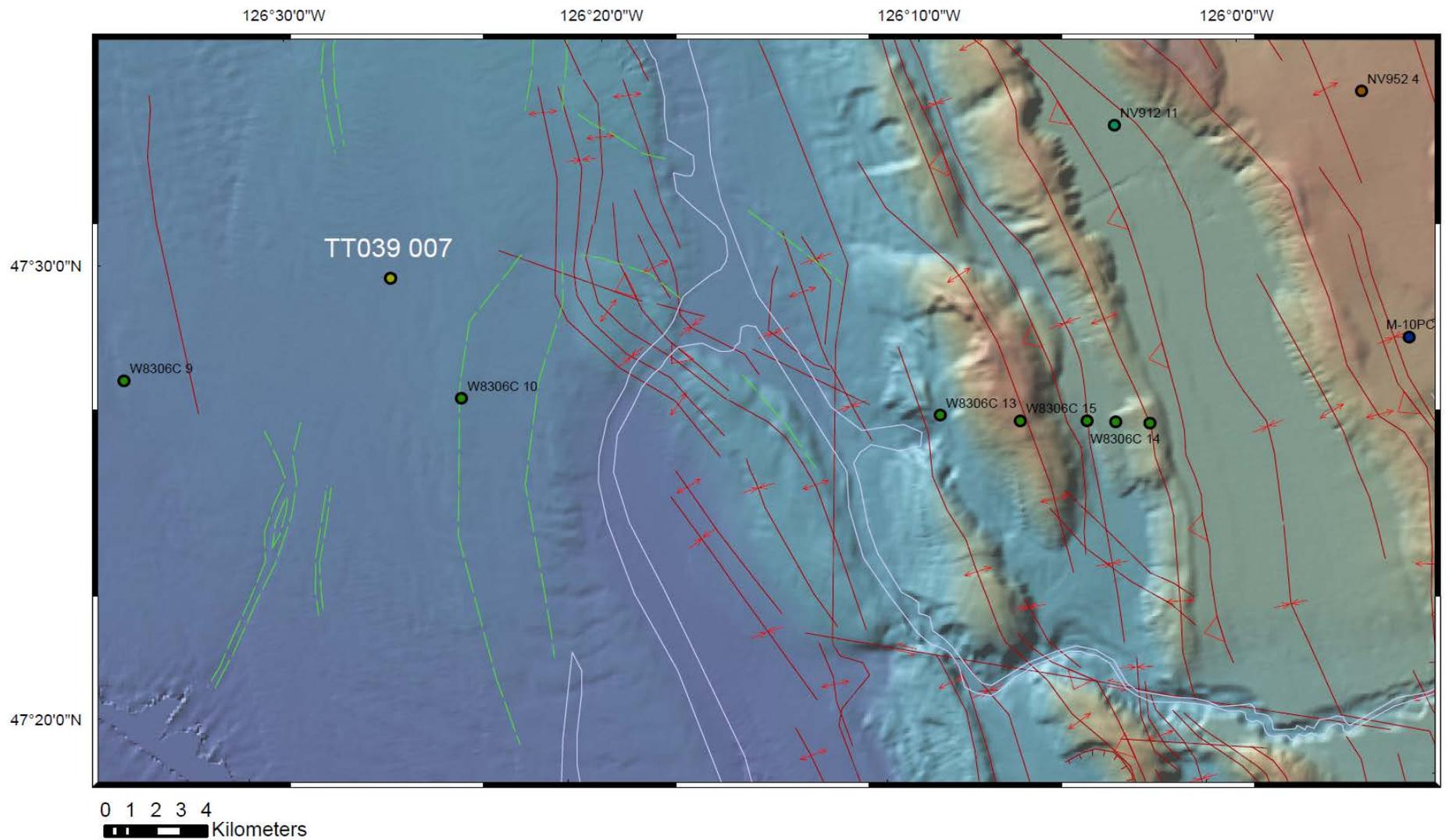
Juan de Fuca Canyon



Projection: UTM Zone 10 North WGS 1984



Central WA margin: Quillayute confluence, landslides and fault intersections

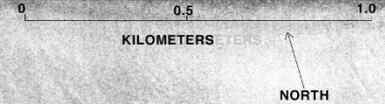


Looking closer, we see a very complex intersection with the frontal landward vergent thrust, partial blockage of Quillayute Canyon and the JDF Channel.

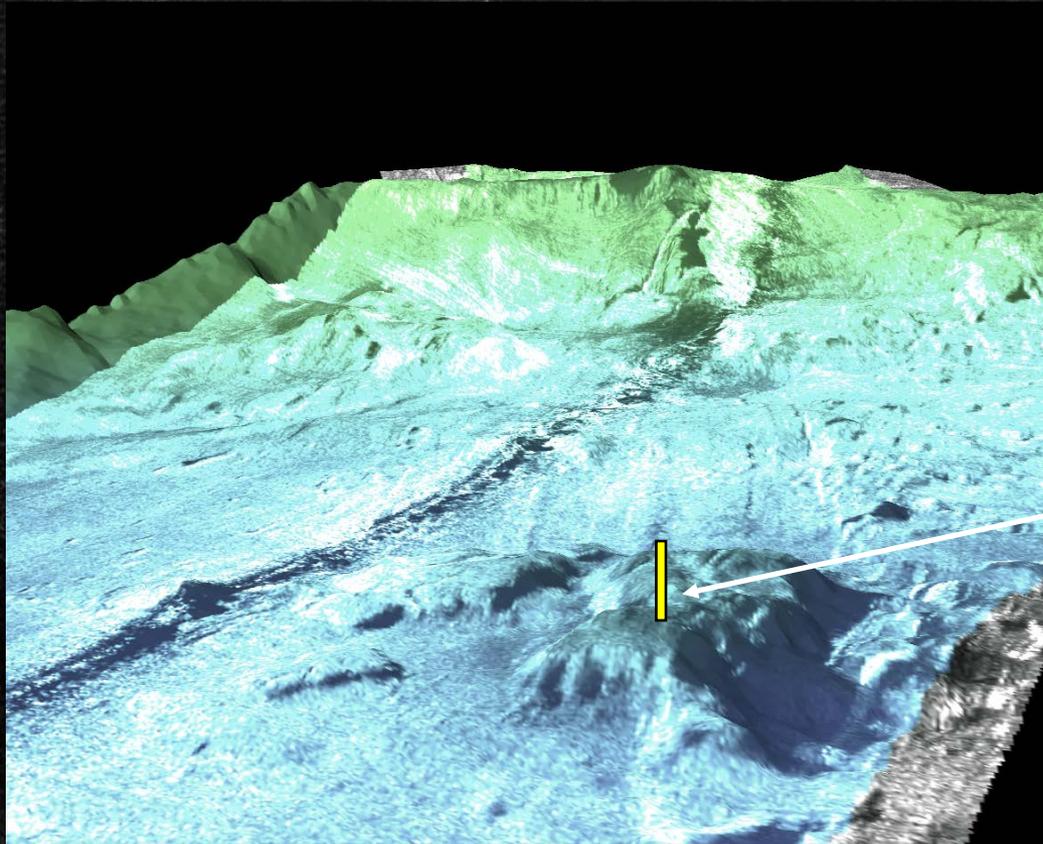
Mud volcano along fault

JDF Channel

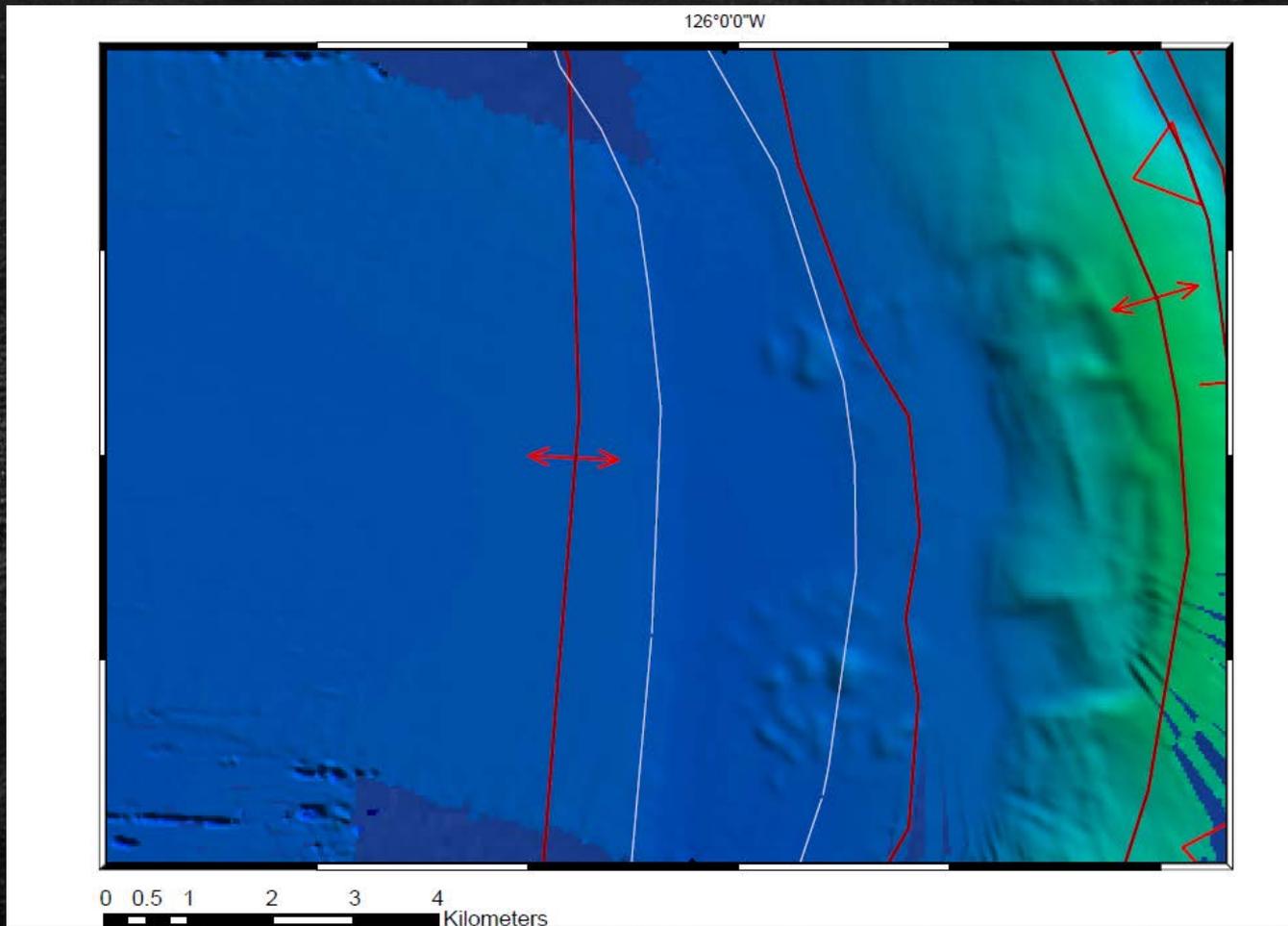
N. Nitinat Fault



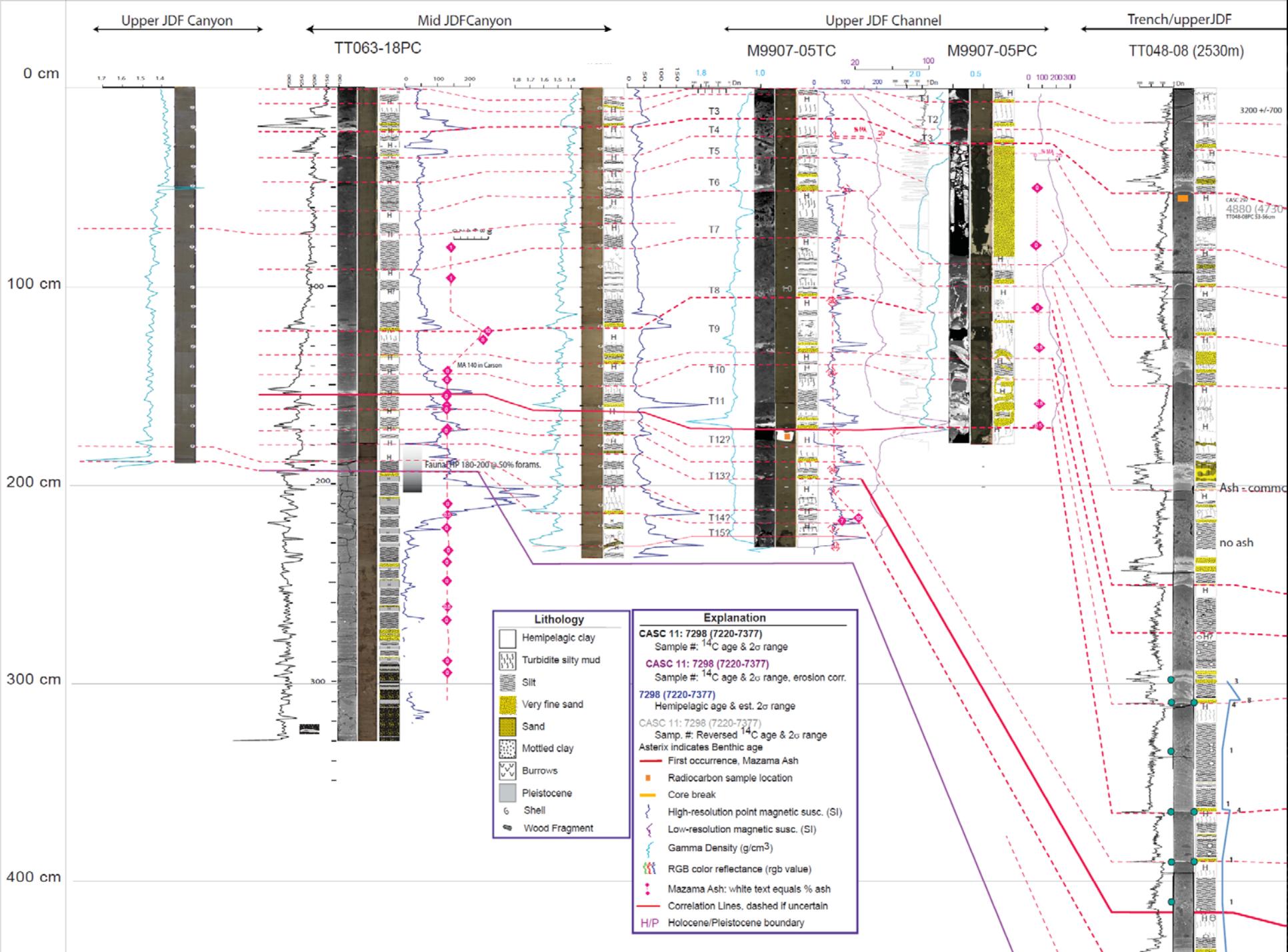
North Nitinat Fault. SeaMARC 1A image shows fault, mud volcano, and channel deflection

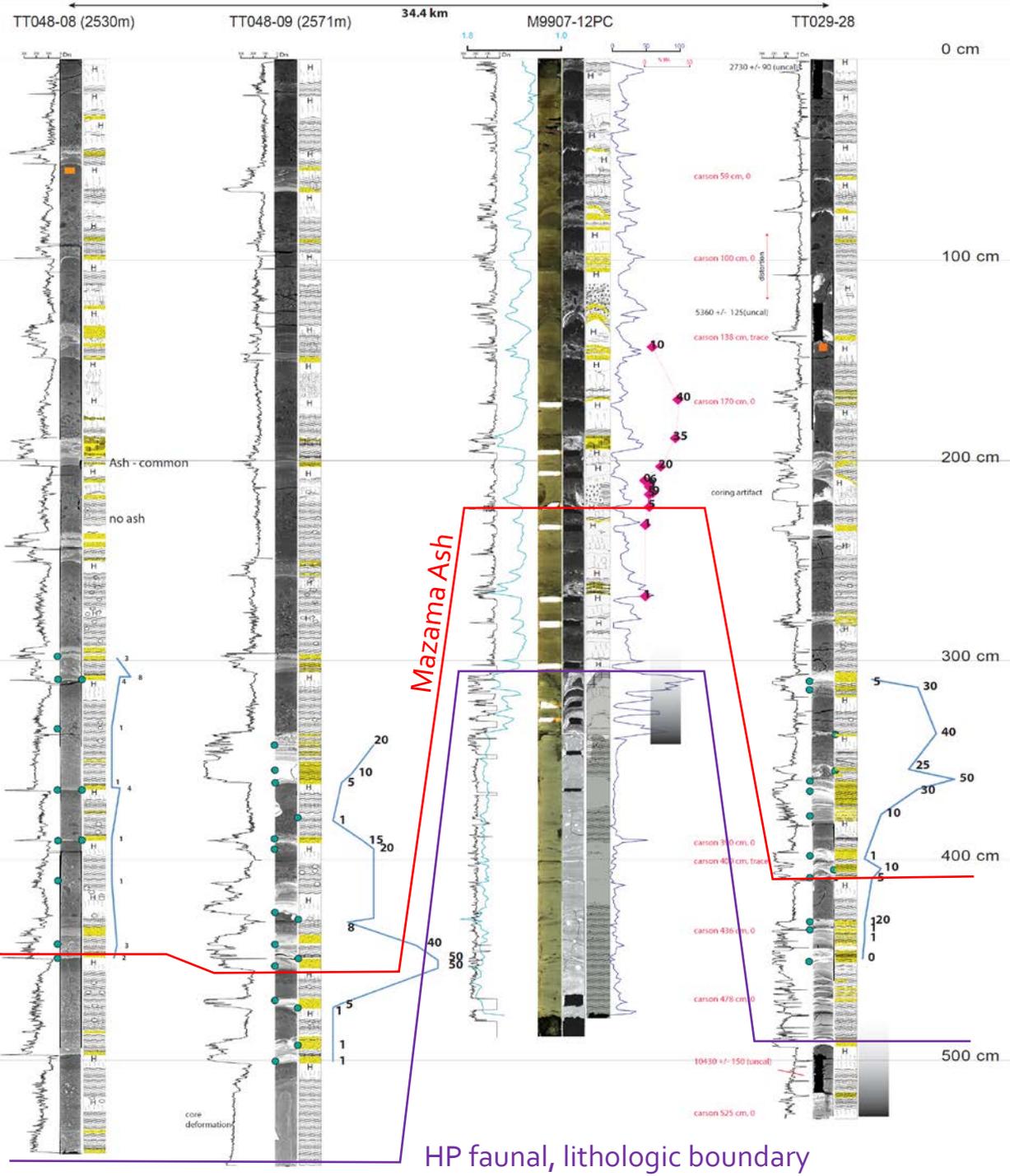


Core W8306c-02.
Hemipelagic
sediment
overlying this slide
block (just 1 km
south of the upper
image) suggests
age is ~ 2000 BP



And further south, two more landslides are in JDF channel. Just downstream, our primary JDF core (M9907-12PC has a relatively thin Holocene section less than 3 m.





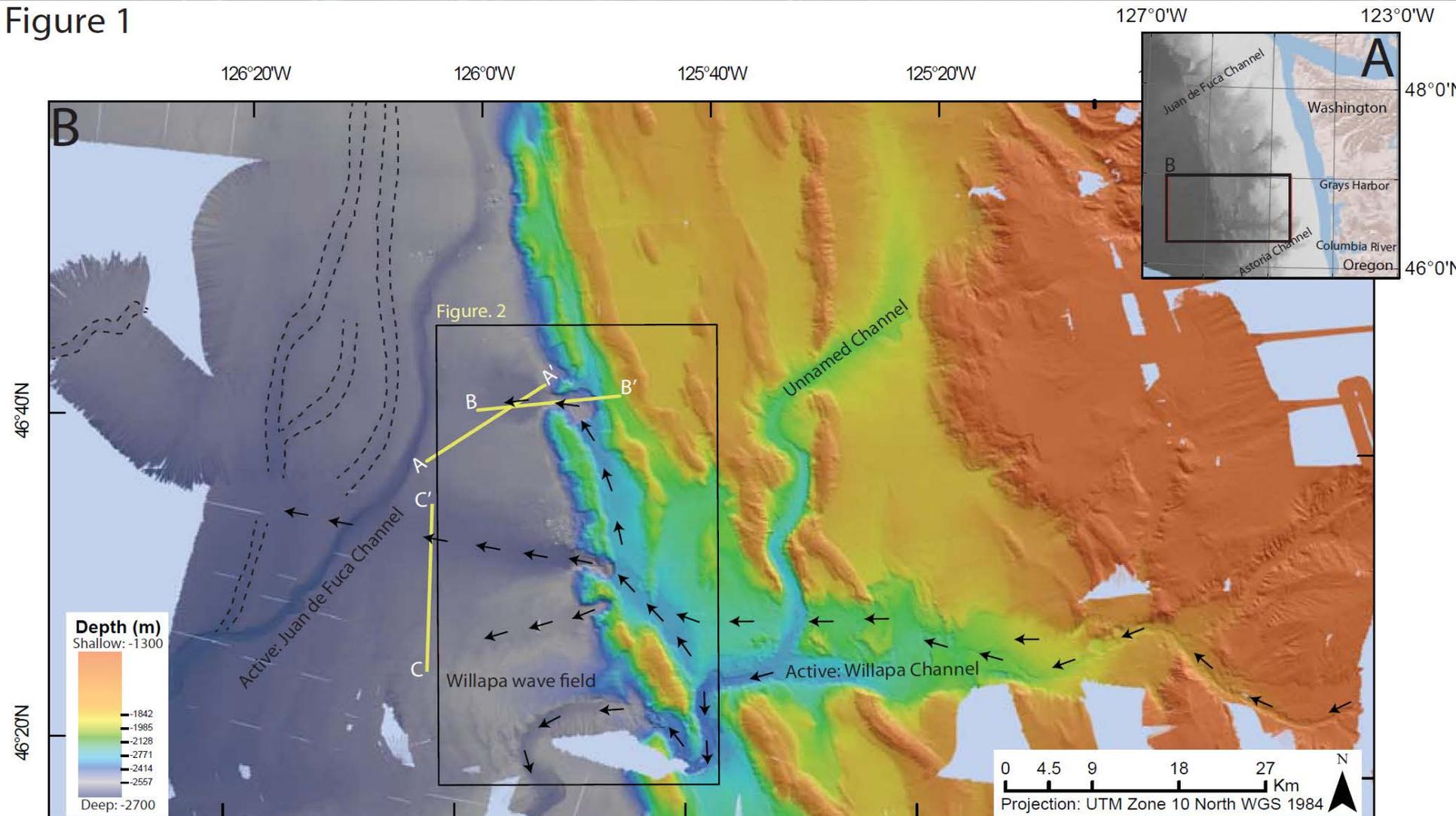
Using CT data and more ash analysis from this project, we observe 18-19 Holocene beds in these cores, and 13-15 above the Mazama ash which is abundant in 3 and sparse in one core.

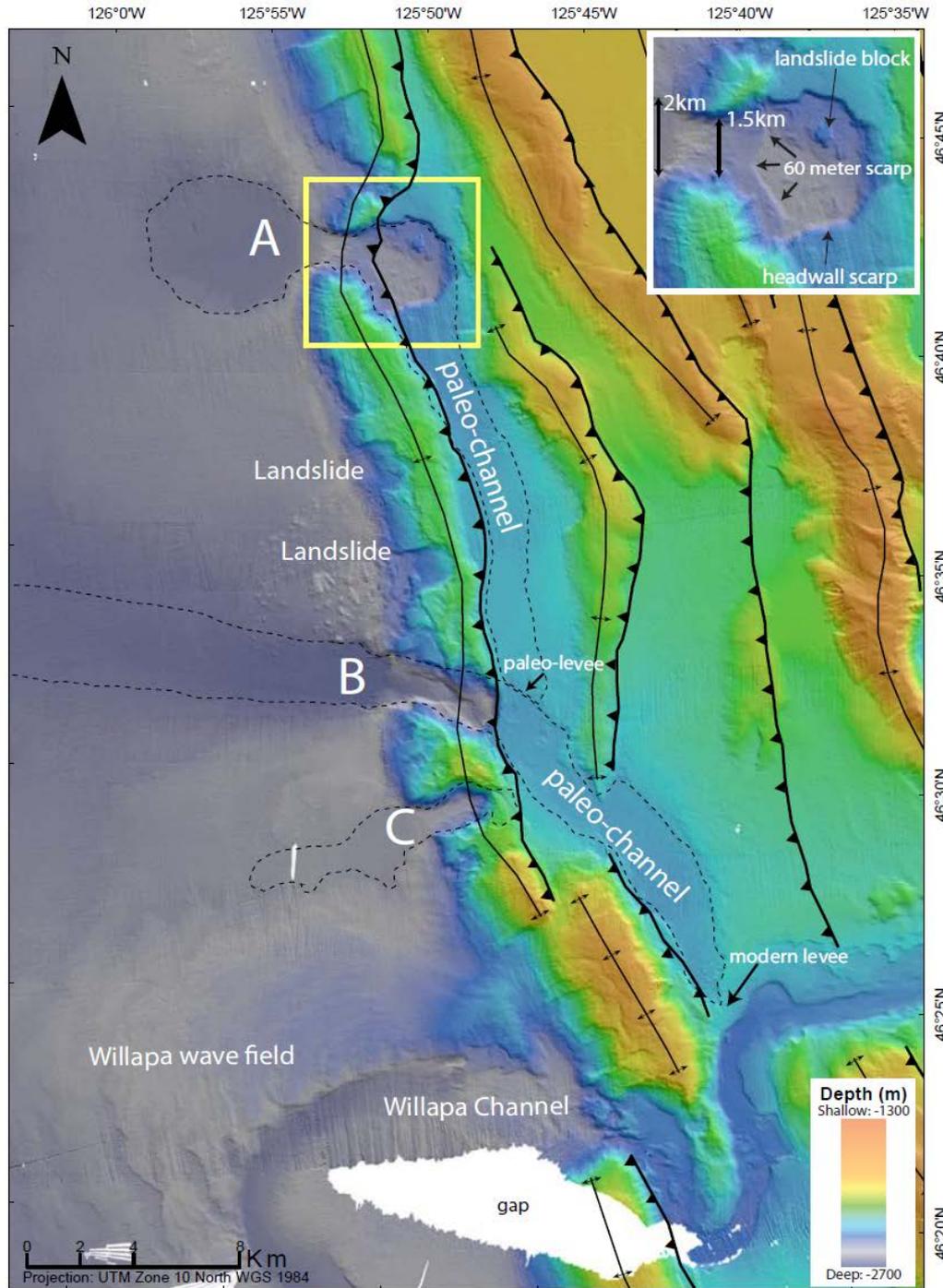
The difference with the old visual logs is due to using CT data. Visual logging is inadequate (unless beds are coarse grained).

Additionally, ash presence and abundance is very spotty and irregular. Note Carson (1971) missed the ash in TT29-28 altogether.

These two paleochannels were mapped as inactive in 1999...

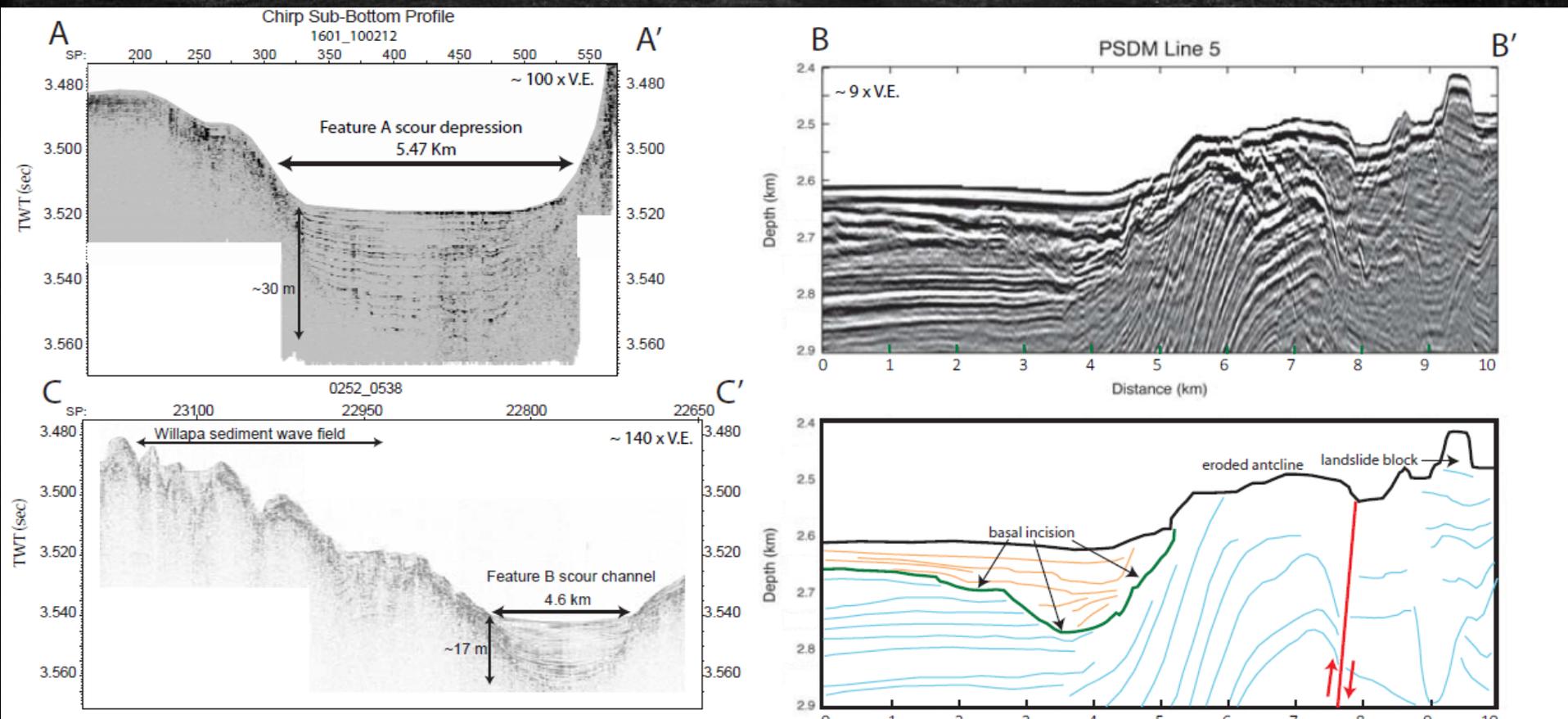
Figure 1

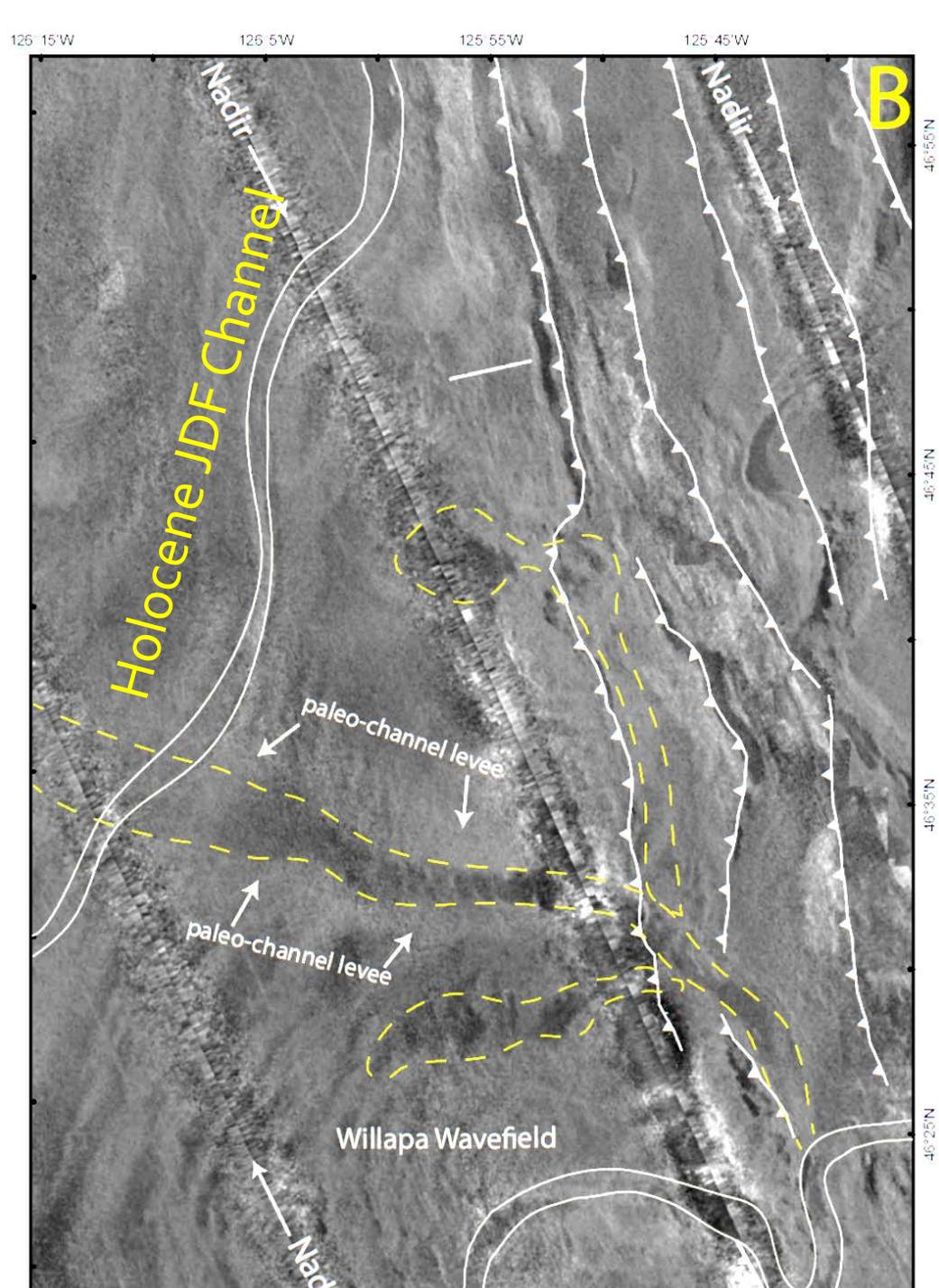




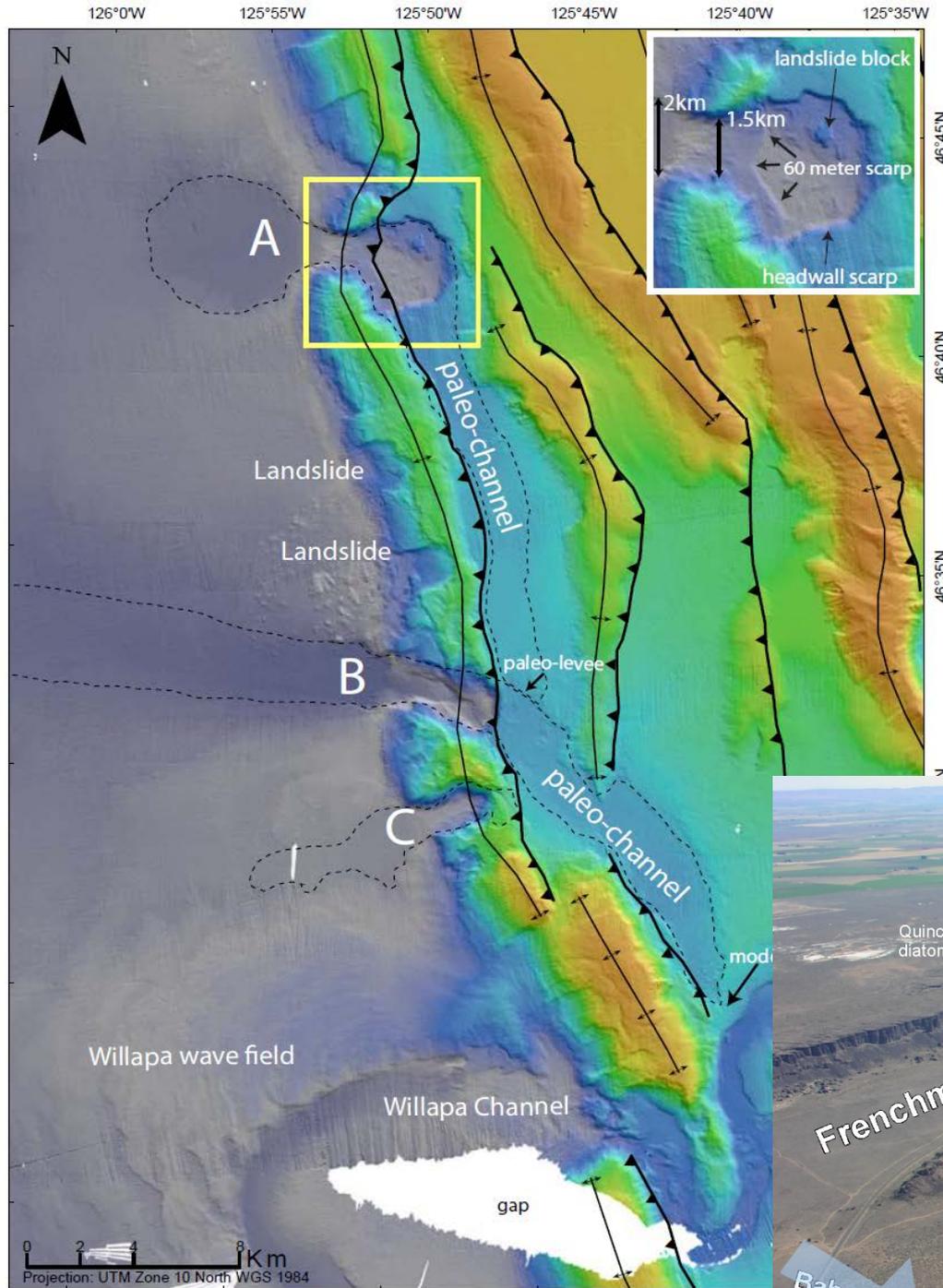
- The two features "A" and "B" extend westward across the fall line at nearly right angles, indicating a high flow regime
- J. Beeson et al 2013, in revision.

These features are not landslide related, they are incised into the abyssal plain, and have no landslide blocks. They most resemble plunge pools, but do not have the vertical drop required of plunge pools.

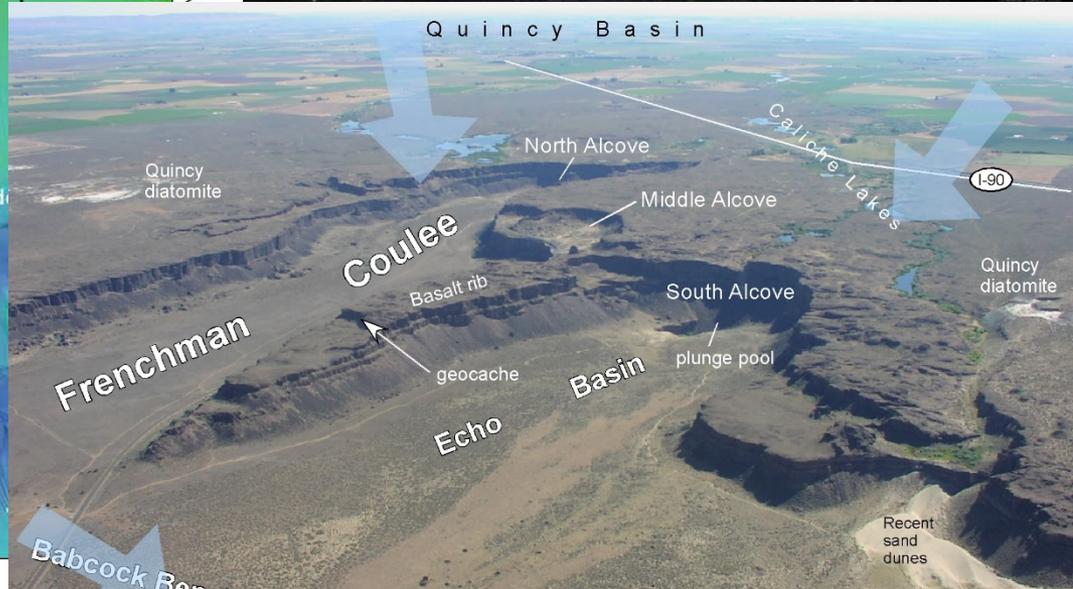


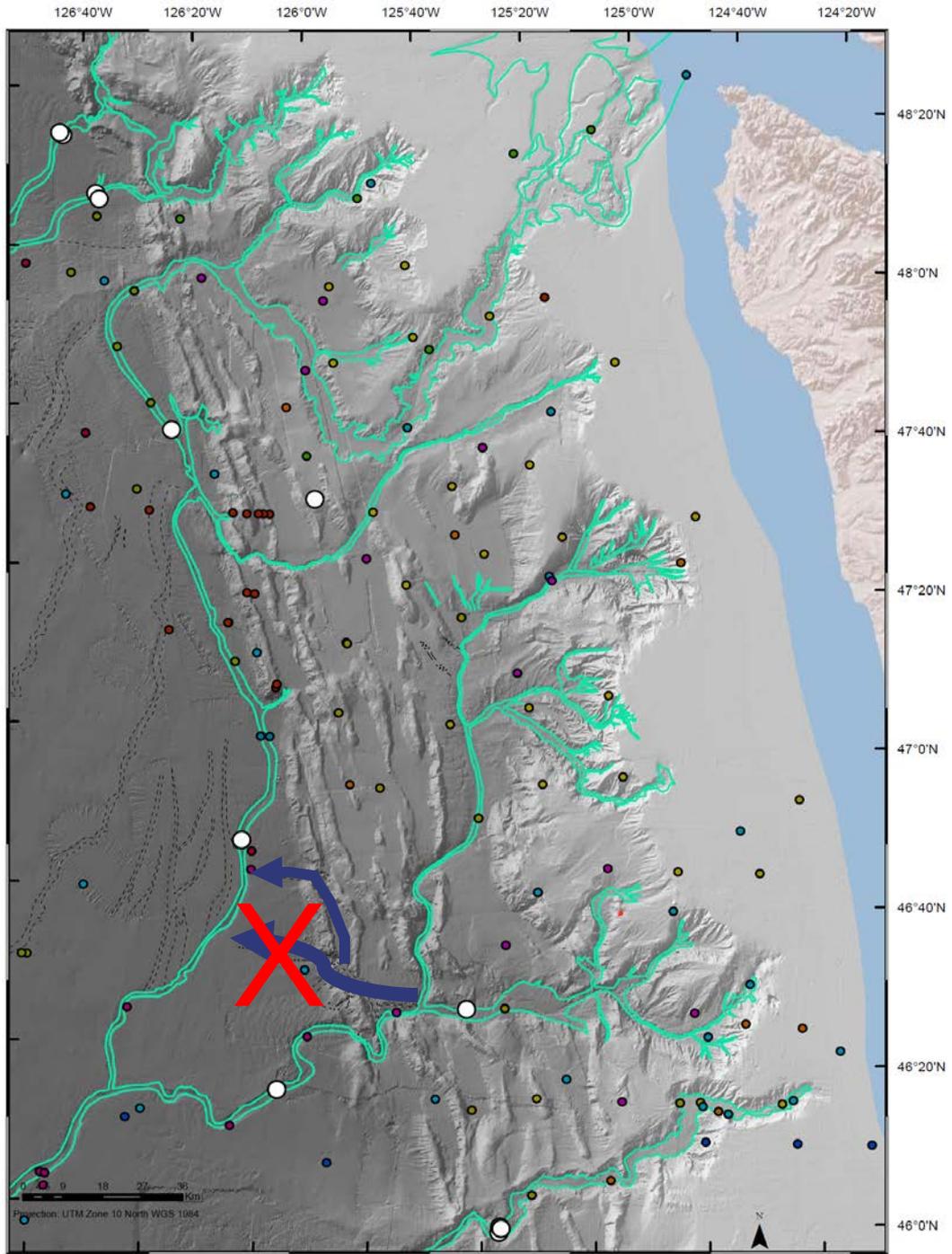


Ages are constrained by crosscutting relationships to be late Pleistocene. They are inactive in the Holocene, as originally mapped.



- The two features "A" and "B" are interpreted as submarine "coulees" similar to those observed onshore and most likely related to the Missoula flooding events of the deglacial period.



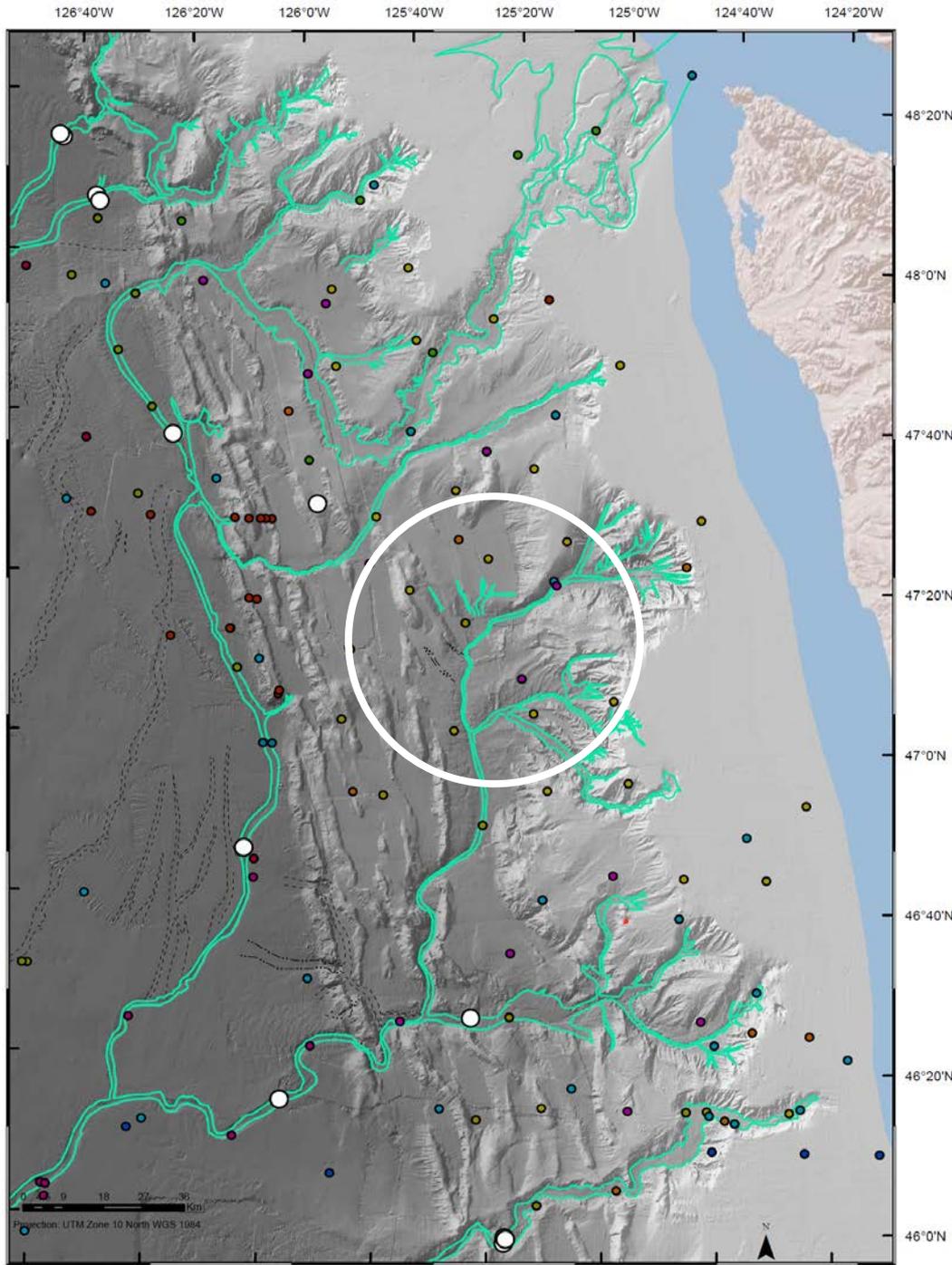


Cascadia

Alternative Holocene channel?

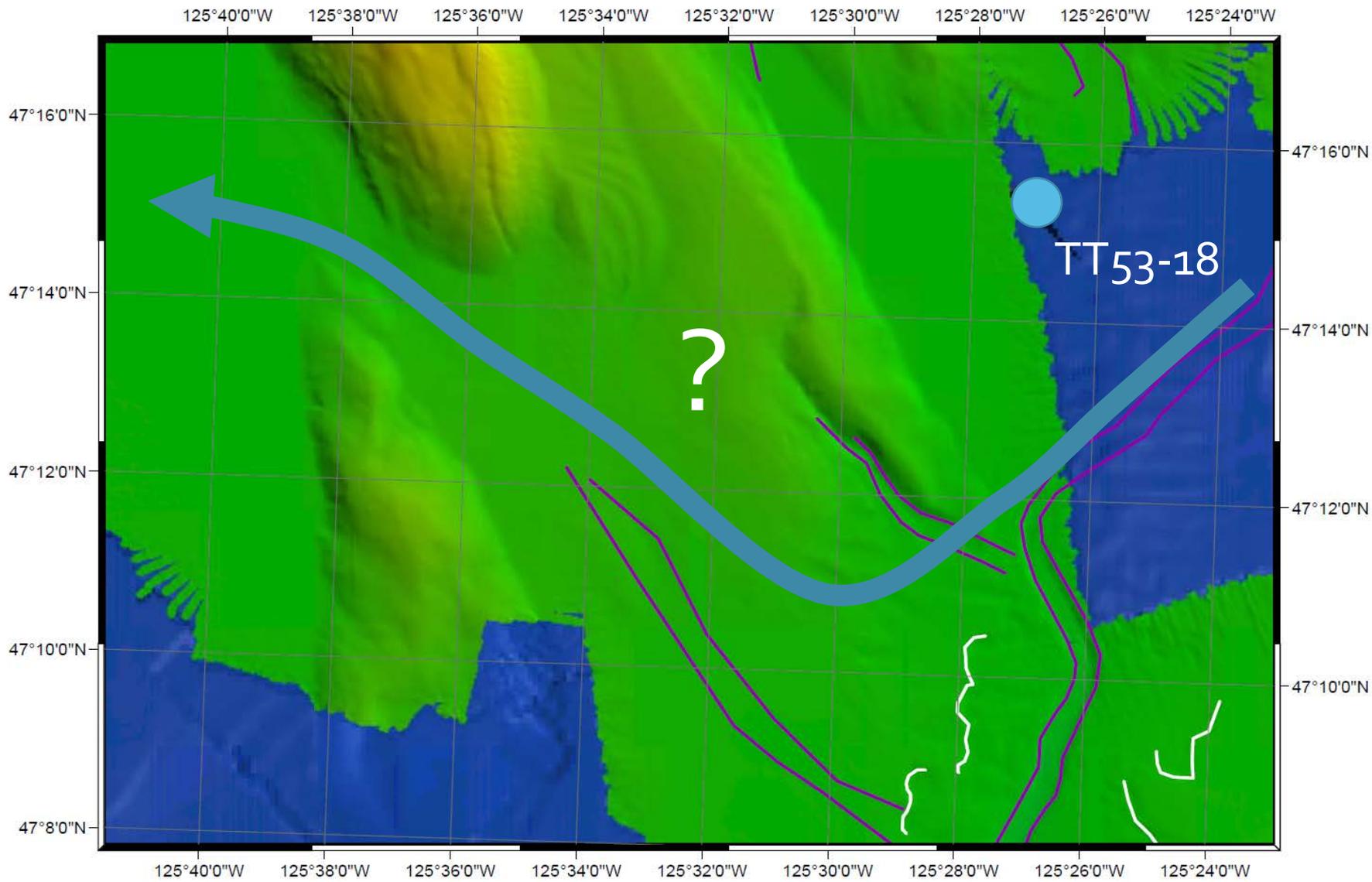
In the Missoula outbursts yes, not significant in the Holocene.

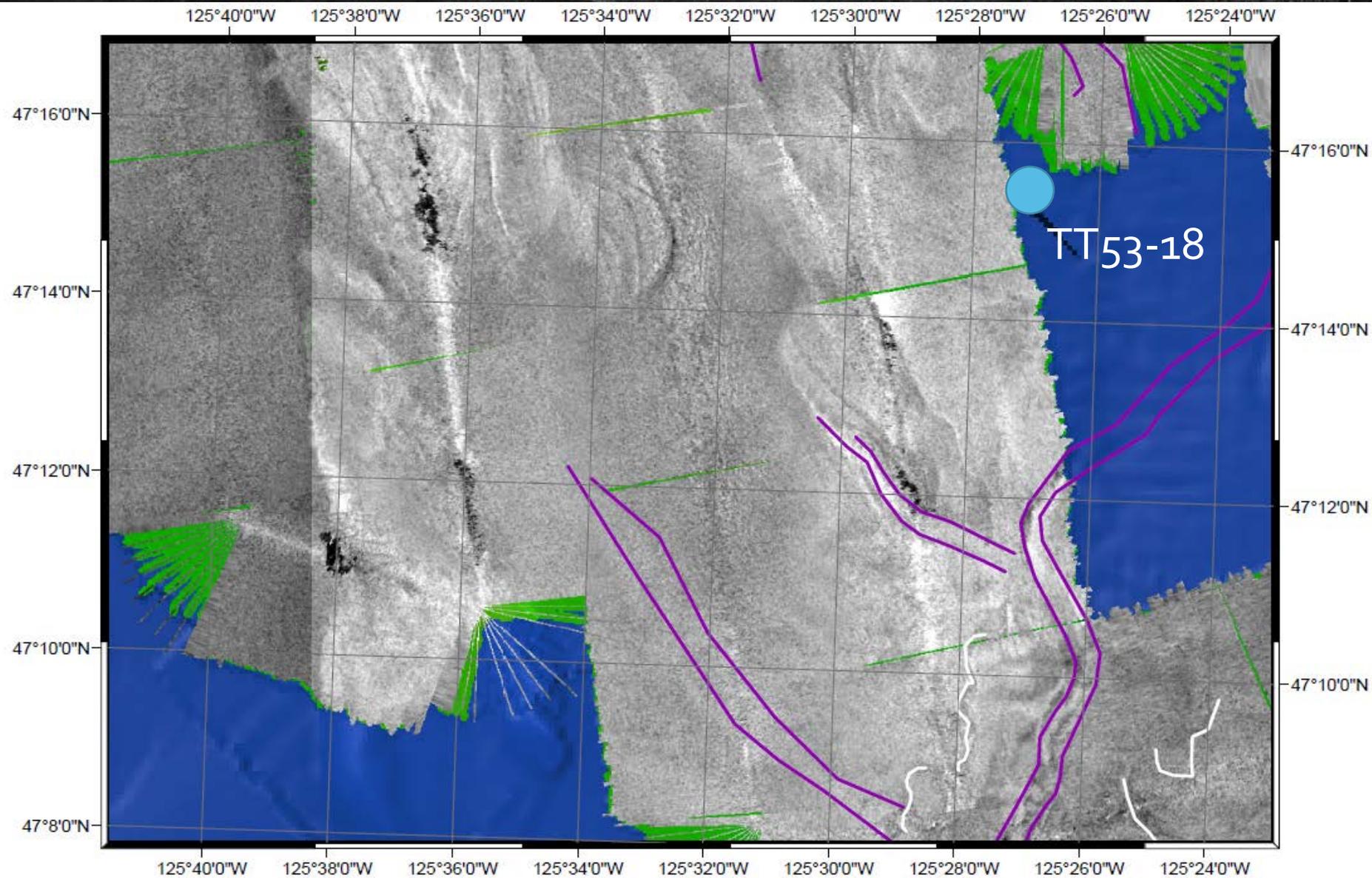
But in any case, not relevant to the confluence test as this notional input is below the key core sites.

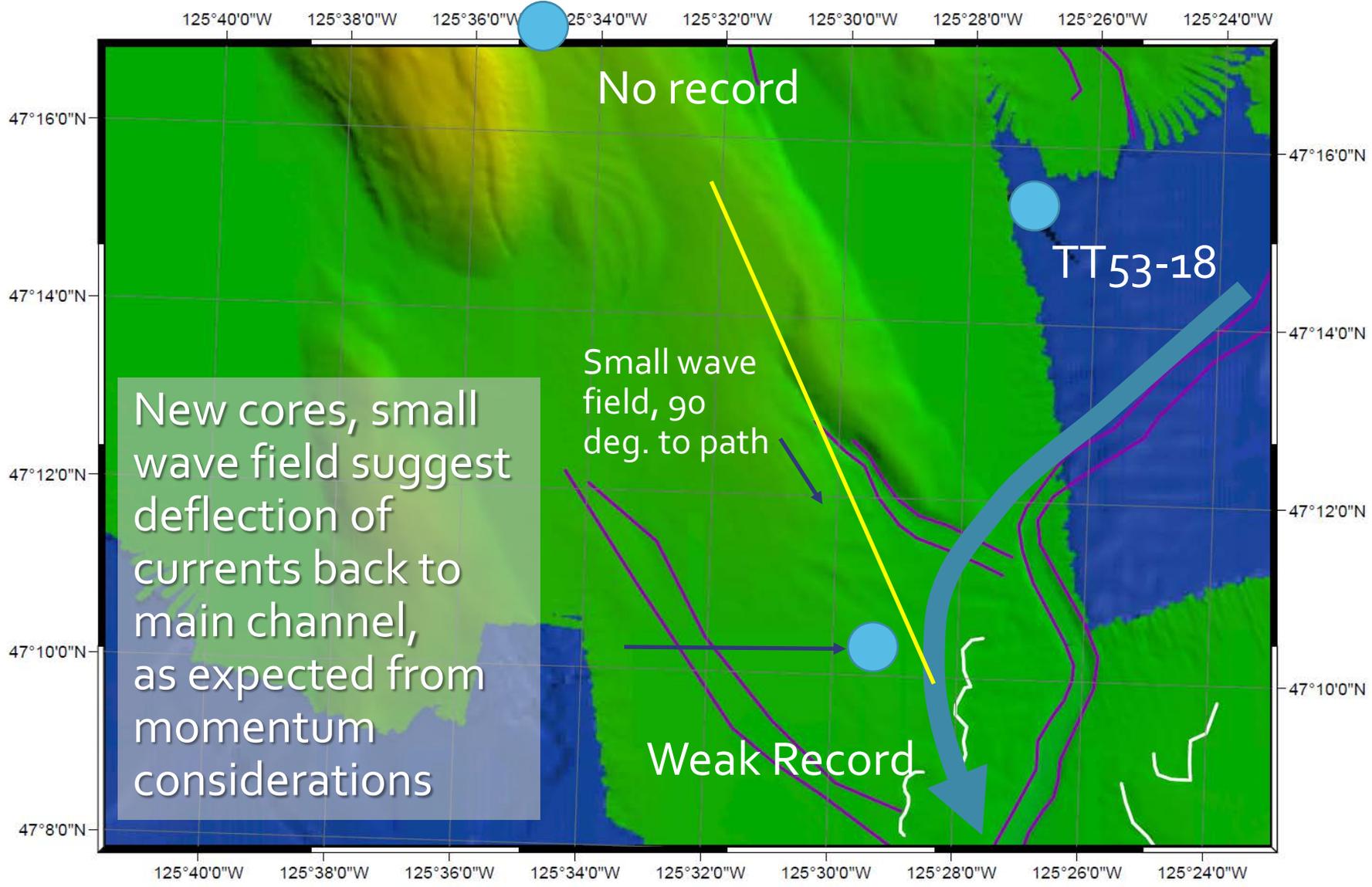


Cascadia

Upper slope canyons funnel turbidity currents into a margin parallel channel as wedge growth has made the upper slope/lower slope break the lowest point in the wedge.







No record

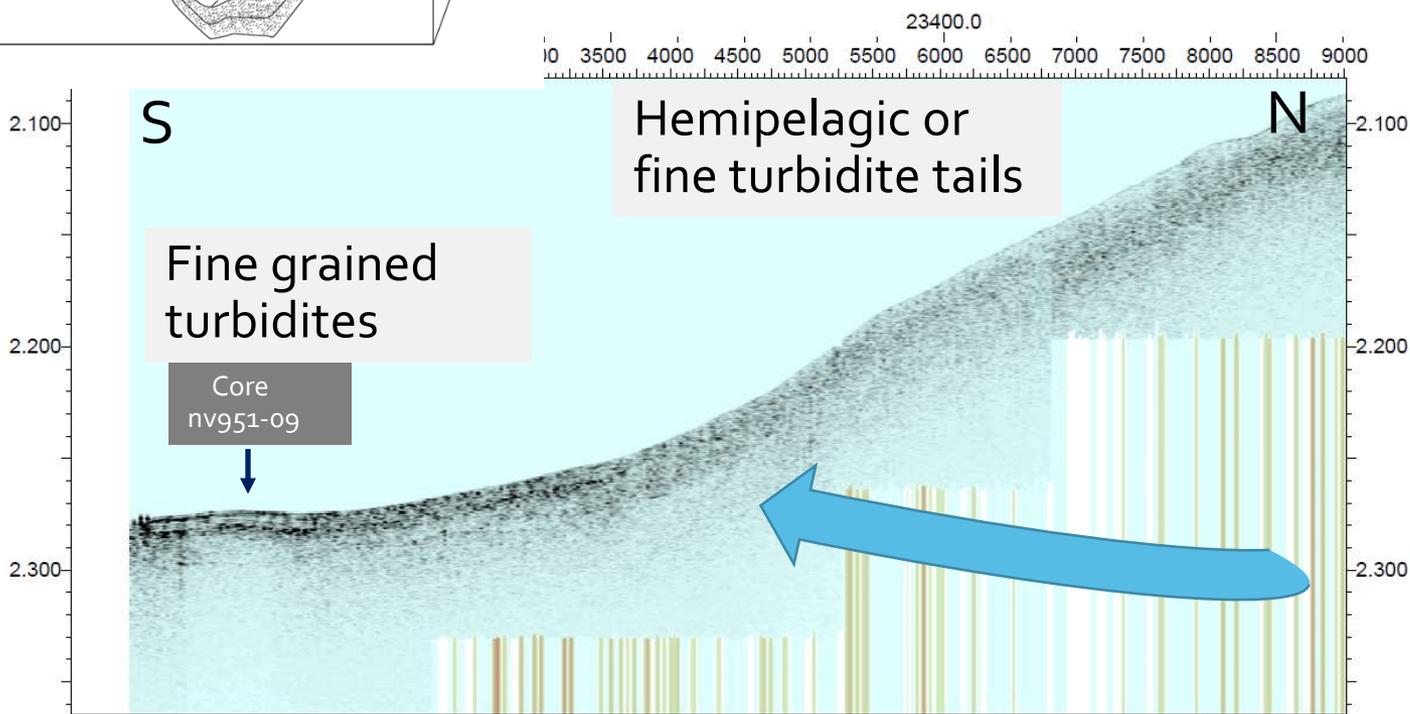
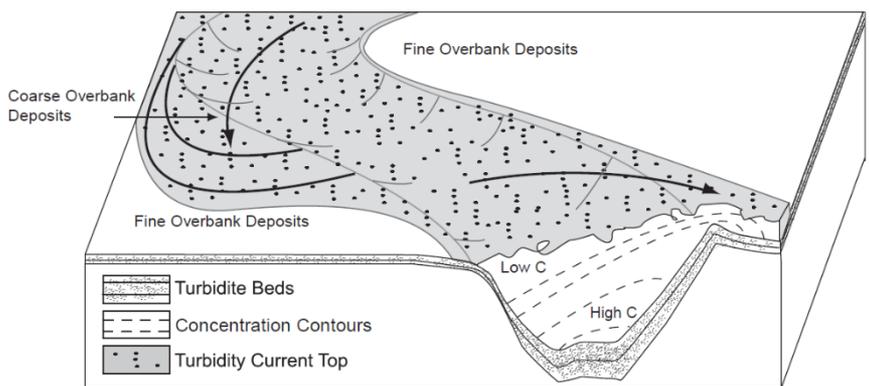
TT53-18

Small wave field, 90 deg. to path

Weak Record

New cores, small wave field suggest deflection of currents back to main channel, as expected from momentum considerations

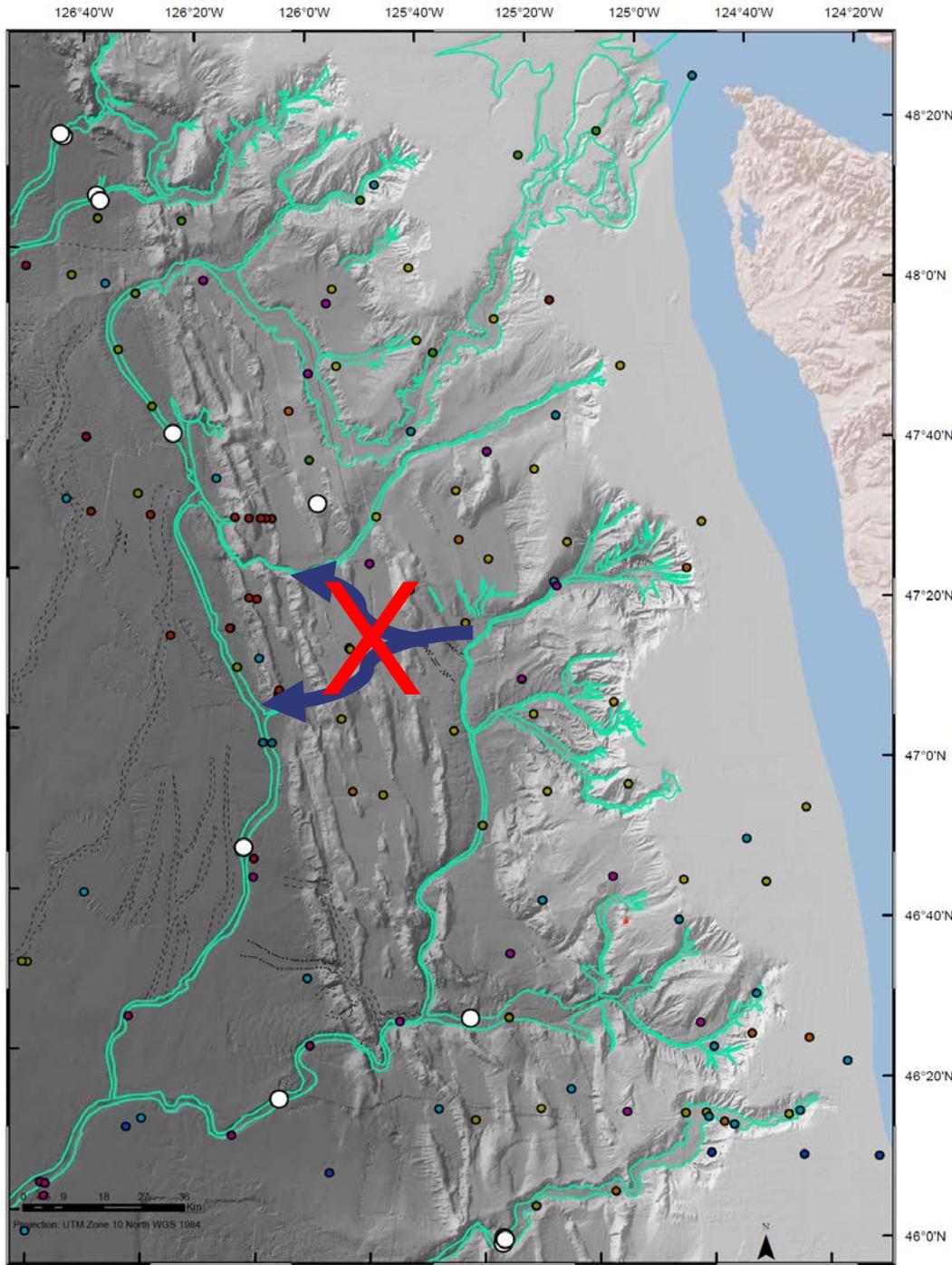
Classic overbank wavefield at a submarine channel bend



Seismic Micro-Technology, Inc.
 Project: West_Coast_CHIRP
 Project Location:
 Line: 0043_2011_151_2249_100212_CHP3.5_RAW_095_ED.96Y.NAV.
 Amplitudes [Filter ON 500, 1000, 2000, 2500]

Monte Carlo search inversion for best fit to core stratigraphy

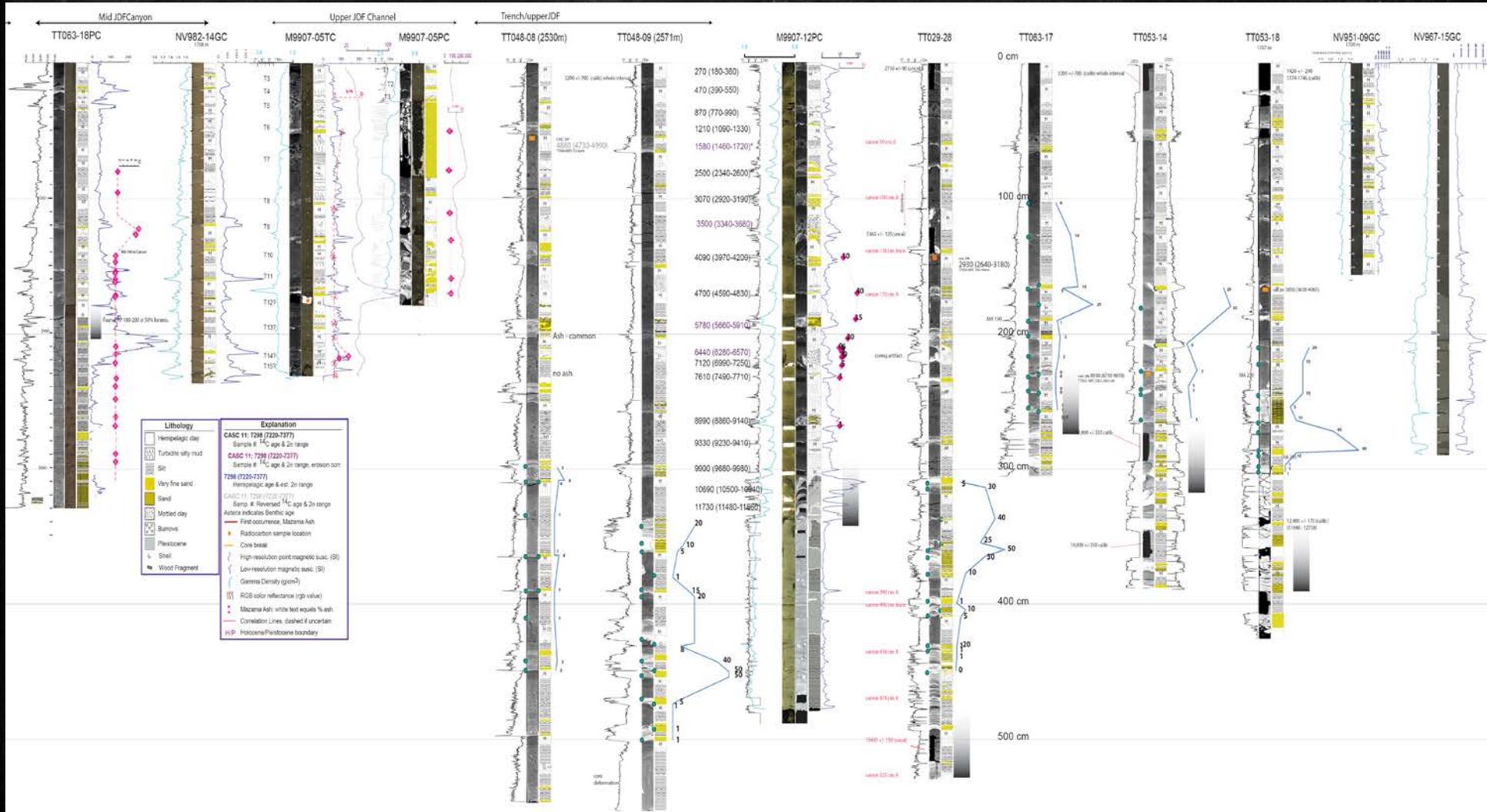




Cascadia

It might be possible to feed a little flow-stripped material into the Quillayute Basin.

Over the top of the accretionary prism and into the JDF Channel is not a possible pathway.



Thin sections
in upper
canyon

Expanded
above
obstruction

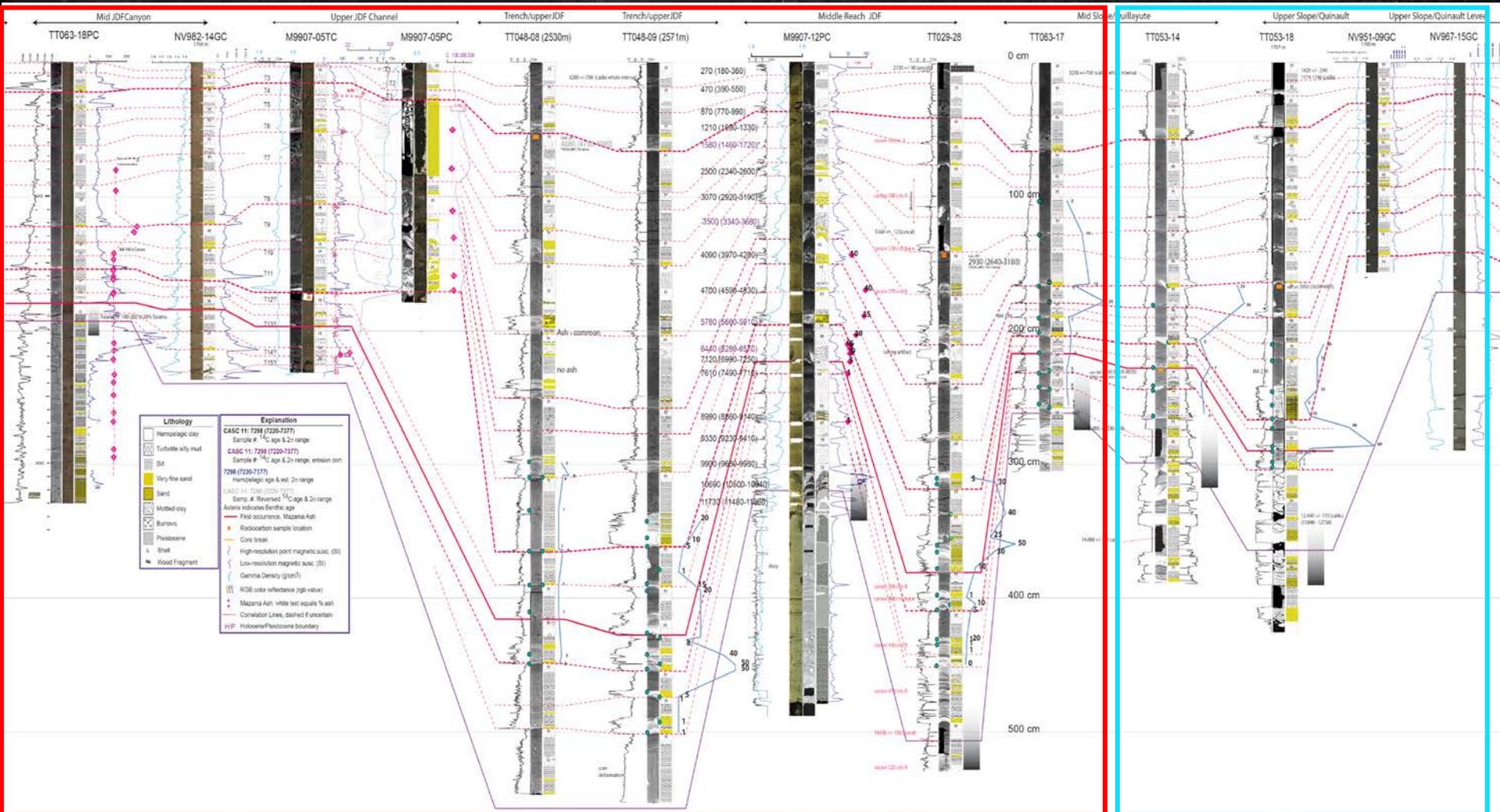
Compressed
below
obstruction

Thick sections
in lower
channel

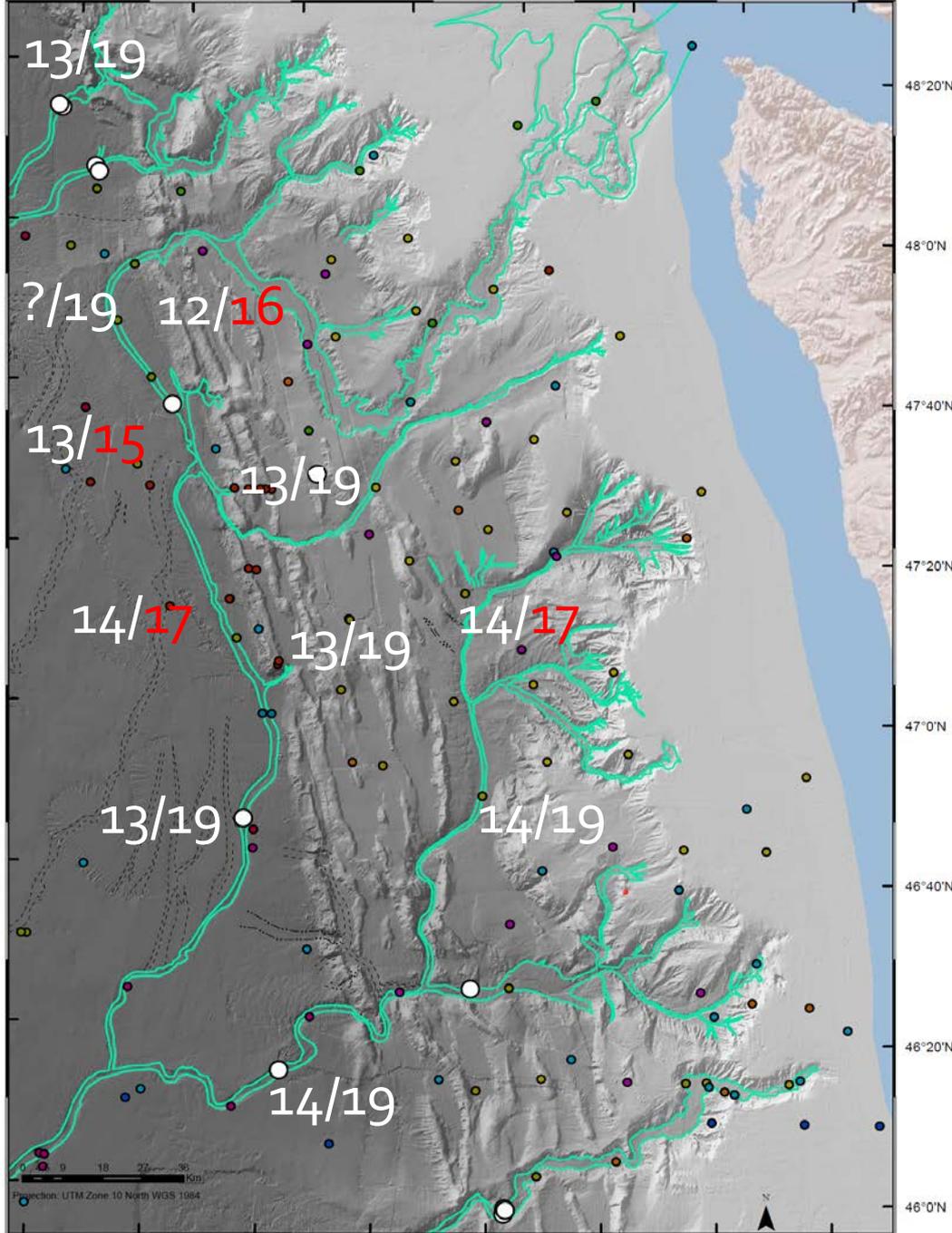
Thickening and coarsening downstream! Consistent with seismic trigger (but not required)

Western systems

Eastern systems



New expanded analysis suggests a rather simple signal of 19-20 Holocene turbidites in the along the northern margin in independent canyon/channel systems, supporting the original confluence test



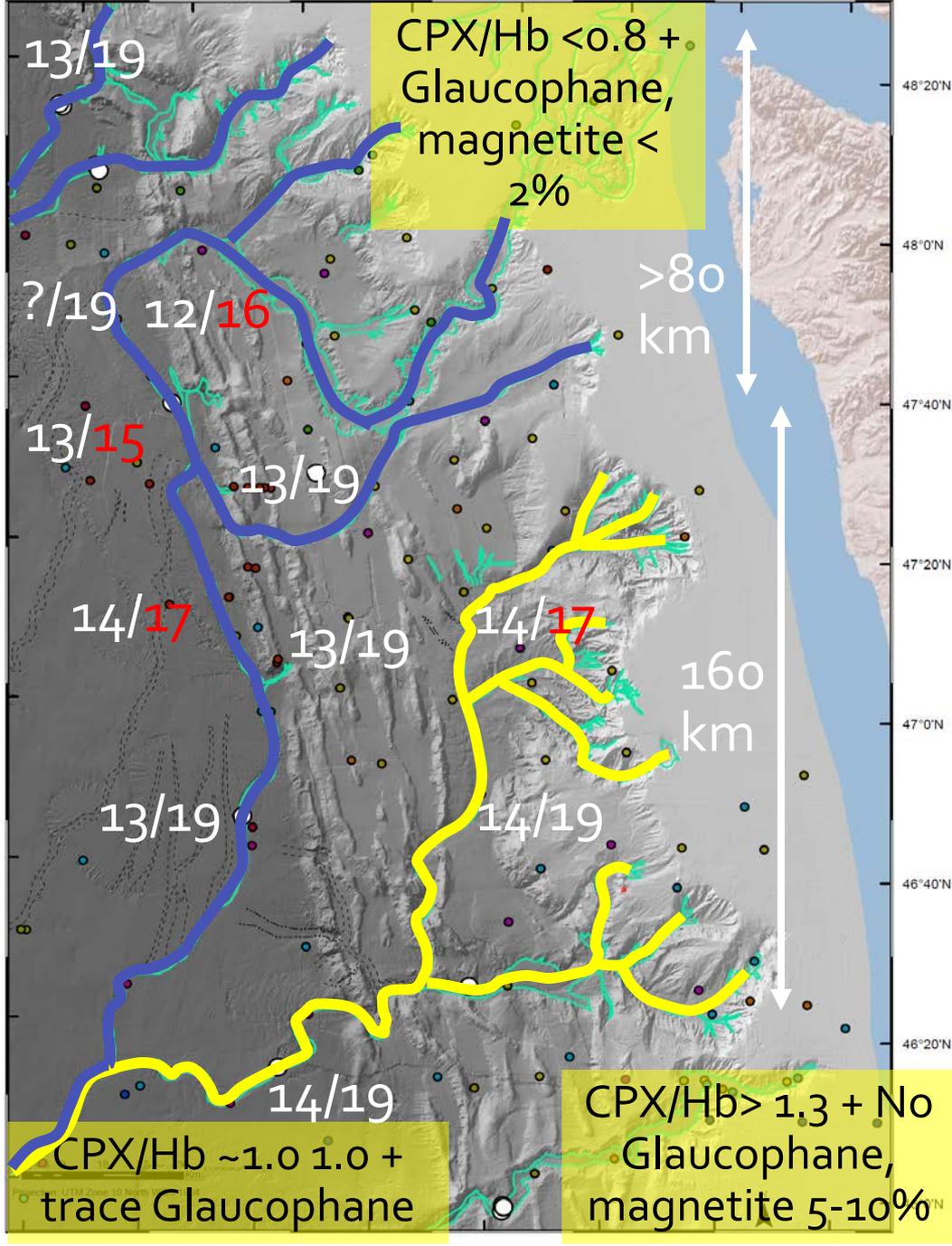
Two major independent Systems:

A consistent record for the most part, with occasional missing section at the Holocene-Pleistocene boundary.

First number is turbidites above the Mazama Ash, second number is Holocene turbidites, red if erosion detected.

Consistency of the Washington margin turbidite record across multiple turbidite systems

Core	H/P	Mazama	Correl. T13	visual logs	notes
JDF					
<i>TT63-18</i>	16	12	?	4/7	missing section at base of Holocene?
<i>TT63-20</i>	14	11	?	8-9 Holo.	CT analysis not complete
<i>TT39-27</i>	19	?	?	4 Holo.	No ash reported in Carson 1971, new analysis pending
<i>TT39-06</i>	20	?	13	2 Holo.	No ash reported in Carson 1971, new analysis pending
<i>TT48-09</i>	17	14	14	15 Holo.	HP not quite reached
<i>TT48-08</i>	17	14	14	15 Holo.	HP not quite reached
<i>M9907-05TC</i>	15	13	13	6	core overpenetrated? 2 events missing from top. H/P not reached
<i>M9907-12PC</i>	19	13	13	12/17	
<i>TT029-28</i>	18	14	14	10	
Quinault-Willapa					
<i>TT053-18</i>	18	16	14	13/15	possible missing section at base of Holocene
Quillayute					
<i>TT063-17</i>	18	16	14	5/11	possible missing section at base of Holocene
<i>TT053-14</i>	18	16	14	8/12	possible missing section at base of Holocene
Cascadia					
<i>M990722-25PC</i>	19	13	13	12/18	

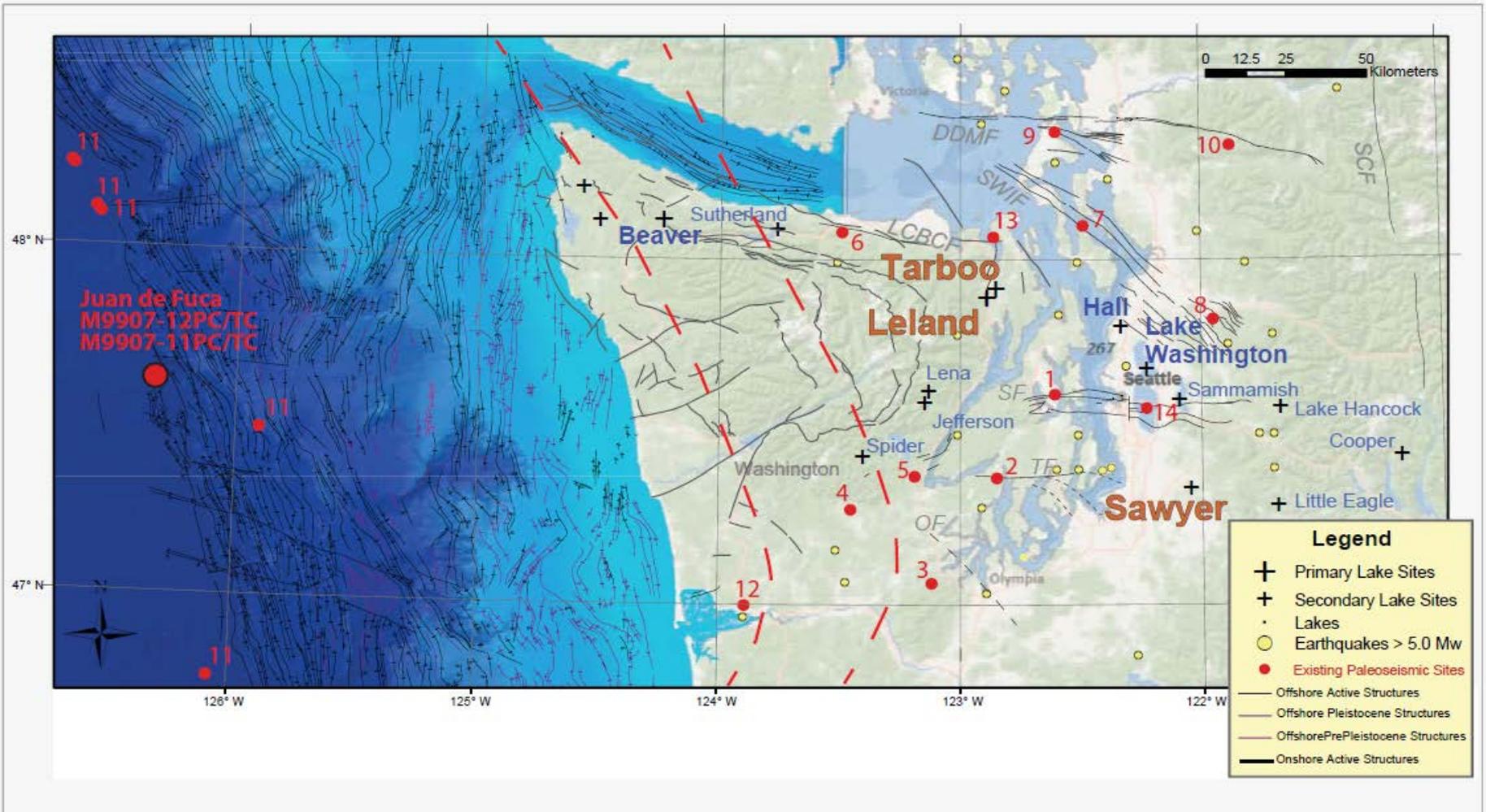


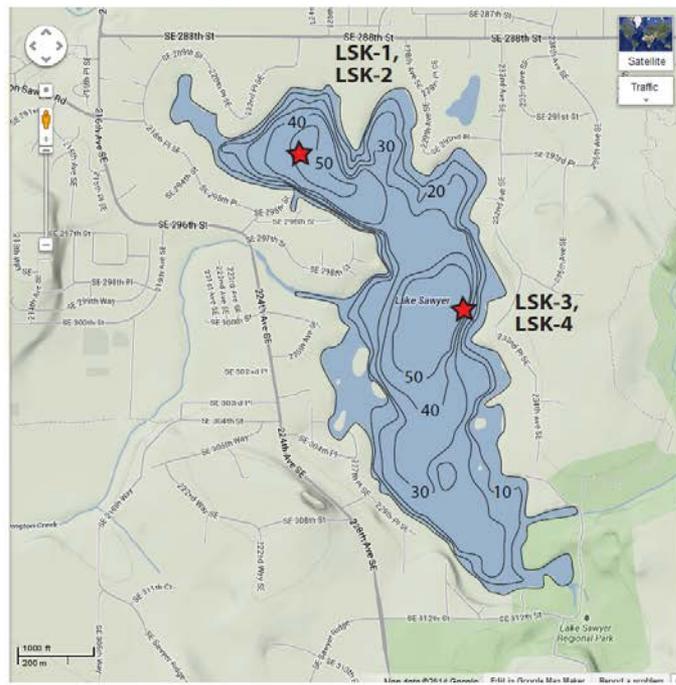
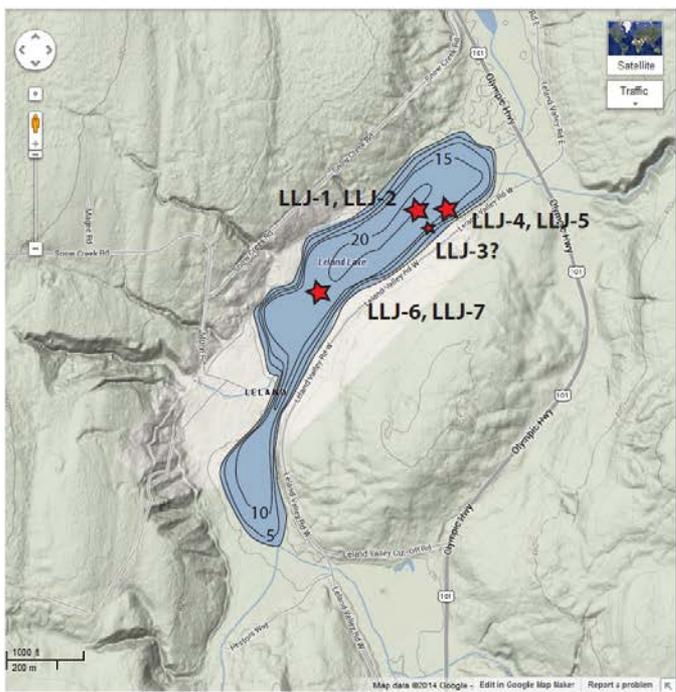
The confluence:

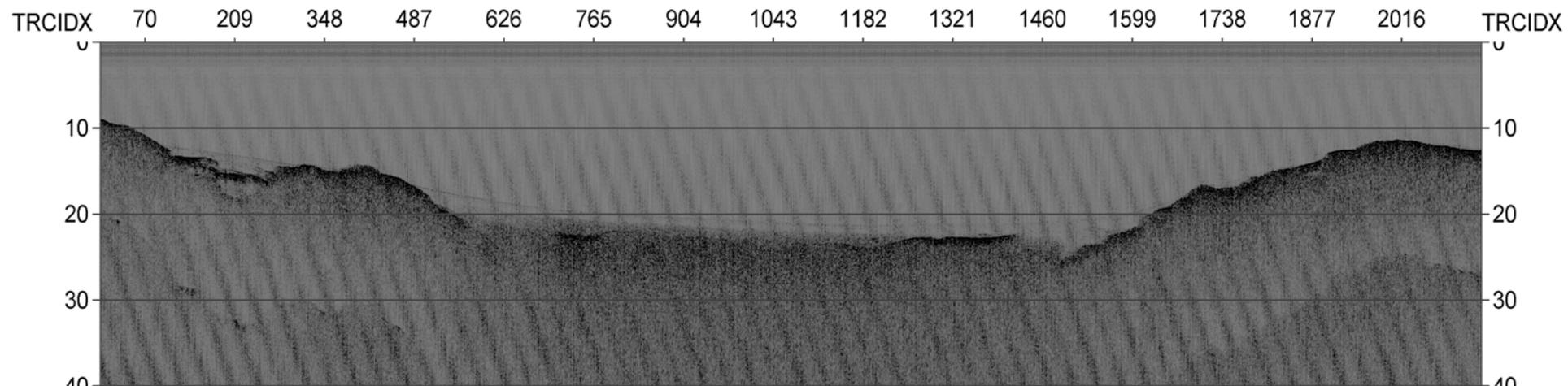
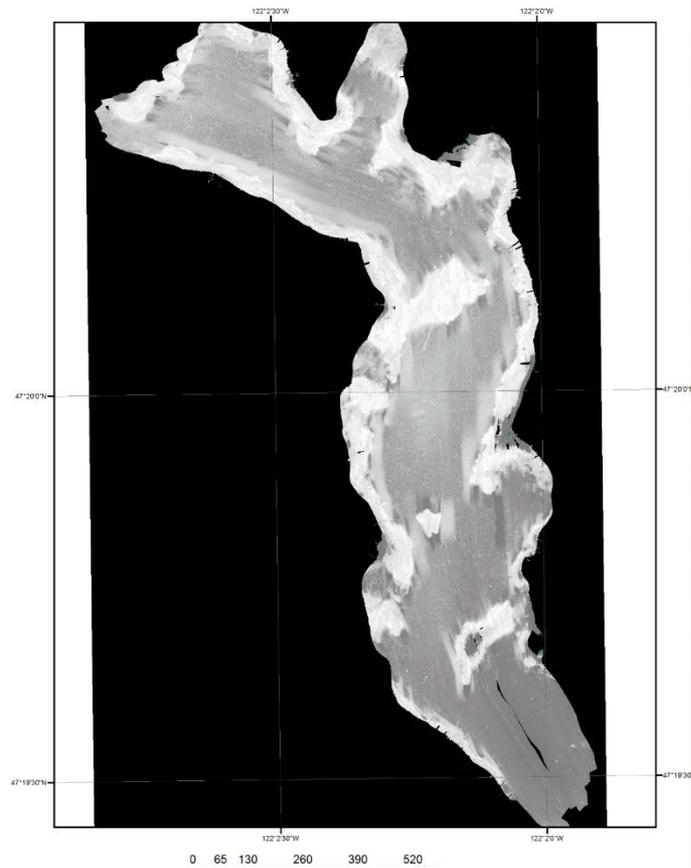
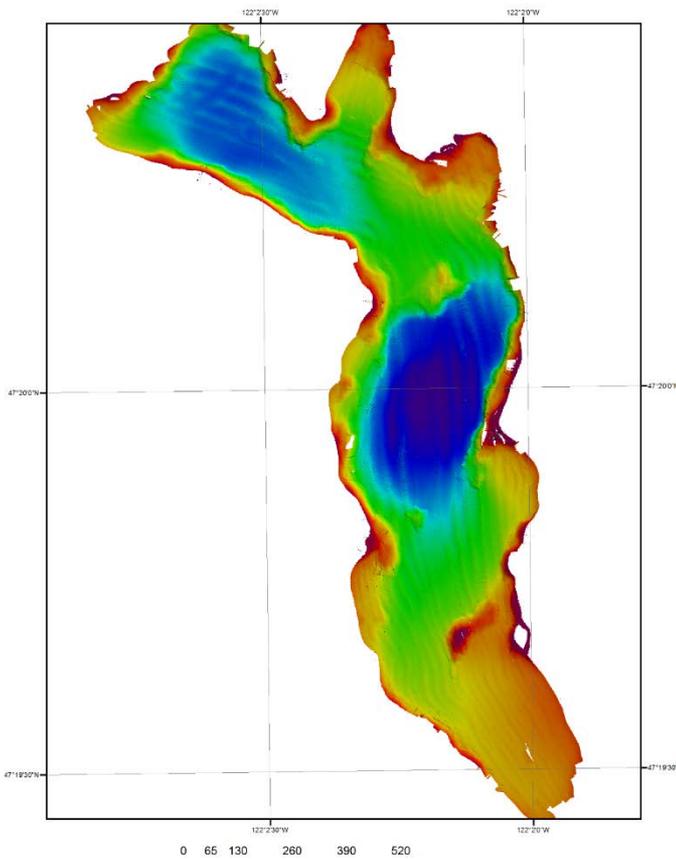
A mixed signal.

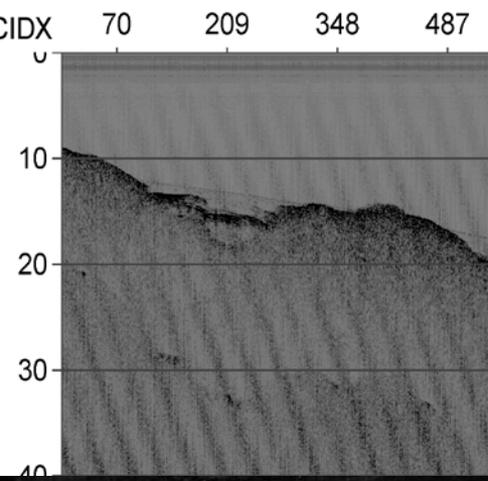
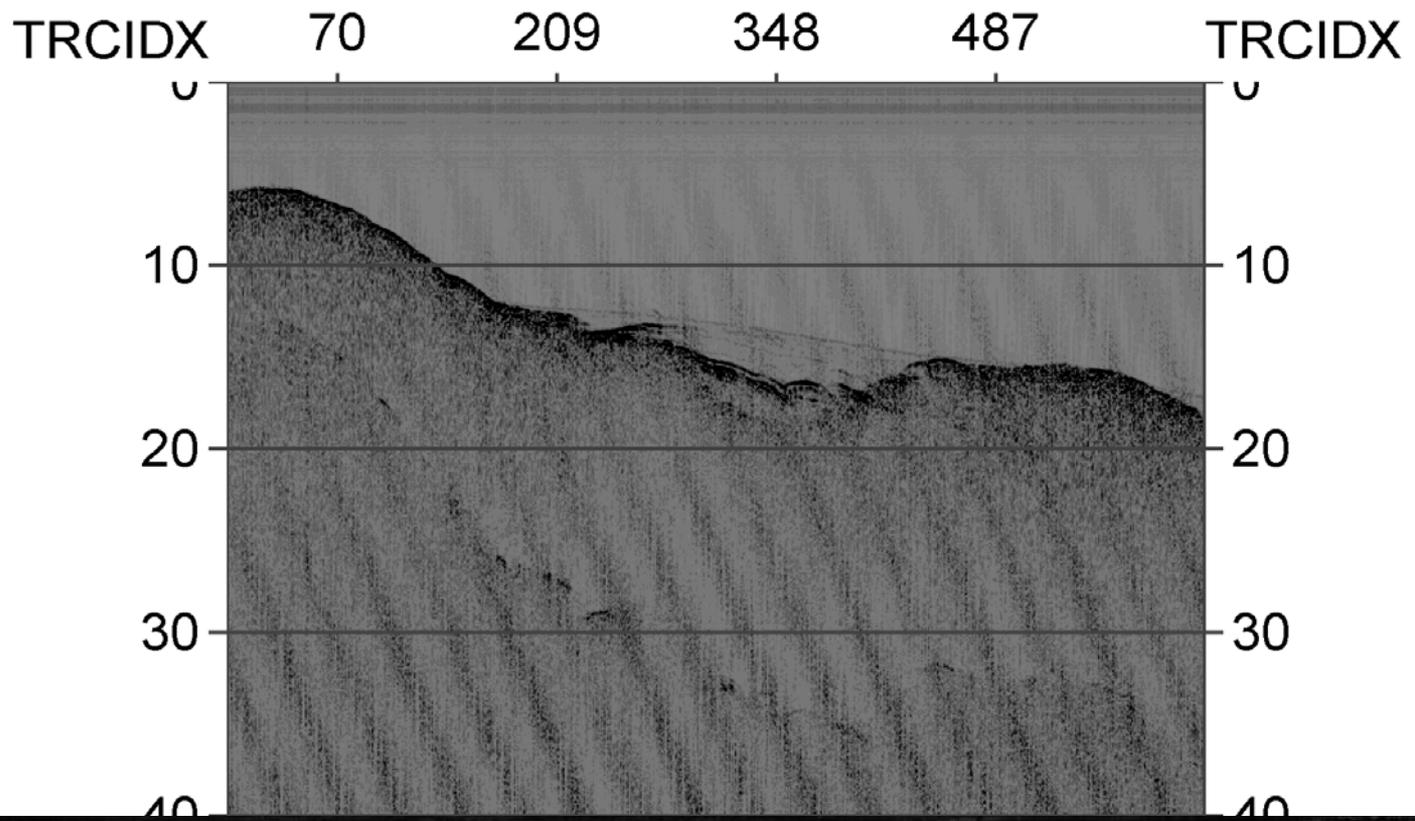
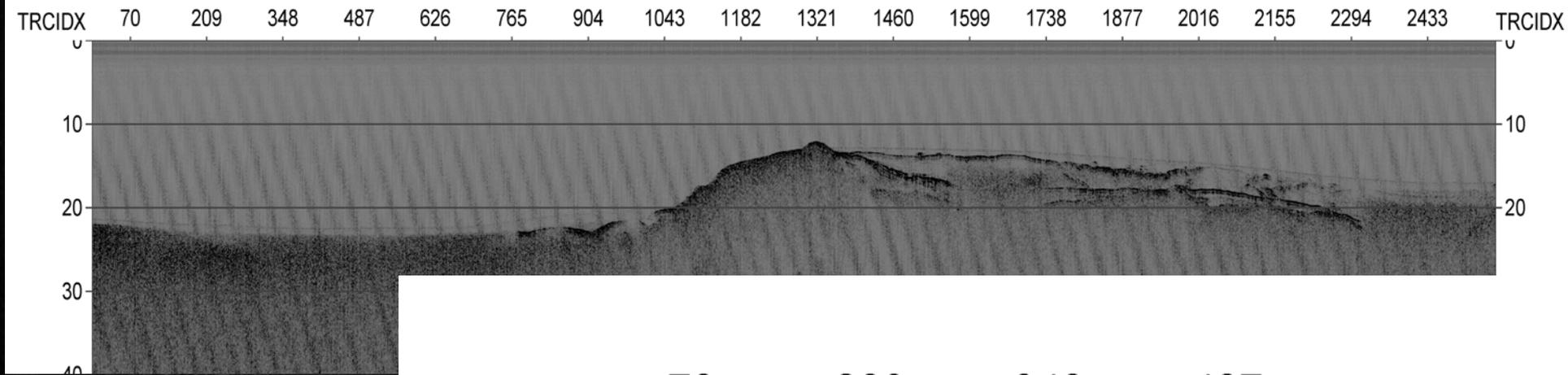
Heavy mineral tracers distinguish northern and southern sources and preclude significant input to JDF from Quinault Canyon.

Direct comparison between offshore and onshore evidence in forearc Lakes.









Marine Core

W

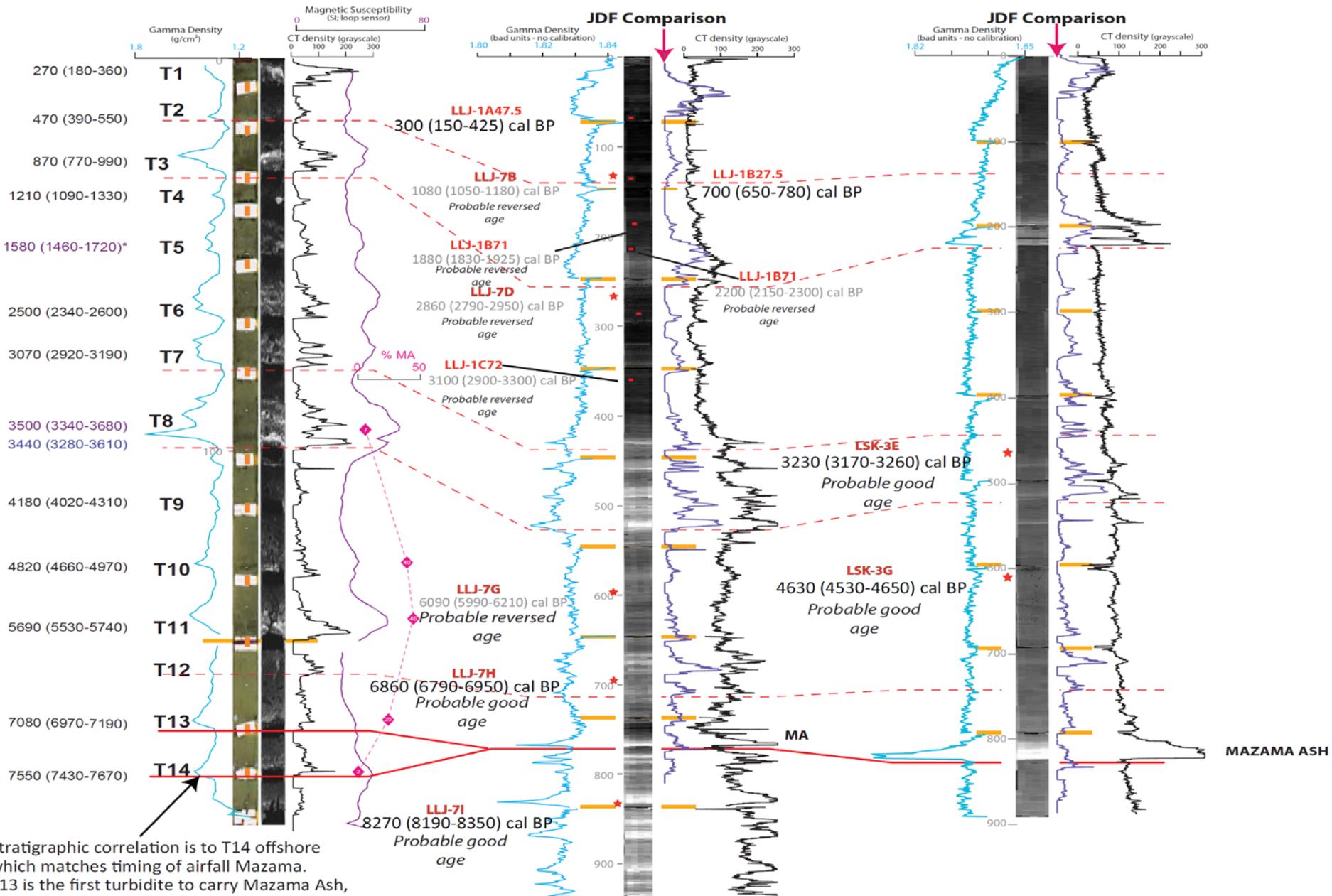
Forearc Lakes

E

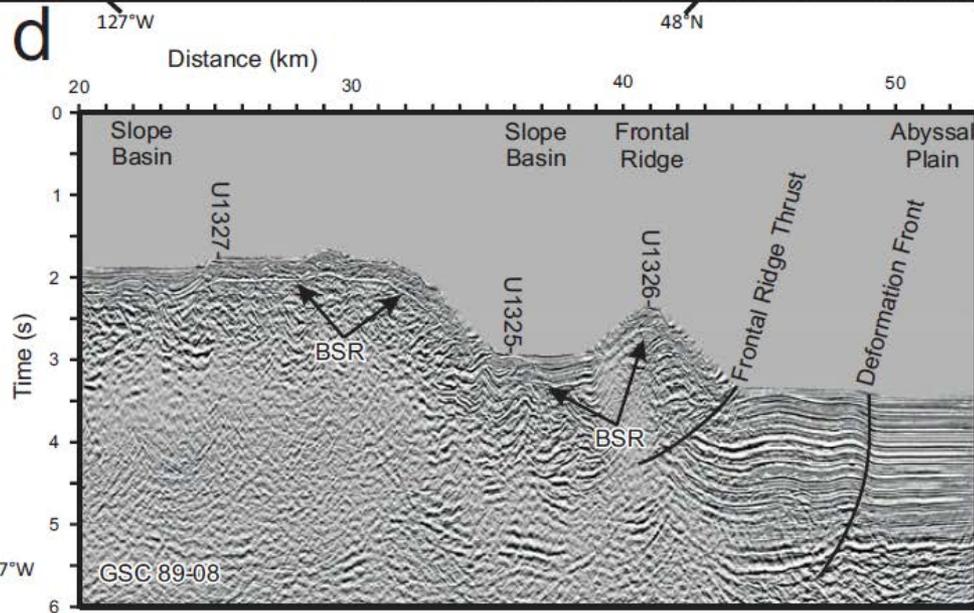
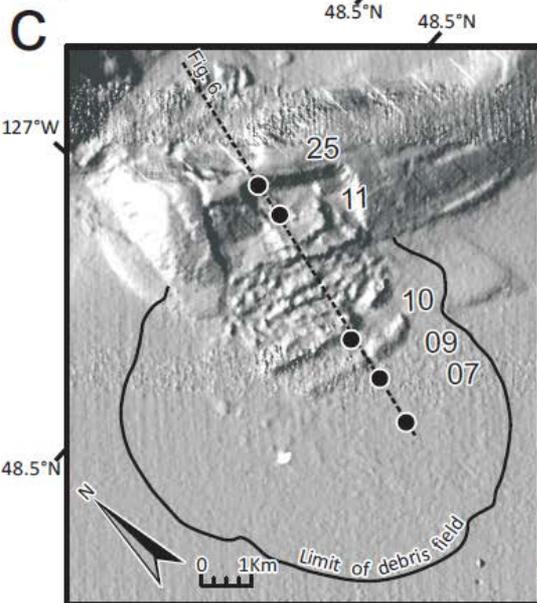
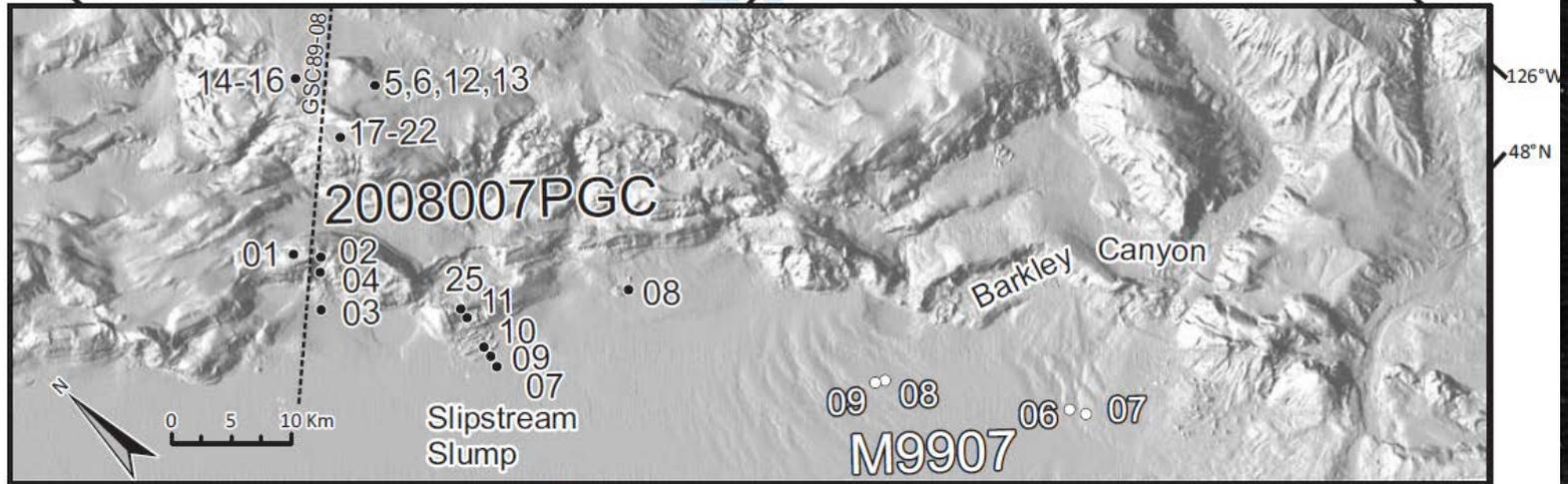
Juan de Fuca (JDF) Channel M9907-11TC

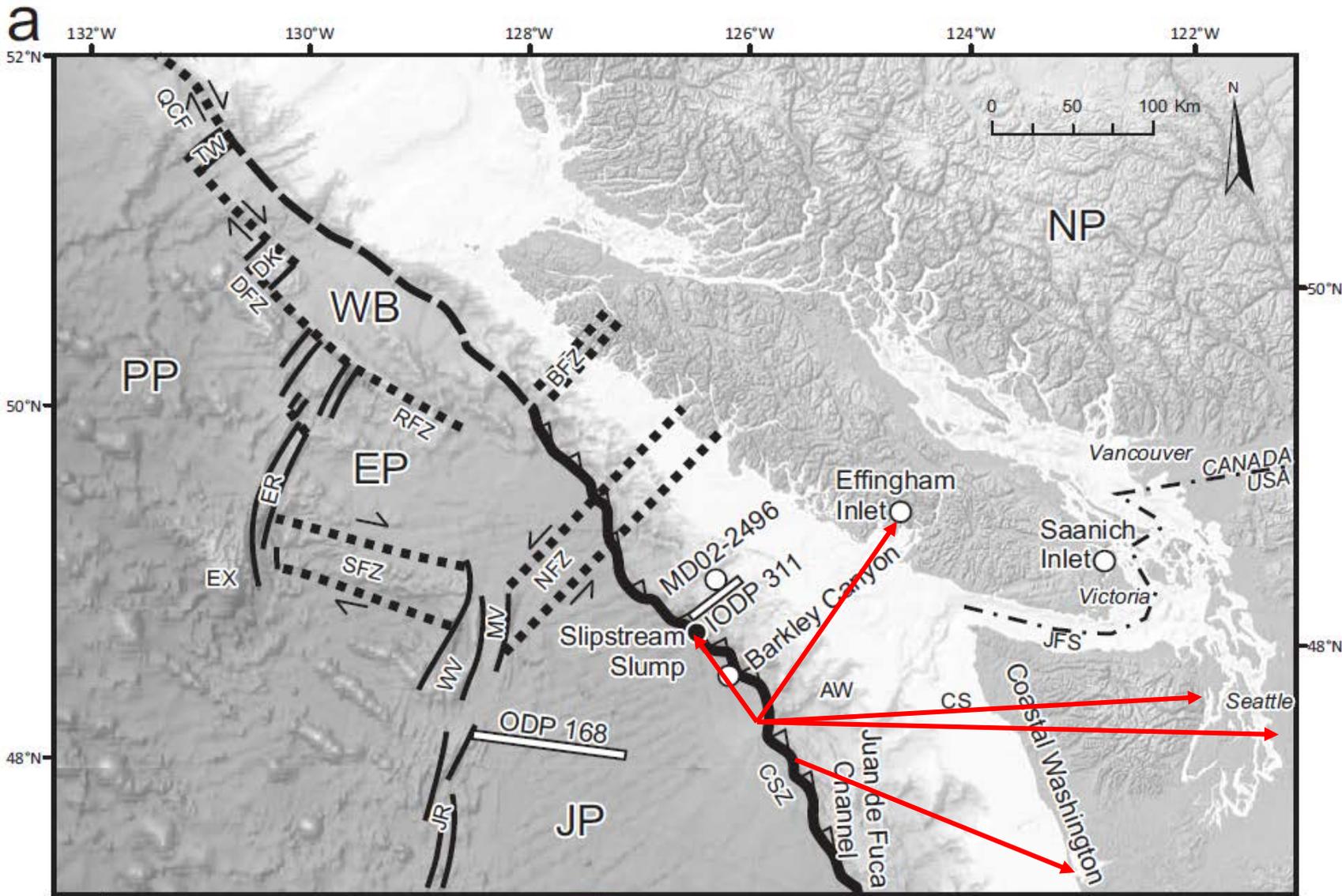
Leland Lake, Jefferson Co. WA LLJ-7

Lake Sawyer, King Co. WA LSK-3



Slipstream slide: Hamilton et al. 2014





Land Data

- Deserted Lake²⁰
- ✕ Port Alberni¹¹
- ✕ Tofino^{7,17,18}
- ✕ Effingham^{11,45}
- ✕ Catala Lake⁹
- △ Kakawis Lake²⁵
- ✕ Saanich Inlet⁶
- ✕ Saanich Varves⁶
- Discovery Bay⁴¹
- ✕ Swantown⁴⁰
- Cultus Bay⁴⁶
- ◇ Copalis River²
- Johns River³⁸
- ✕ Willapa Bay^{2,3,47}
- Long Beach WA³⁷
- Ecola Creek^{14,34}
- ✕ Ecola 2007⁴⁴
- ◇ Netarts Shennan^{39,24}
- Netarts Marsh^{12,13}
- △ Salmon River²⁹
- Yaquina Bay¹⁴
- ◇ Alsea Bay^{14,31}
- Coquille River⁴³
- Coos Bay^{230,27,28}
- ✕ Bradley Lake²³
- Sixes River⁴³
- Humboldt Bay³³
- ✕ Eel River³³
- ◇ Lagoon Creek^{15,16}

Onshore-Offshore space-time diagram for the most recent ~ 2800 years.

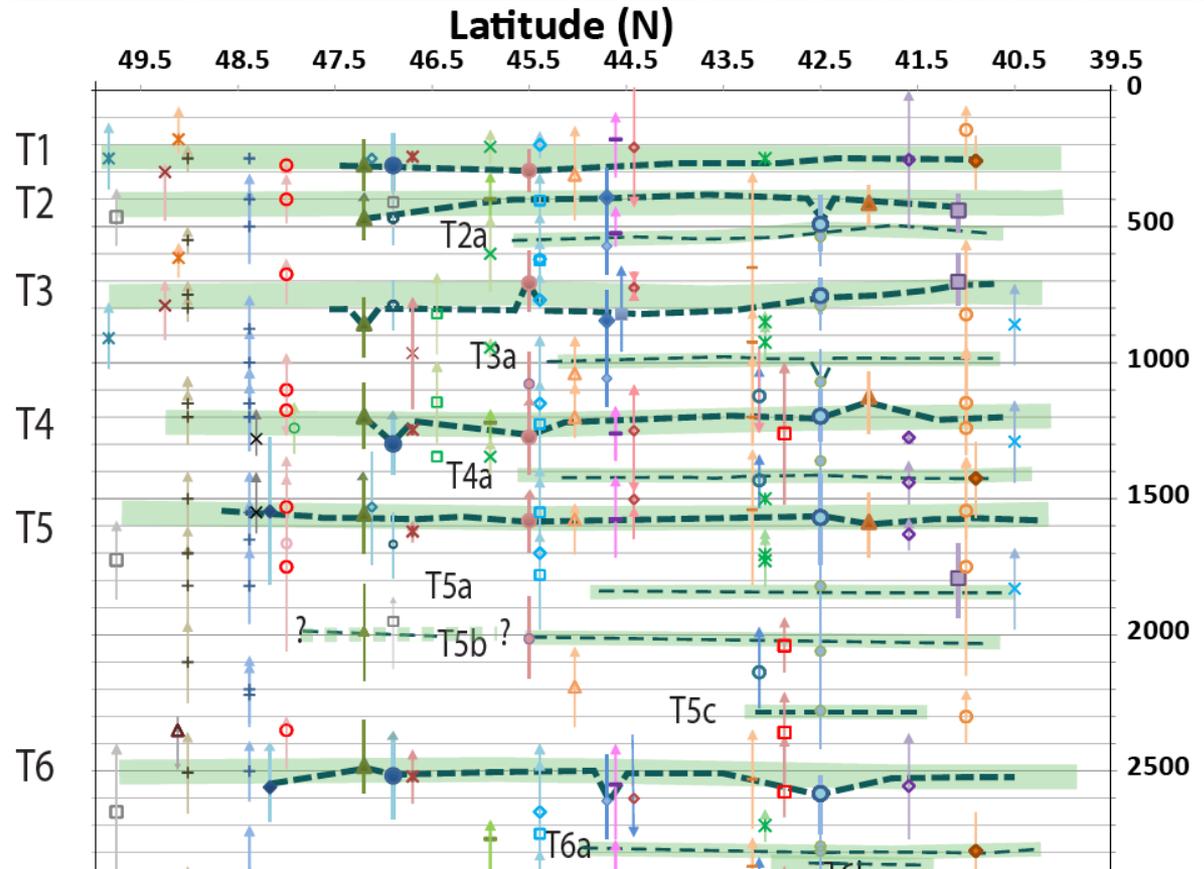
(Filled symbols are marine data, open symbols land data; smaller open symbols are bulk peat ages, given lower weighting here.)

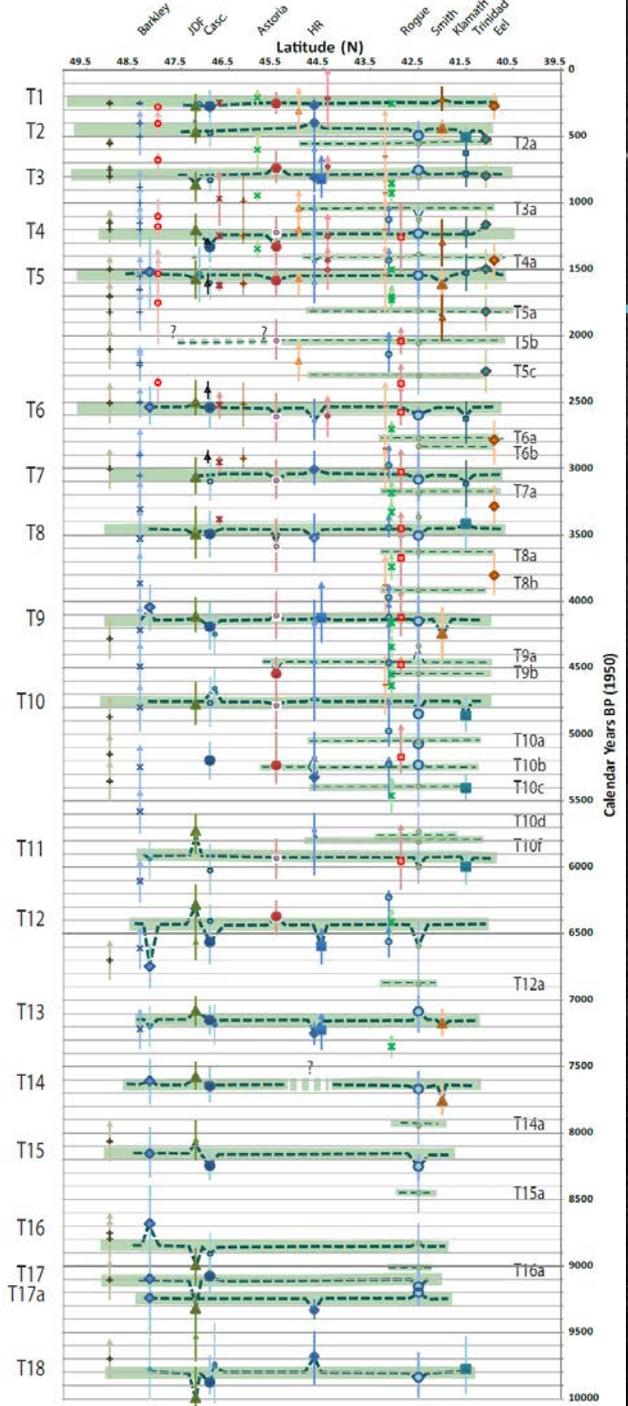
Stratigraphic correlation for offshore data shown in blue dashed lines.

Marine Data

- ◇ Barclay Canyon
- ◇ Barkley Canyon H
- △ Juan de Fuca
- △ Juan de Fuca H
- Cascadia Channel
- Cascadia Channel H
- Cascadia 1996⁴⁸
- Astoria Channel
- Astoria Channel H
- Astoria 1996⁴⁸
- ◇ Hydrate Ridge
- ◇ Hydrate Ridge H
- Rogue Apron
- Rogue Apron H
- △ Smith Apron
- Klamath Canyon
- Trinidad Plunge Pool
- ◇ Eel Channel

Land Data





Onshore-Offshore space-time diagram including only higher precision land ages (Filled symbols are marine data, open symbols land data; smaller open symbols are bulk peat ages, given lower weighting here.)

Stratigraphic correlation for offshore data shown in blue dashed lines.

And finally, back to the model.

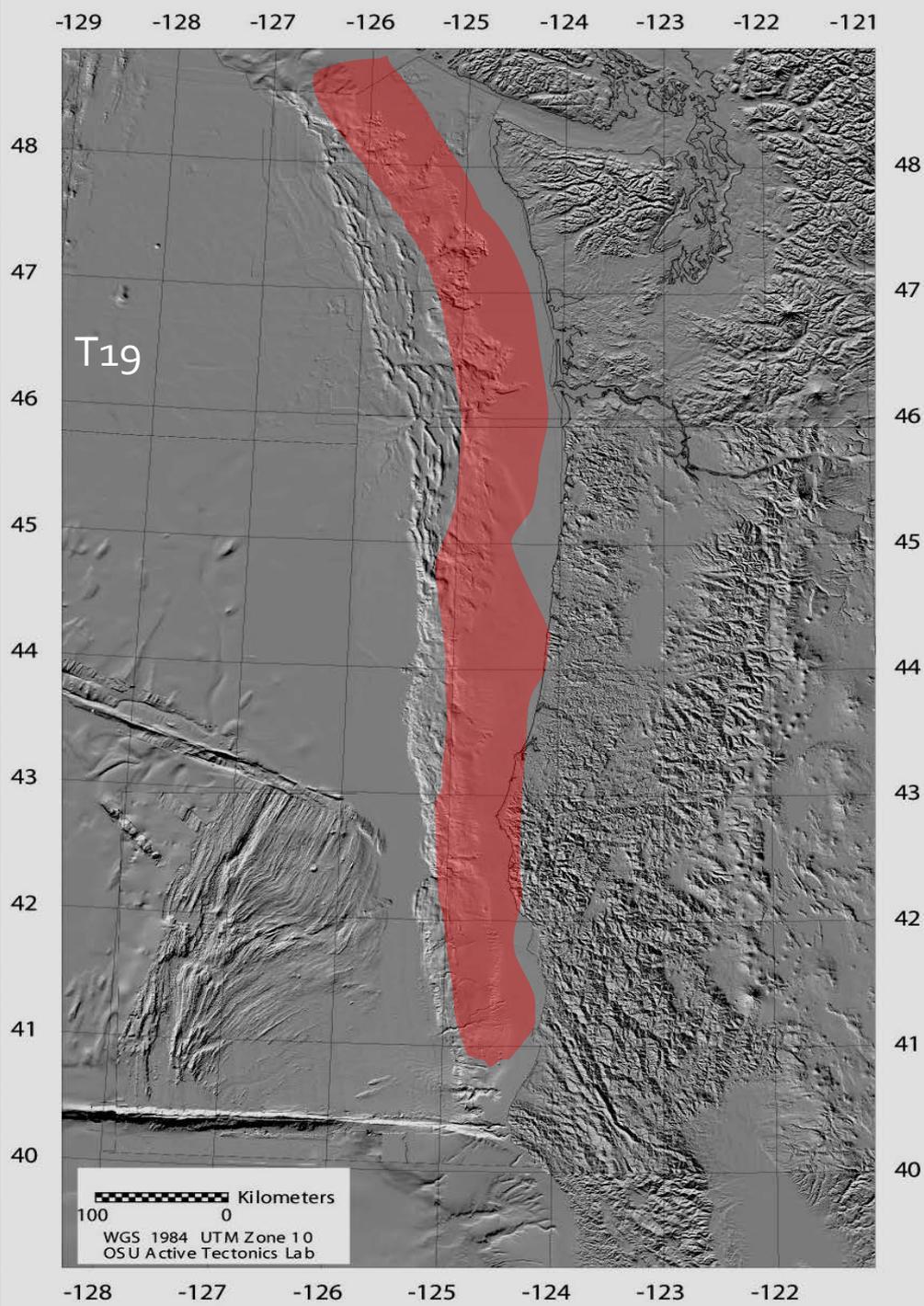
The inputs are quantified where possible according to this scheme:

- OxCal χ^2 and agreement indices as measures of radiocarbon fit, land and marine data.
- Pearson correlation coeff. for downcore series.
- 85% probability assigned (+/- 15) for confluence and coincidence tests based on all relevant data.
- Boolean (yes/no) for turbidite filter criteria.
- Binning of strike length and number of sites observed.
- Seismic correlation criteria not applicable to northern margin (topography too rough).
- Hydrodynamic model fit RMS can be used with caution.
- The model is in development, and the following should be considered experimental.

Preliminary results: probability of correlation given the input data, T1-18 for (JDF, Cascadia, HR, Rogue) and high precision land sites.

Idnum	freq		probabilities of Correl	P(case)
1	1	*	(0.94 0.06)	4.17444e-008
2	1	*	(0.77 0.23)	4.17444e-008
3	1	*	(0.82 0.18)	4.30094e-008
4	1	*	(0.88 0.12)	4.17444e-008
5	1	*	(0.71 0.29)	4.43127e-008
6	1	*	(0.93 0.07)	4.17444e-008
7	1	*	(0.87 0.13)	4.30094e-008
8	1	*	(0.73 0.27)	4.30094e-008
9	1	*	(0.83 0.17)	4.17444e-008
10	1	*	(0.63 0.37)	4.30094e-008
11	1	*	(0.85 0.15)	4.17444e-008
12	1	*	(0.74 0.26)	4.17444e-008
13	1	*	(0.89 0.11)	4.17444e-008
14	1	*	(0.82 0.18)	4.30094e-008
15	1	*	(0.77 0.23)	4.30094e-008
16	1	*	(0.85 0.15)	4.17444e-008
17	1	*	(0.83 0.17)	4.17444e-008
17a	1	*	(0.81 0.19)	4.30094e-008
18	1	*	(0.79 0.21)	4.17444e-008

Average probability long ruptures= 81%, southern ruptures 64%



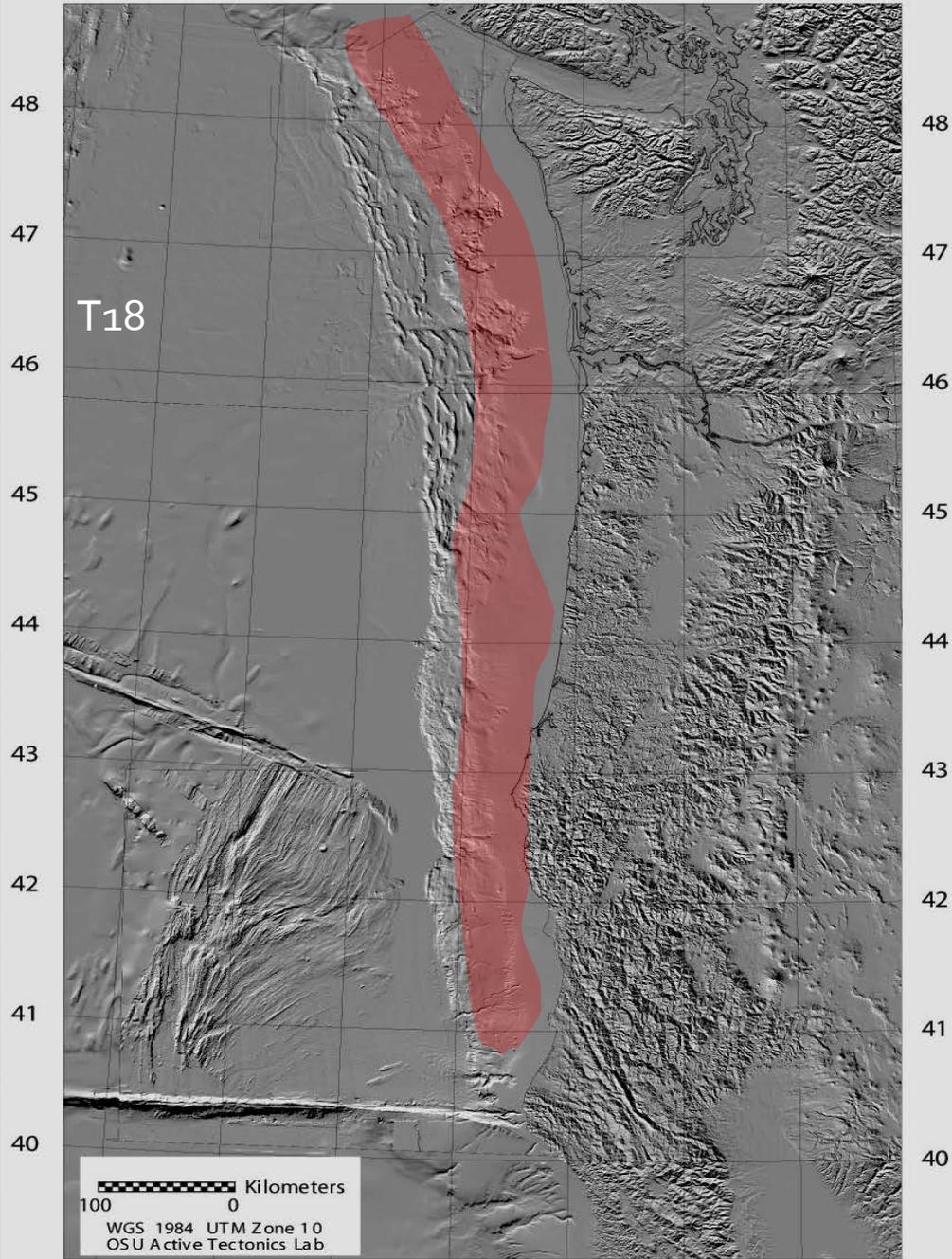
Cascadia: The Movie

This sequence shows the Cascadia Holocene earthquake sequence.

The slides are timed at 1 sec ~ 200 years.

Event pulses that correlate at all sites are shown by flashes of the "locked zone" in red. Event "size" shown by intensity of red shading

-129 -128 -127 -126 -125 -124 -123 -122 -121



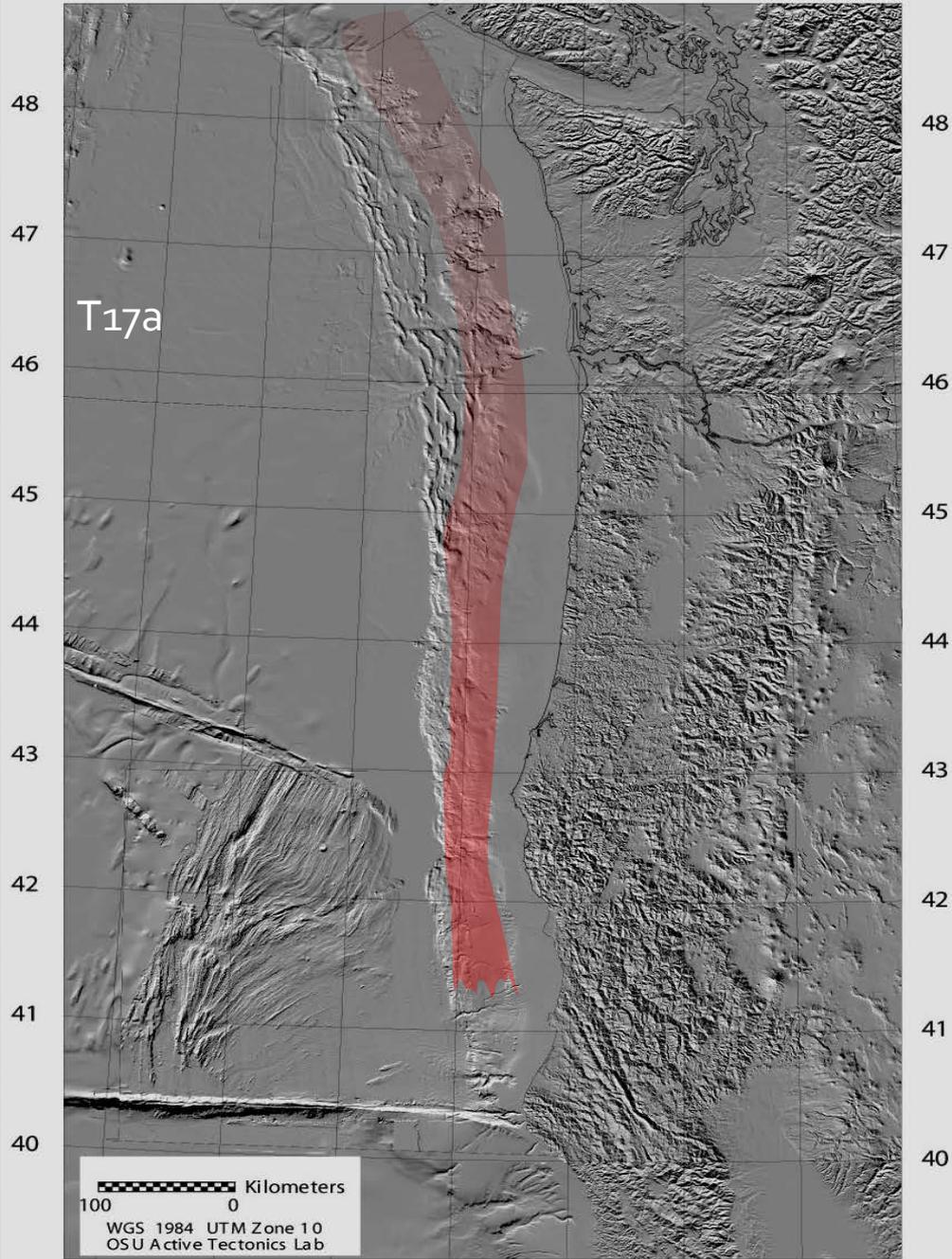
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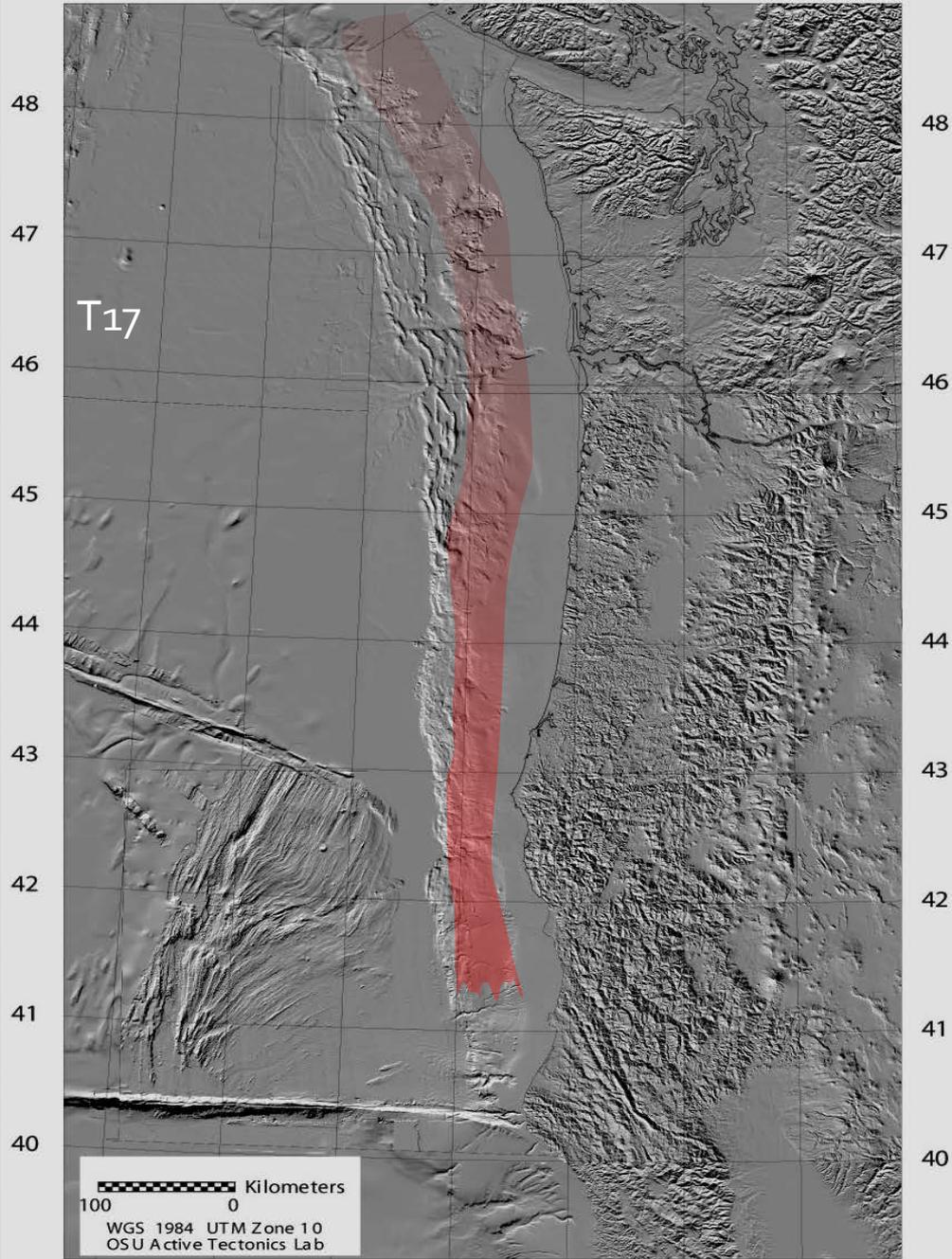
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

-129 -128 -127 -126 -125 -124 -123 -122 -121



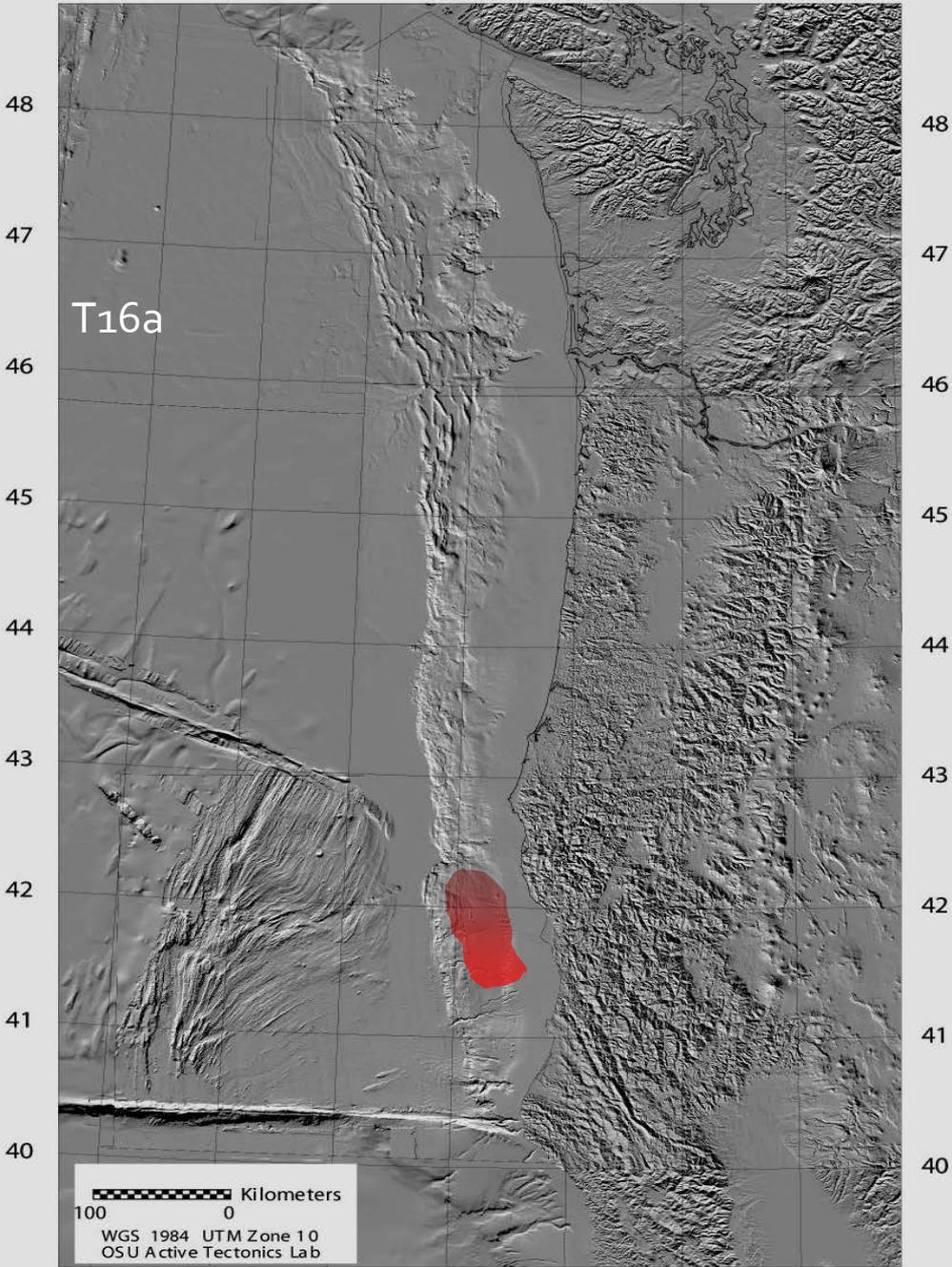
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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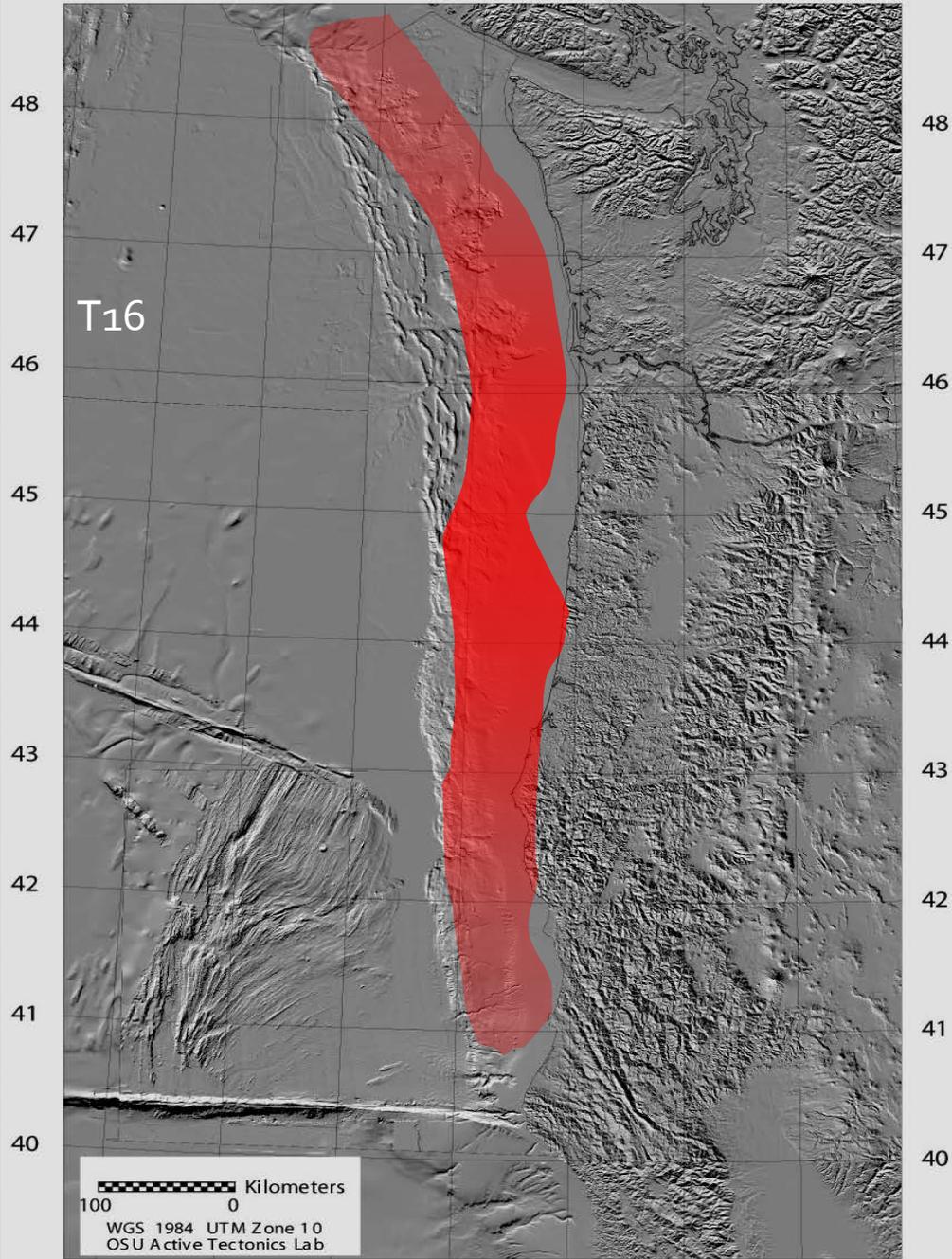


T16a

100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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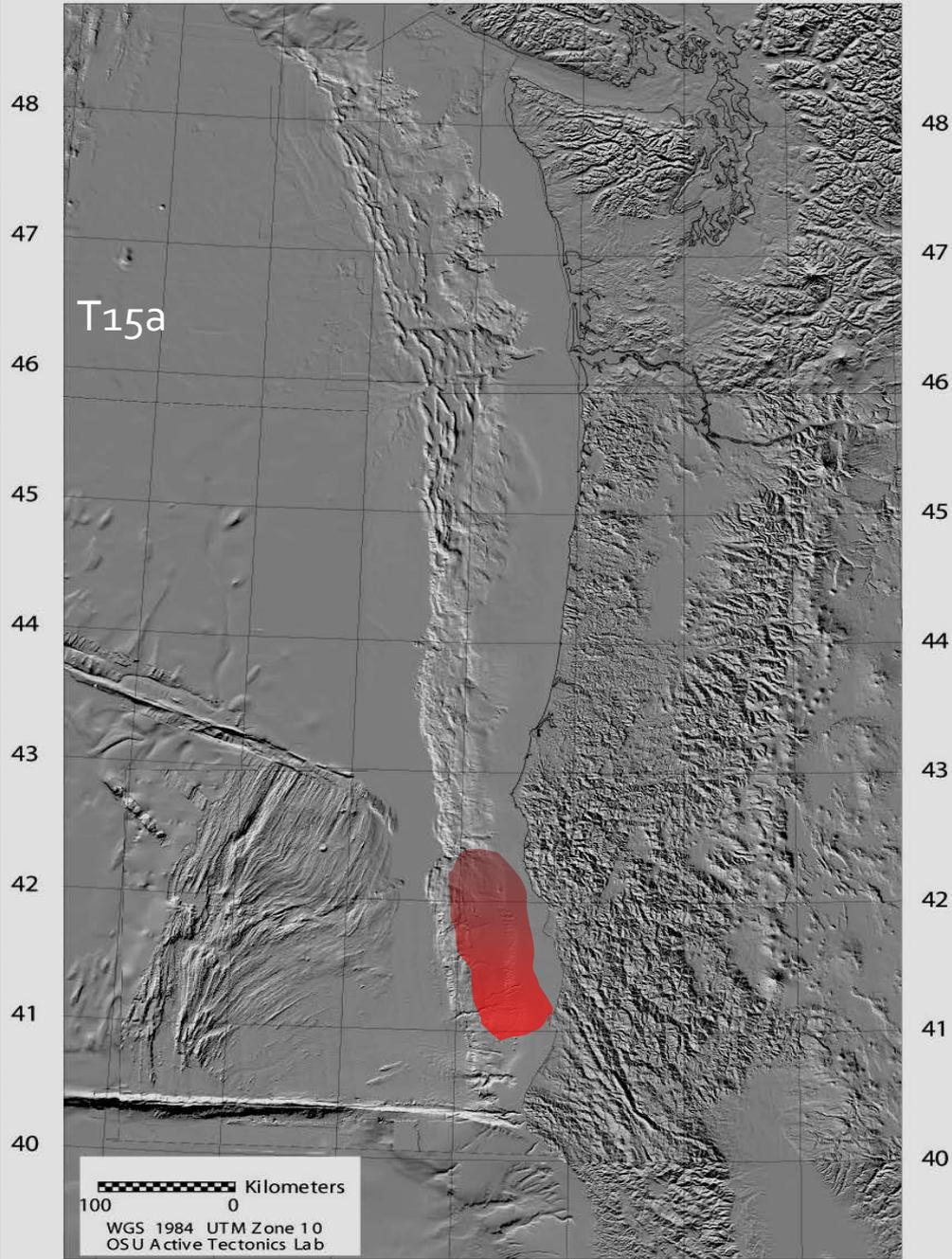
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WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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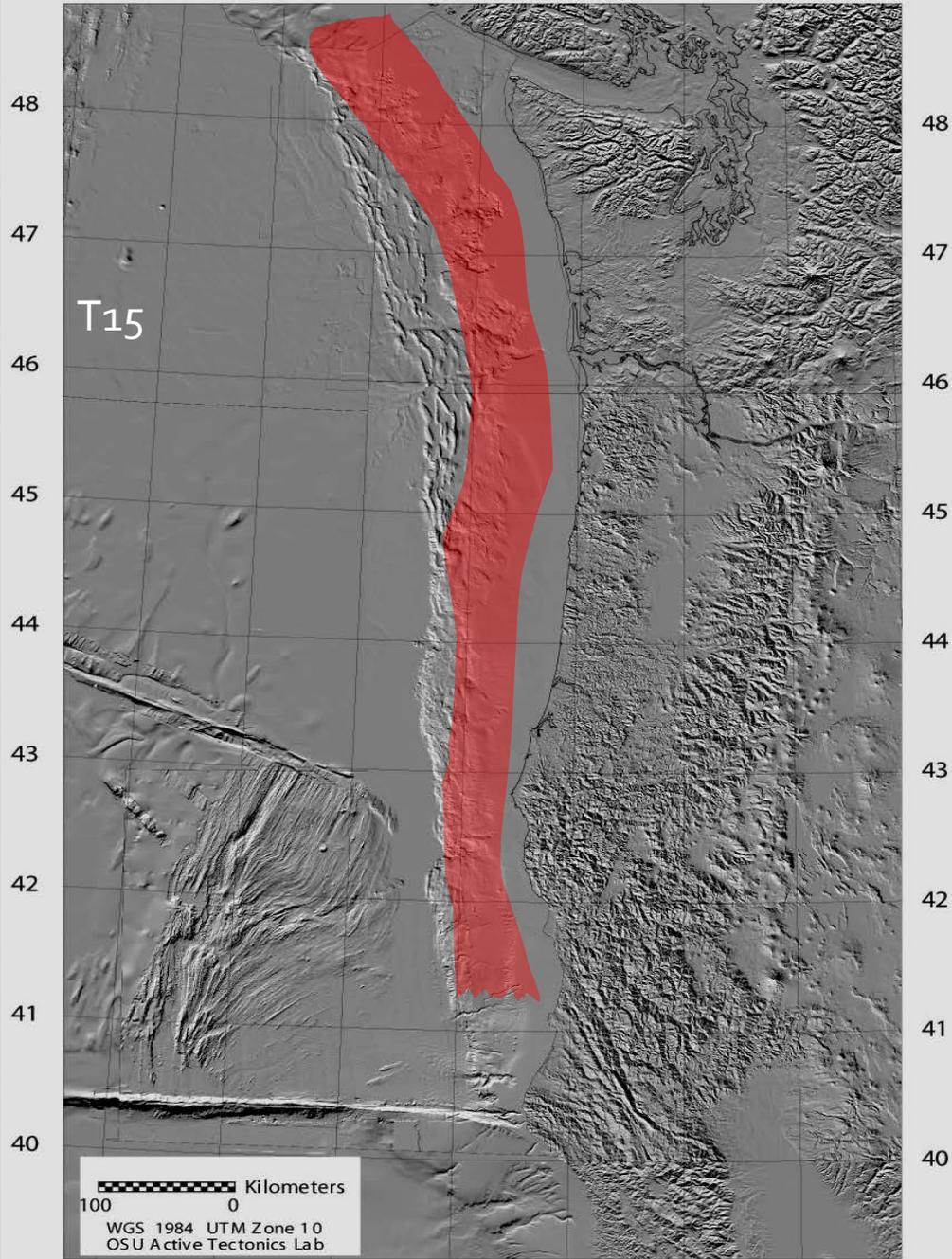


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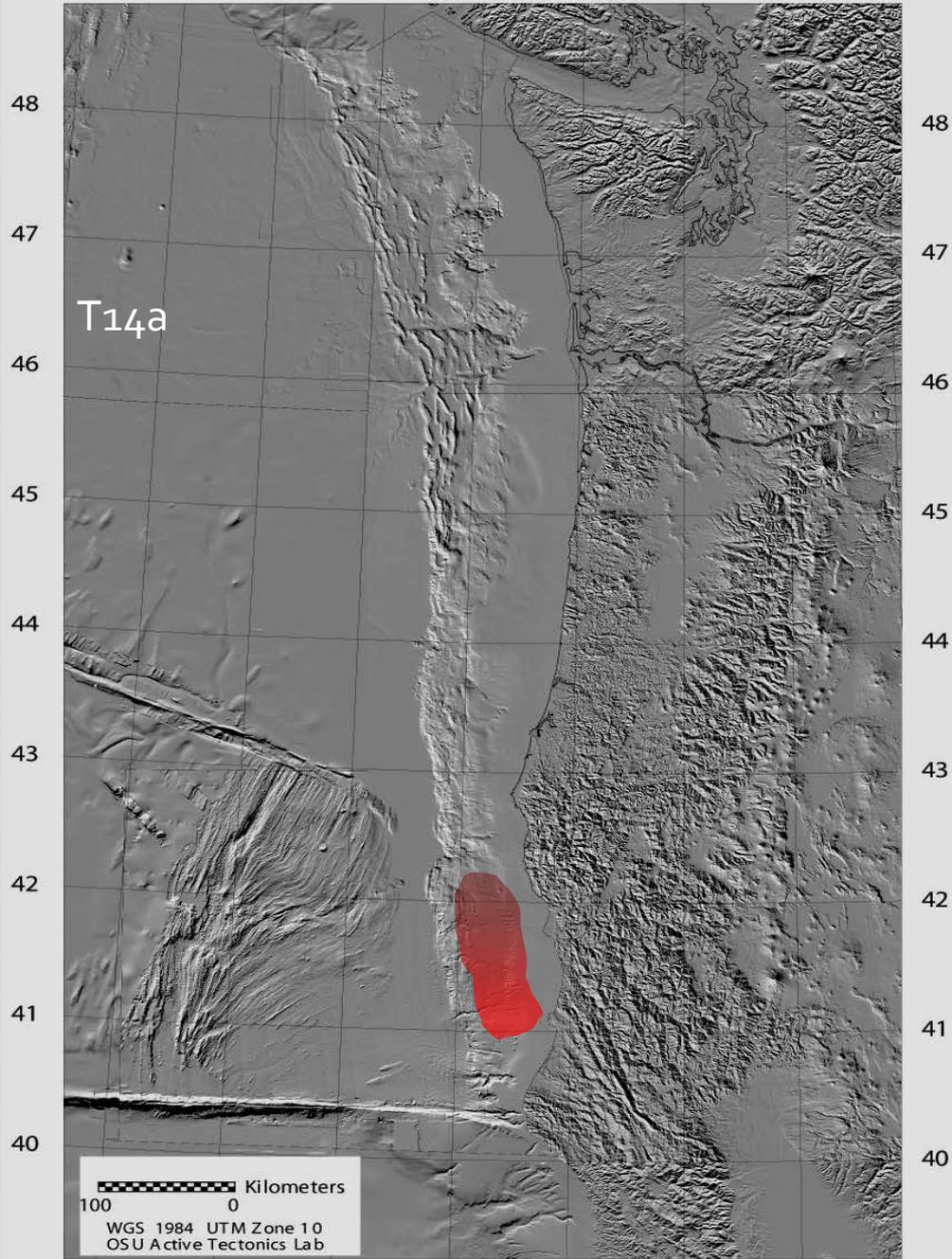
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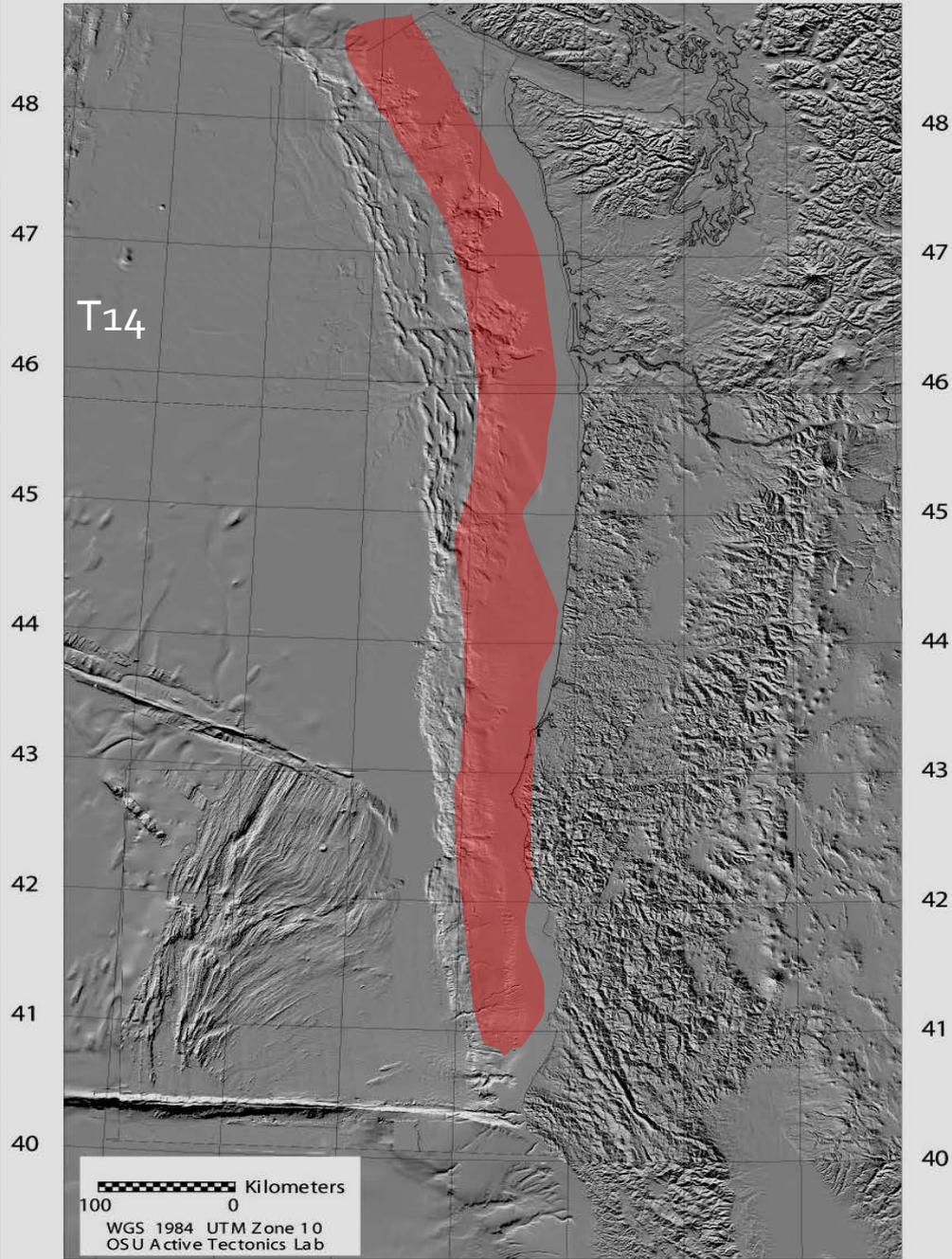
100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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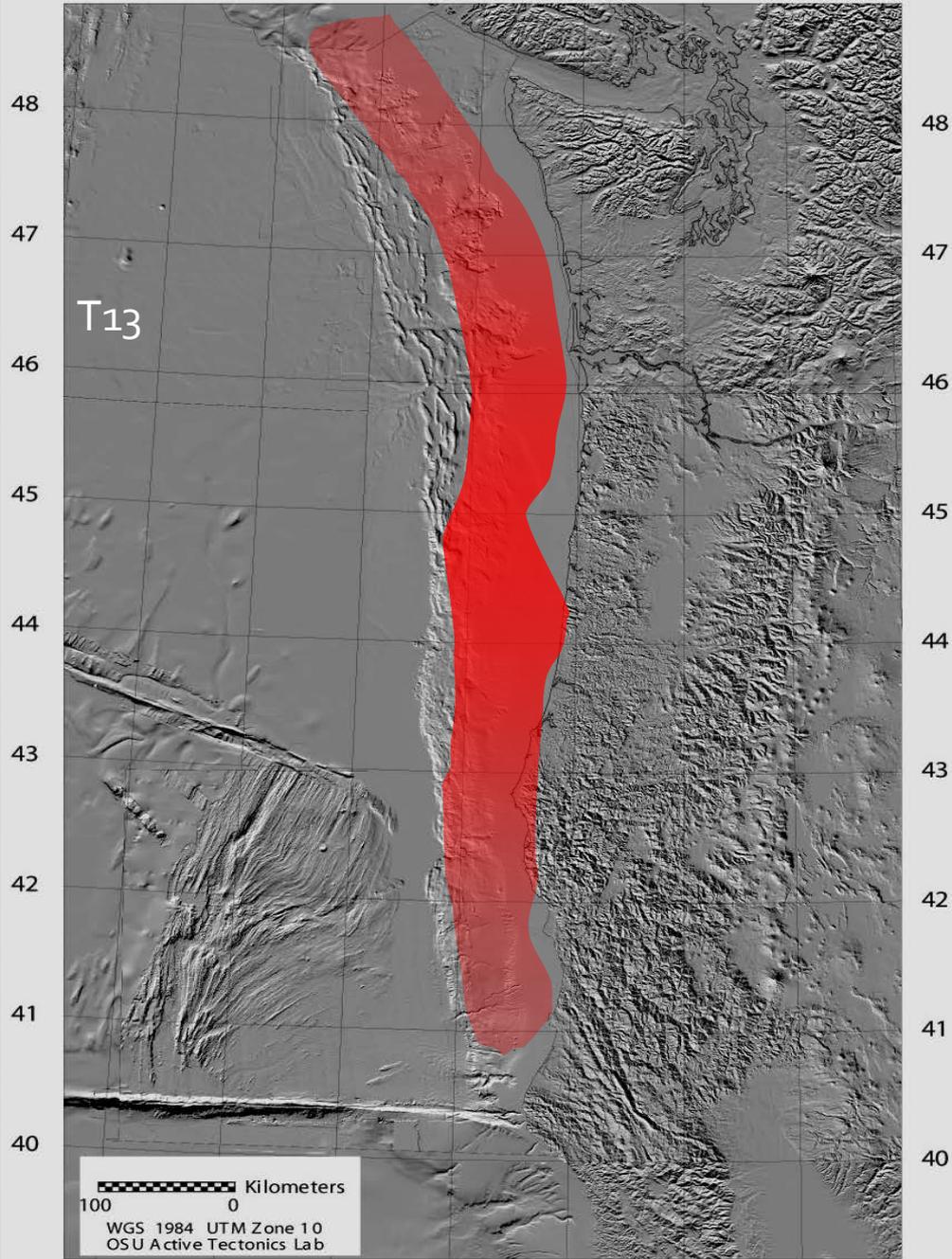


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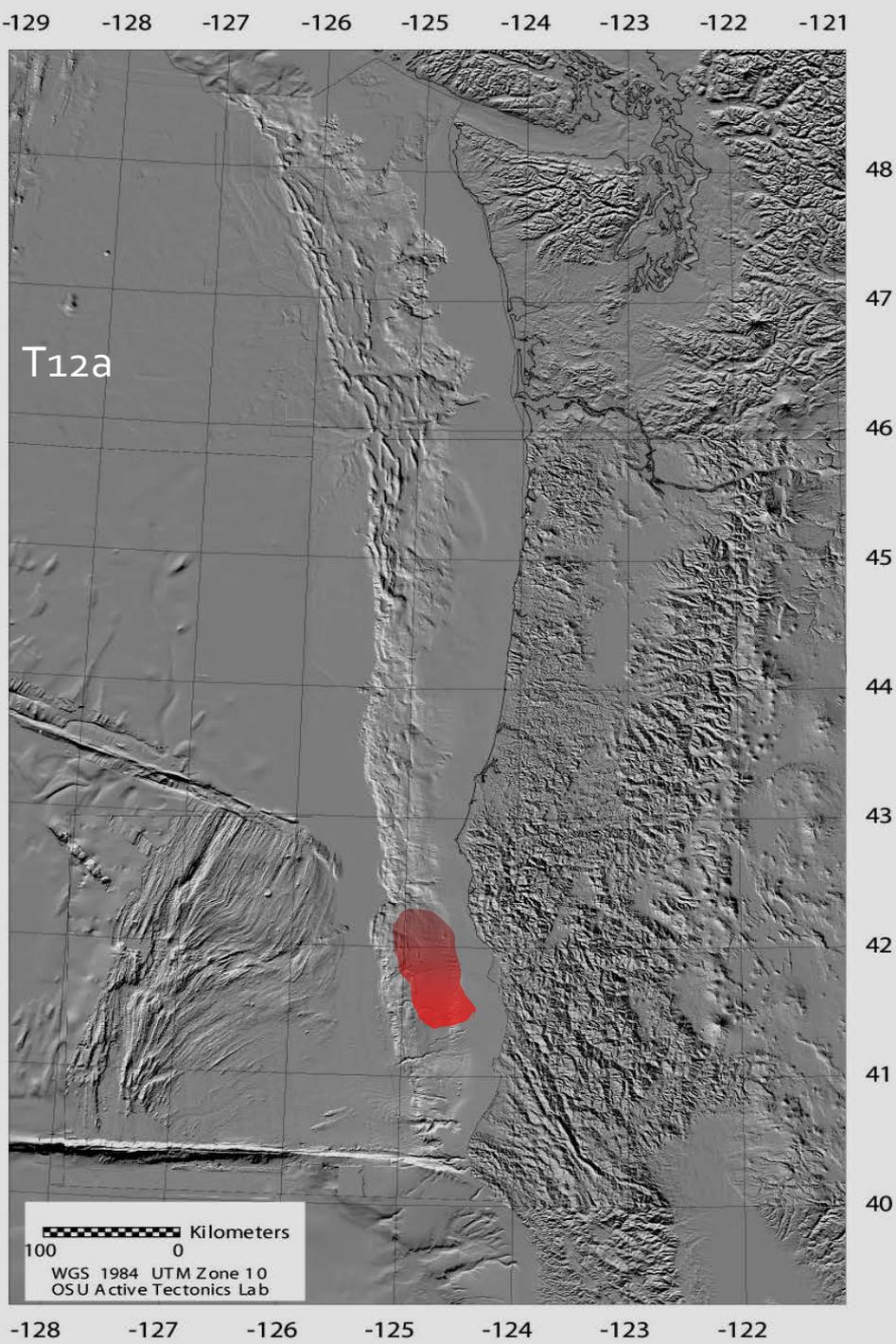
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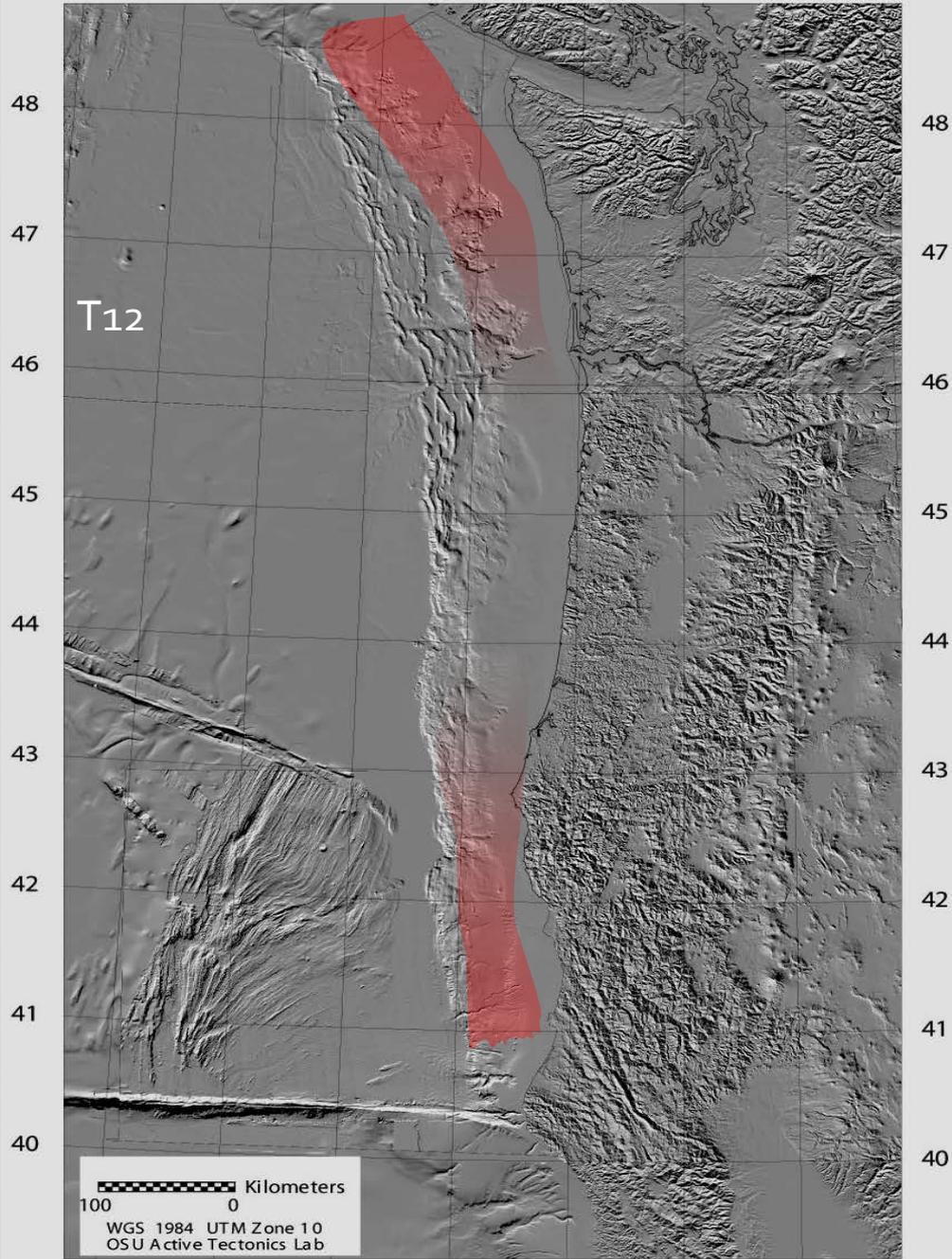
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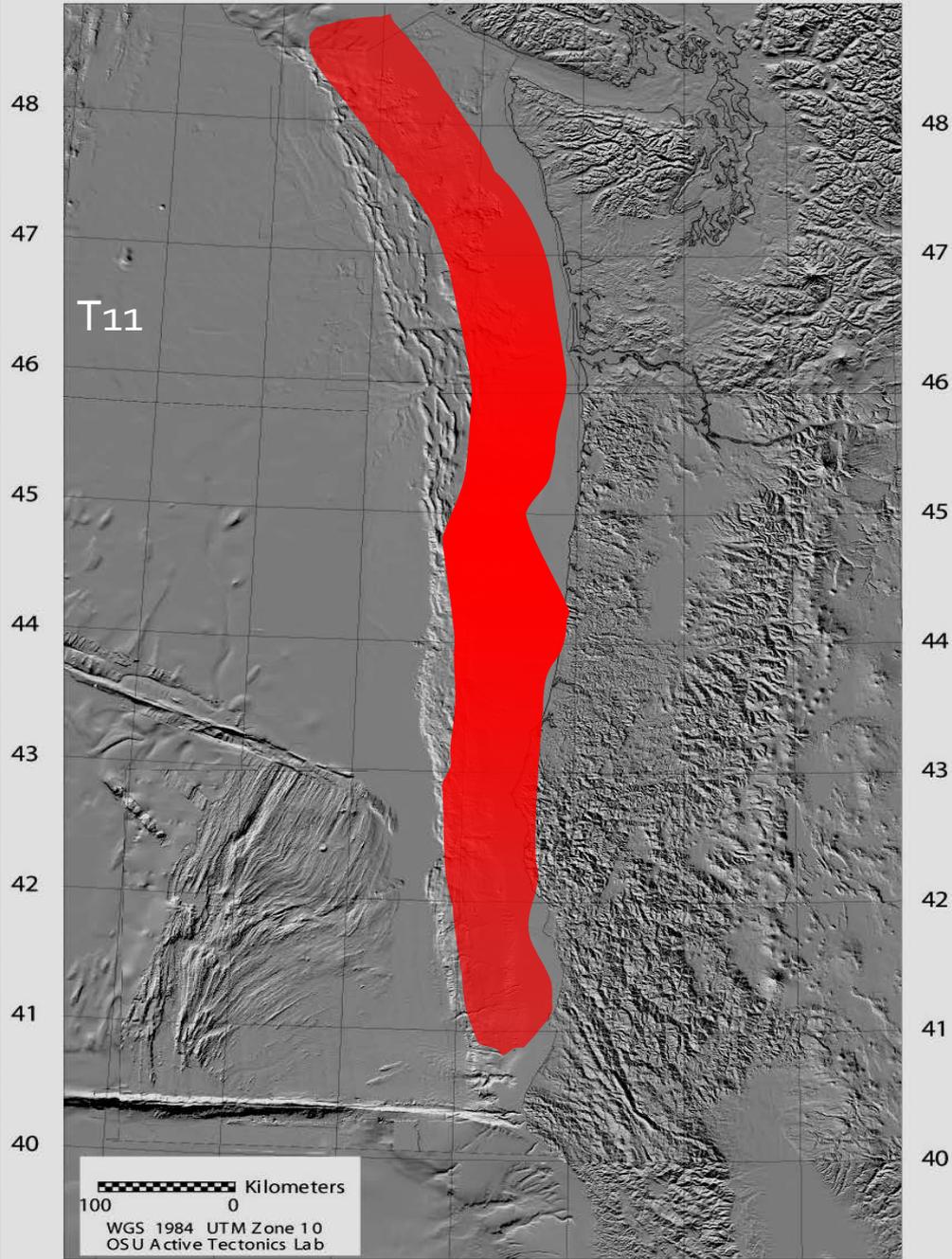
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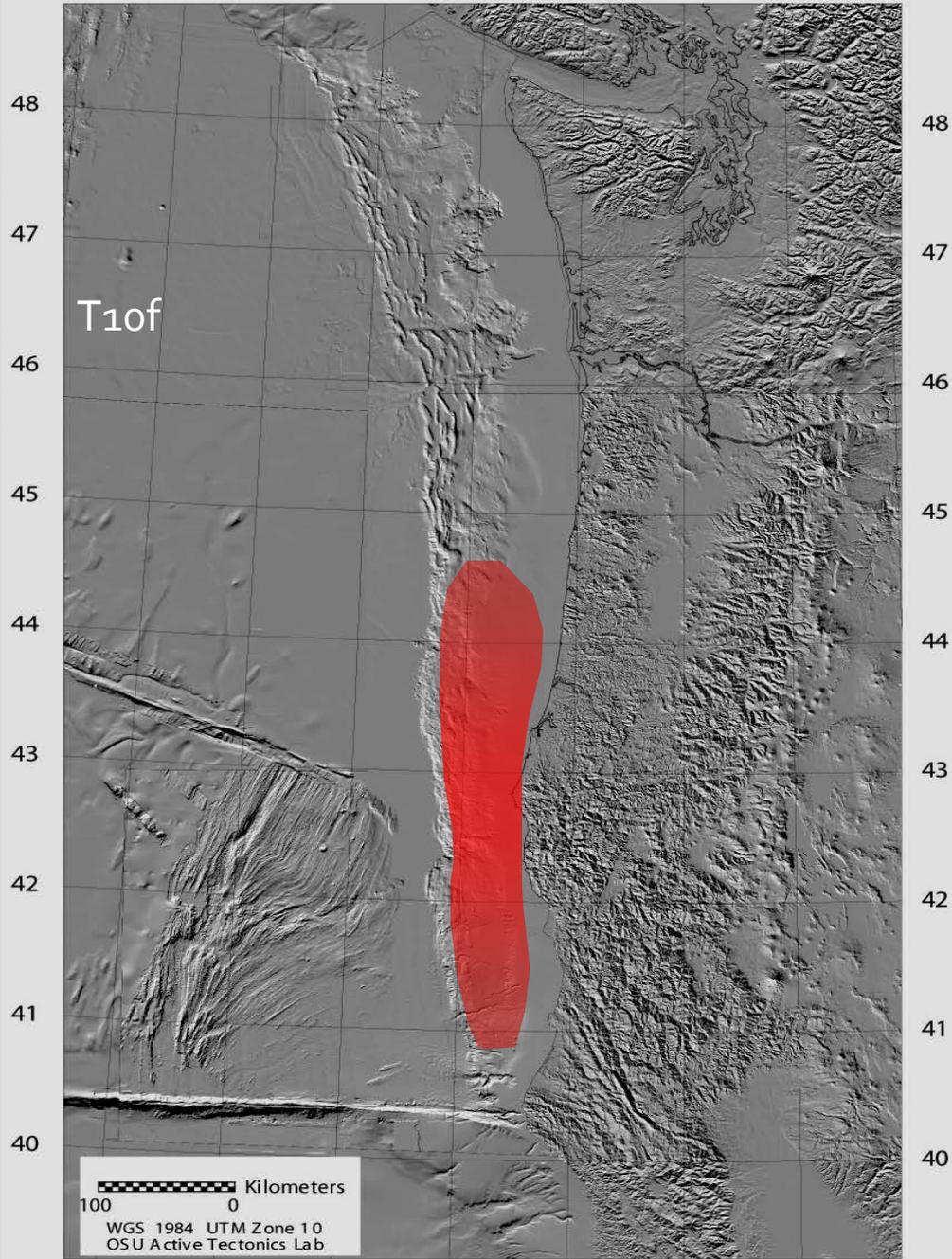
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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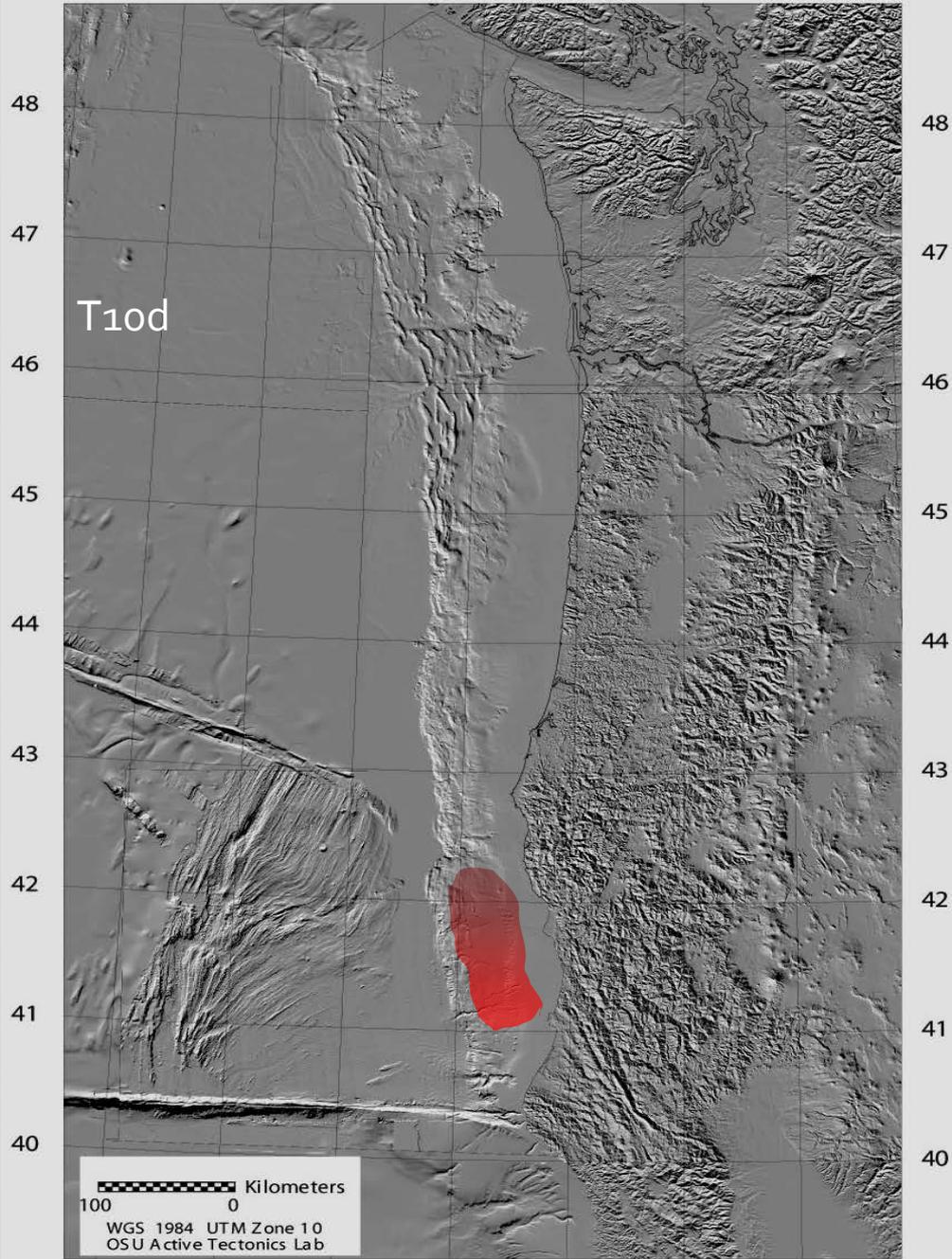


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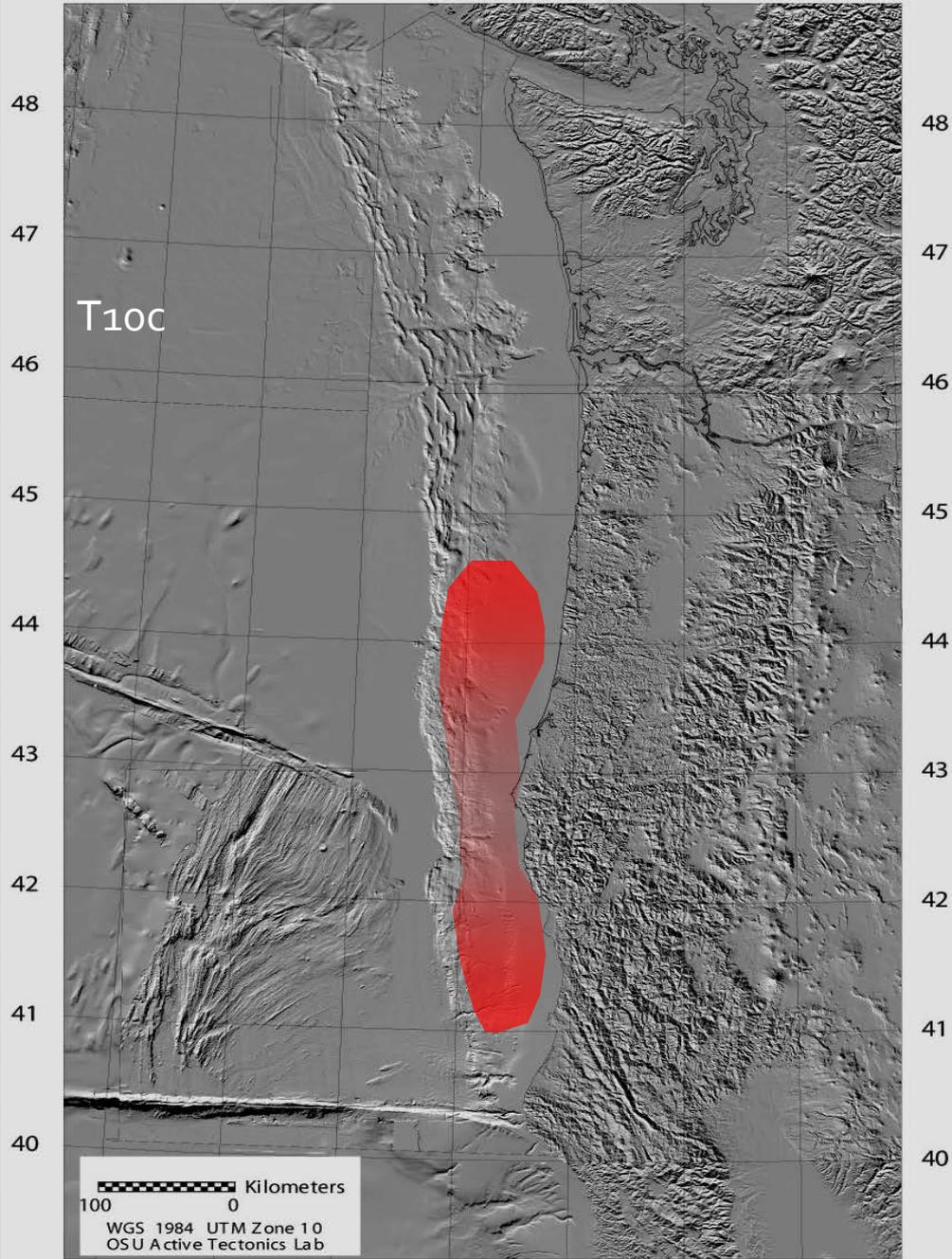
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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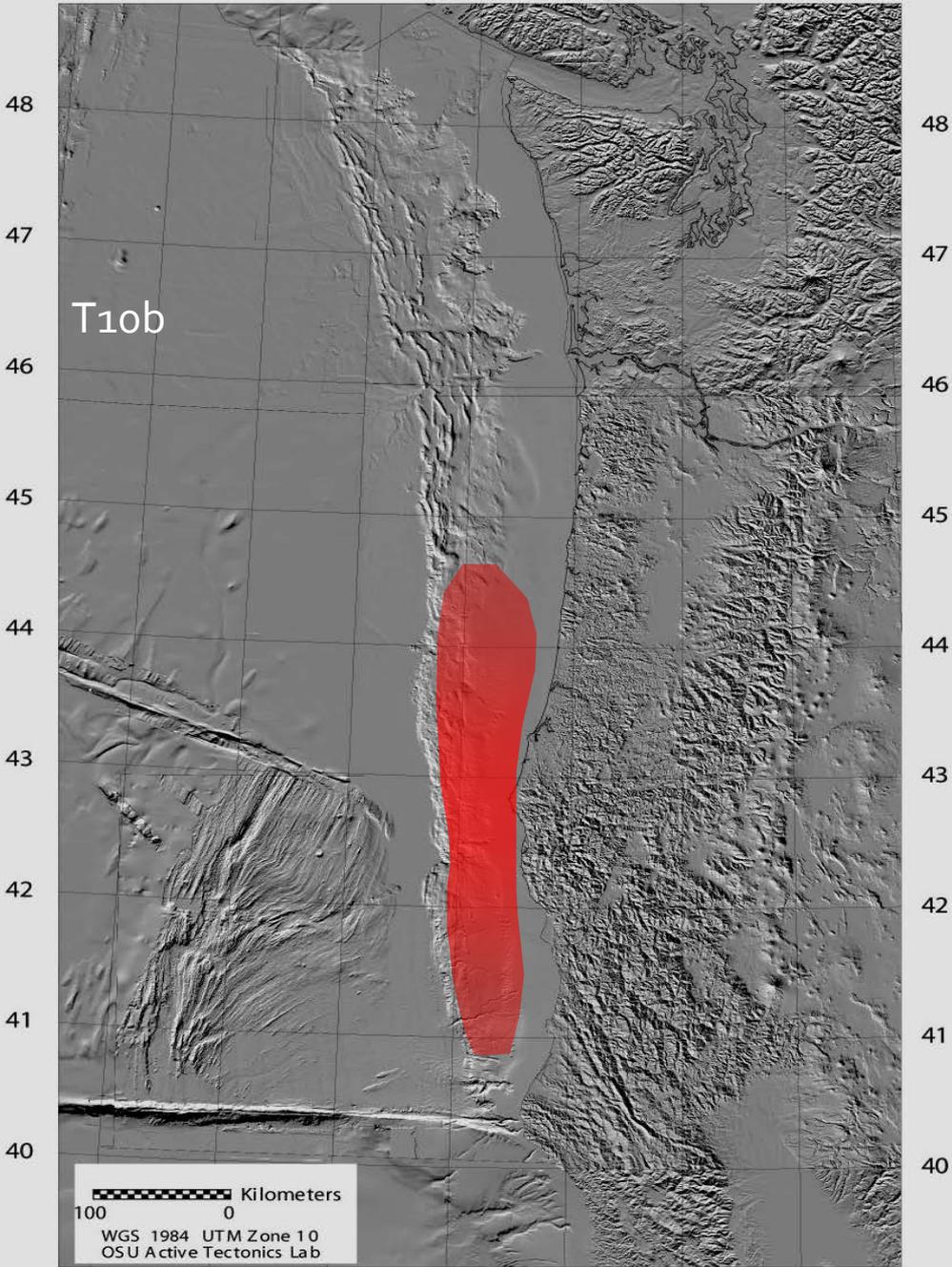
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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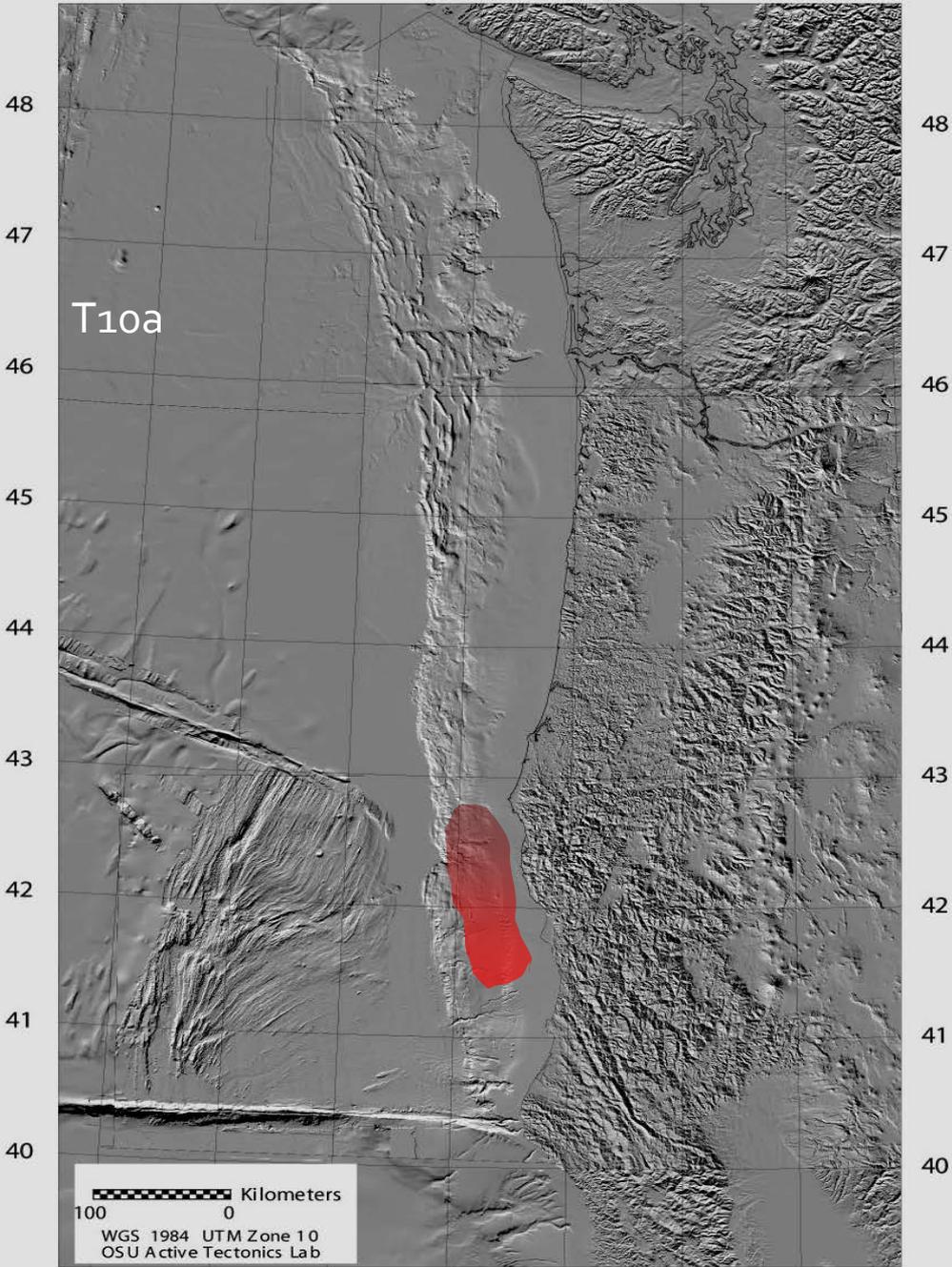


T10b

100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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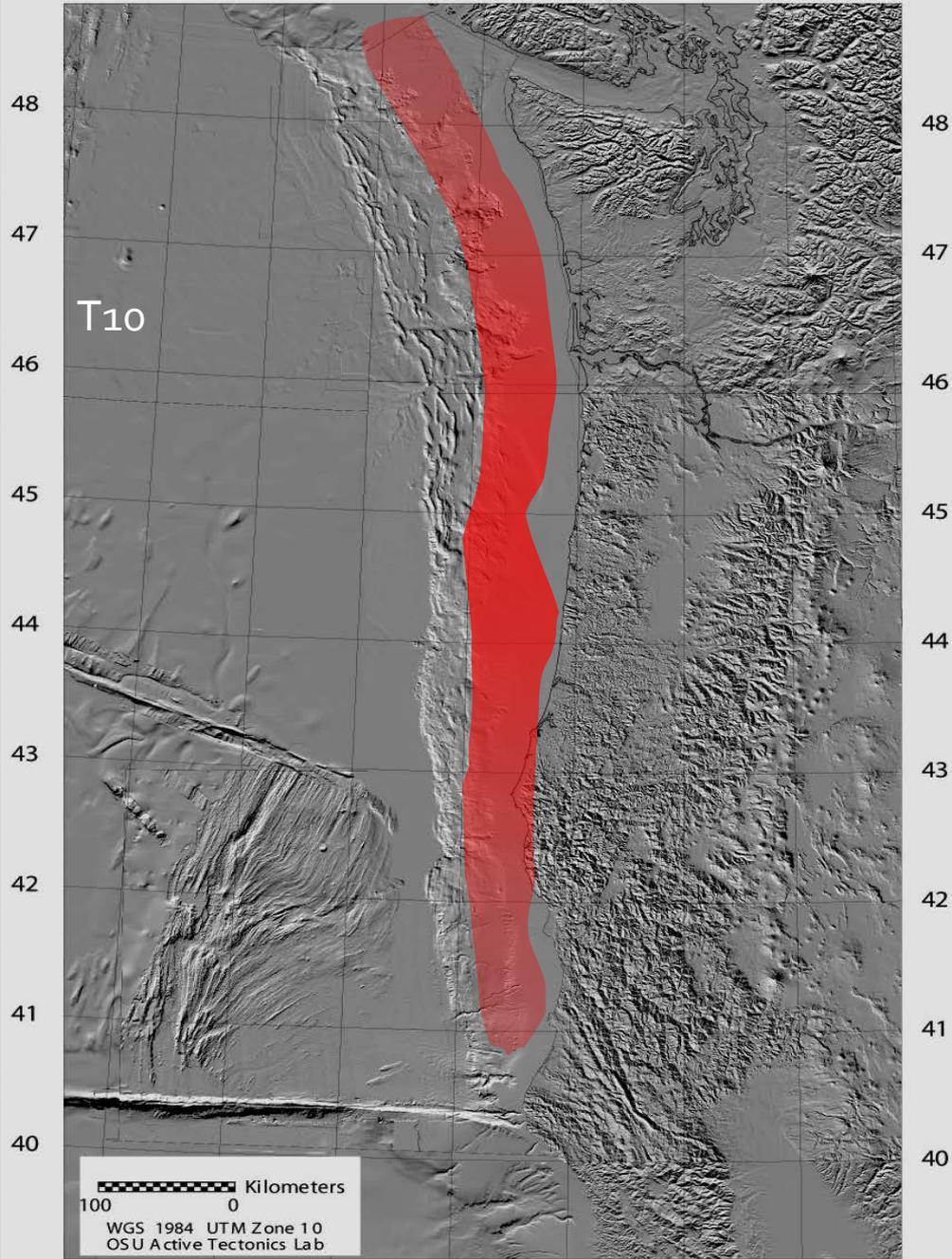
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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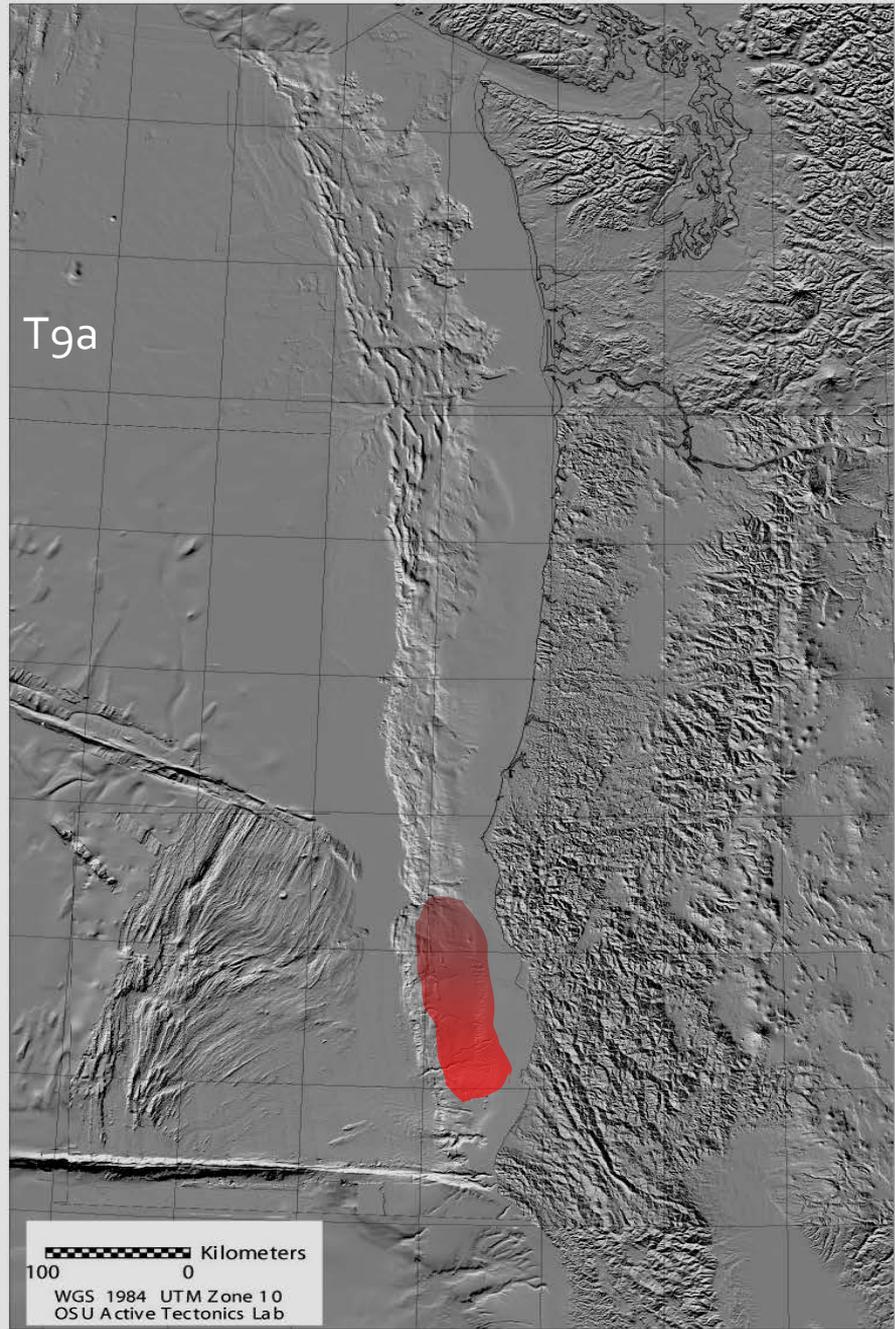
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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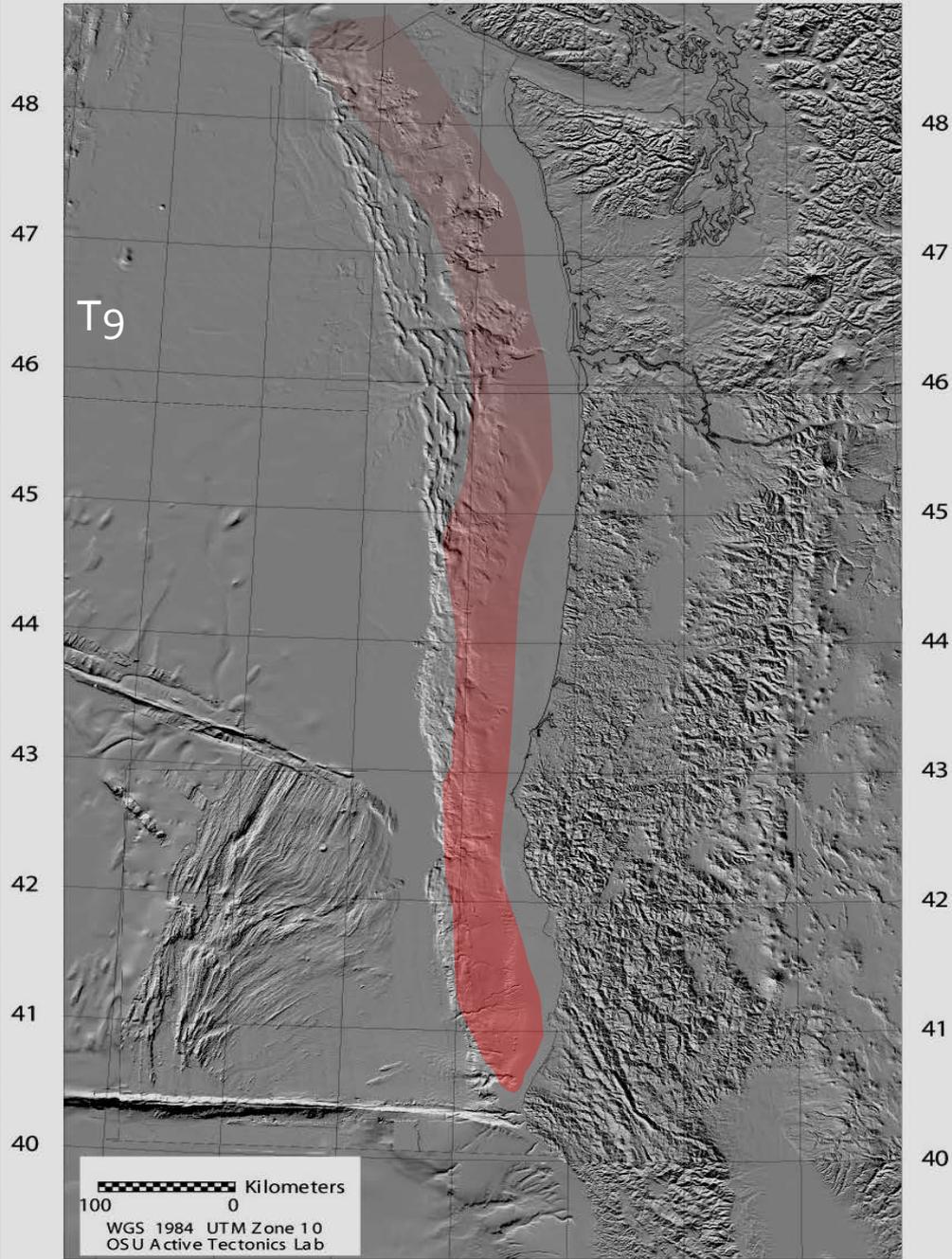
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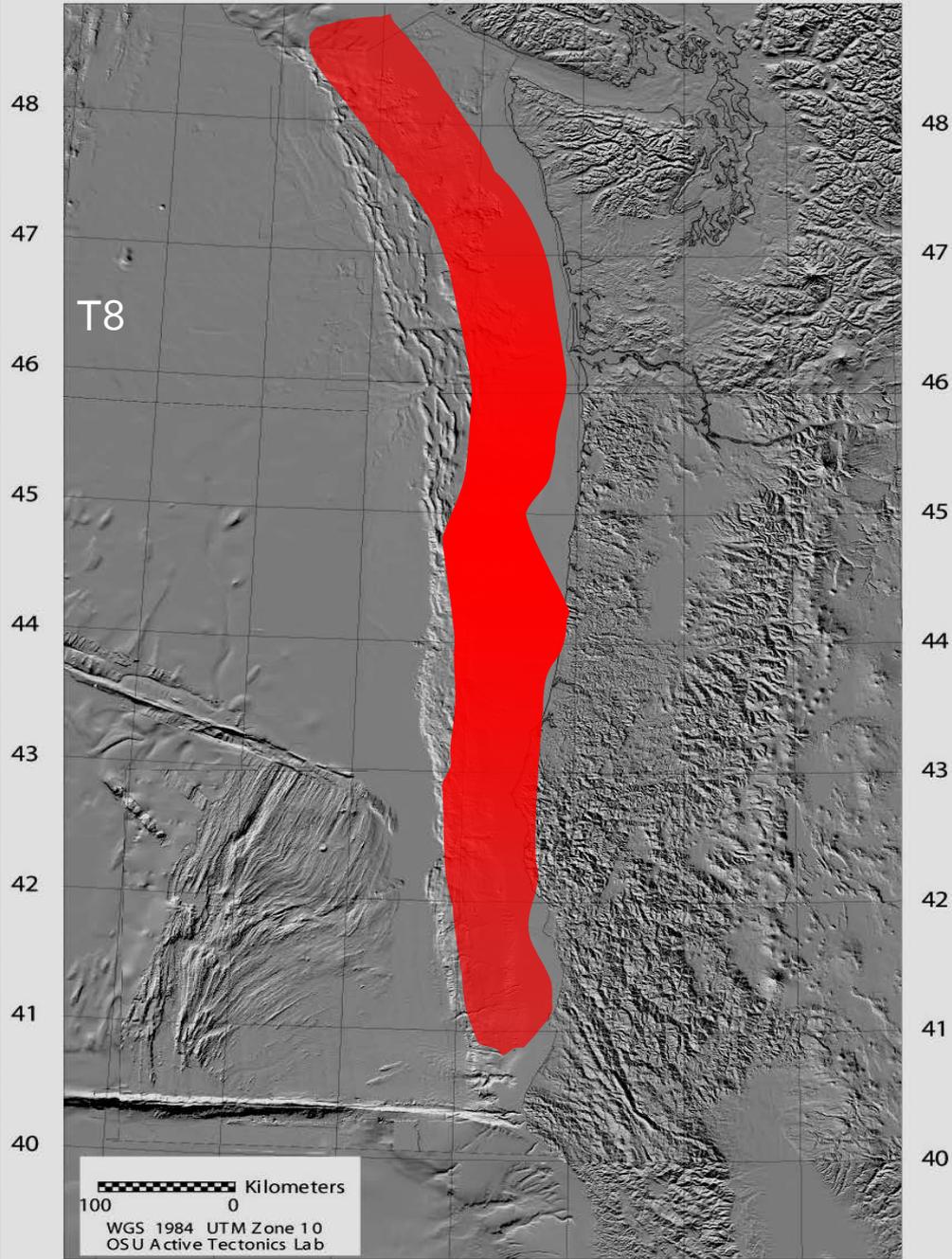


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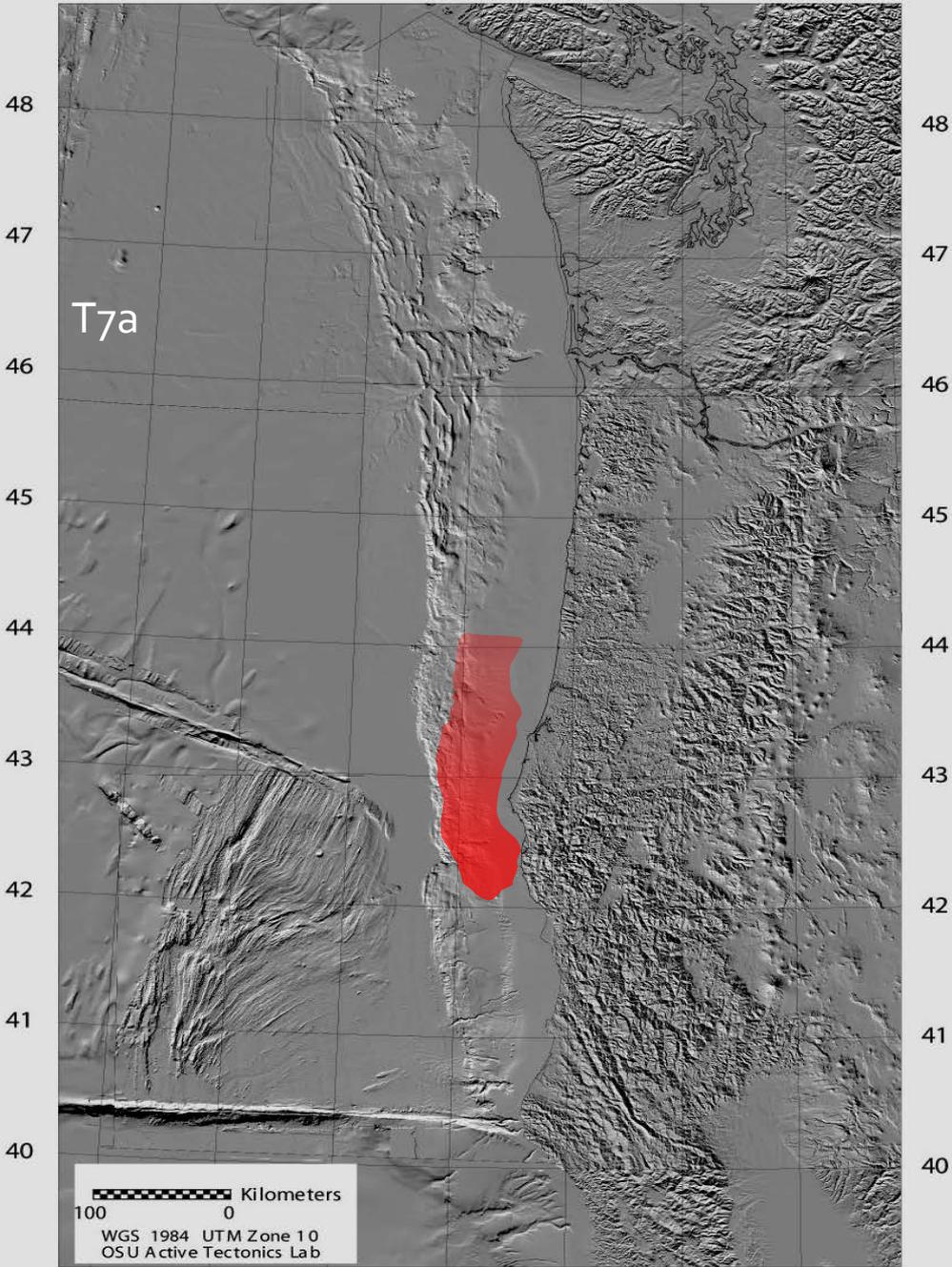
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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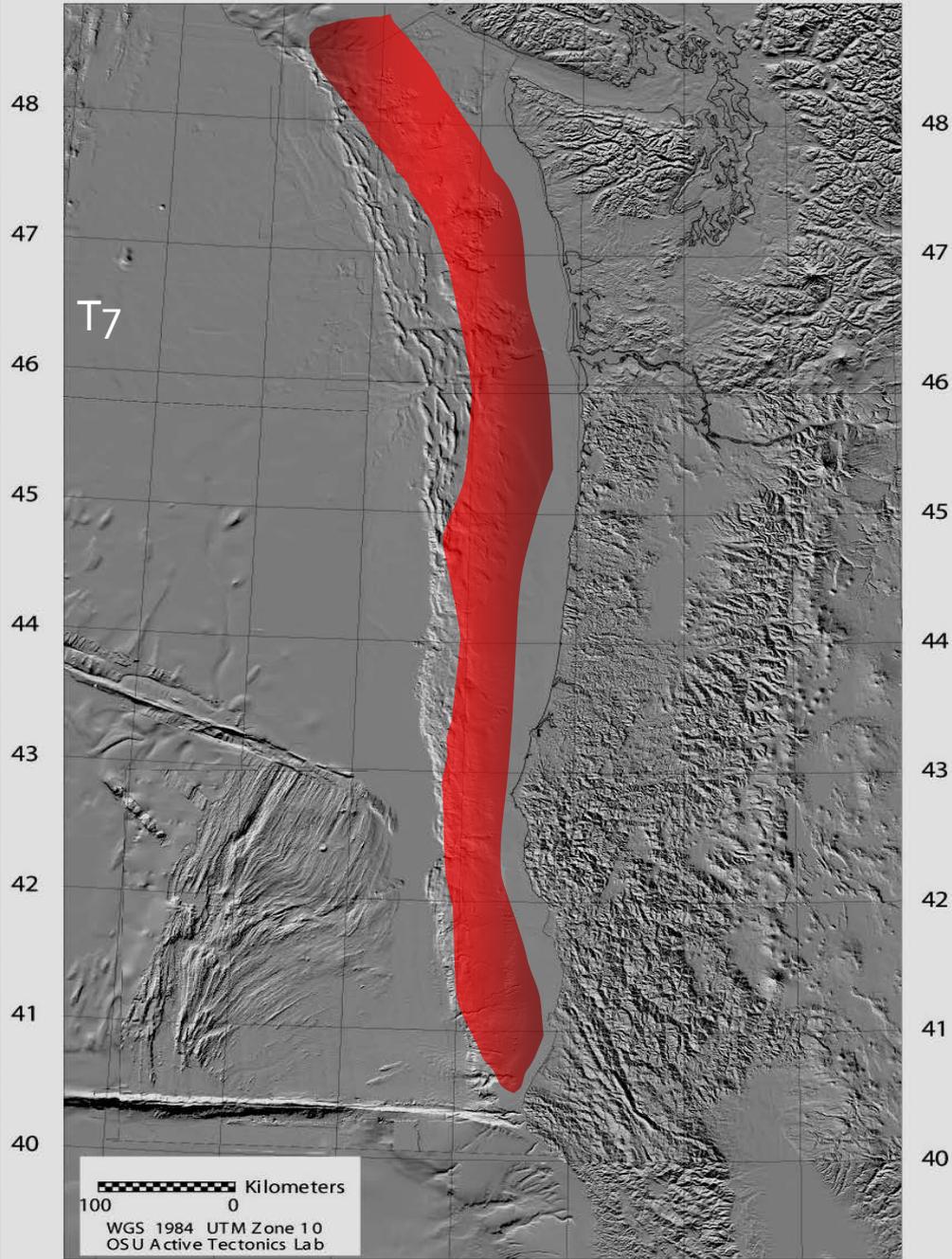
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

-128 -127 -126 -125 -124 -123 -122

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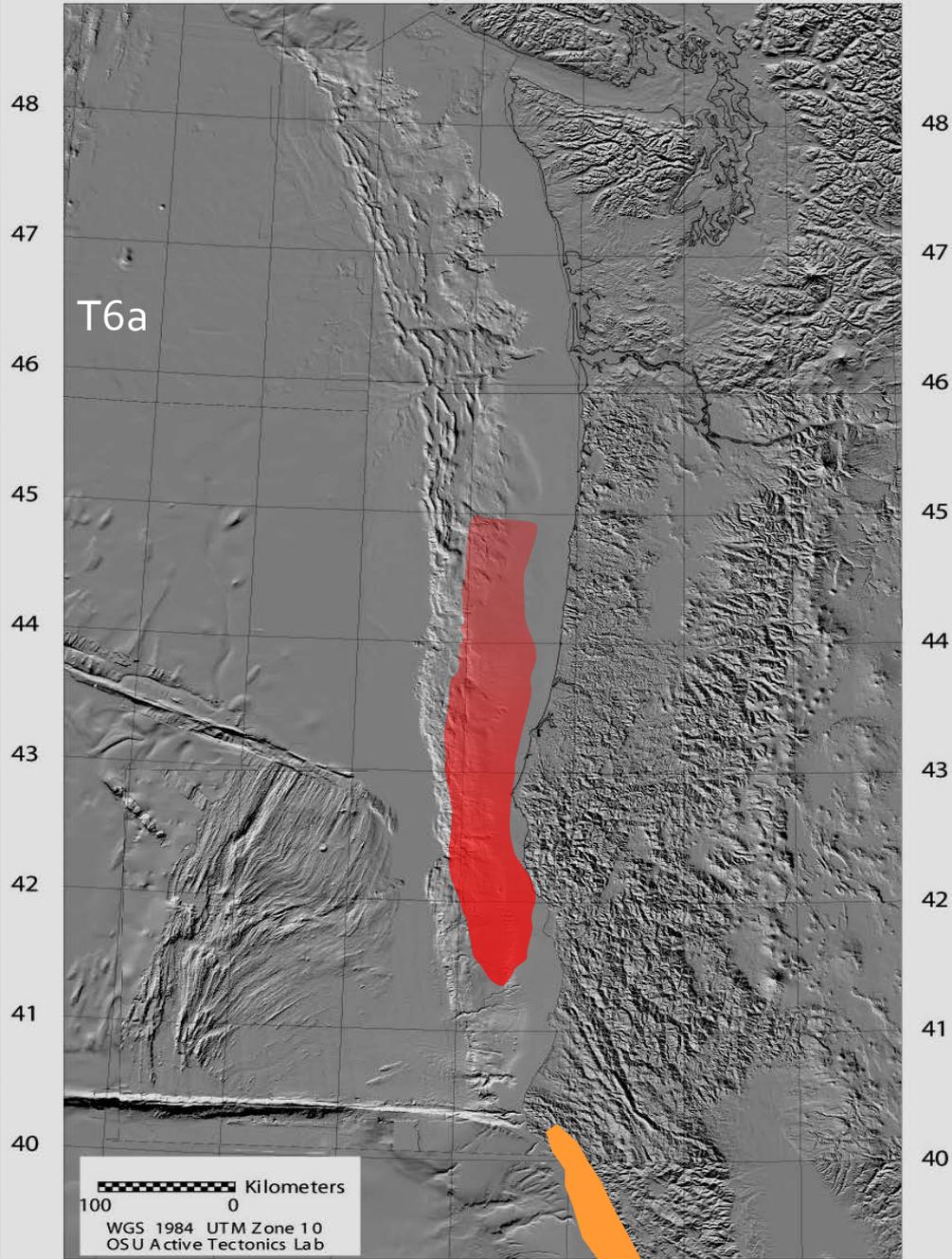
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100 0 Kilometers
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OSU Active Tectonics Lab

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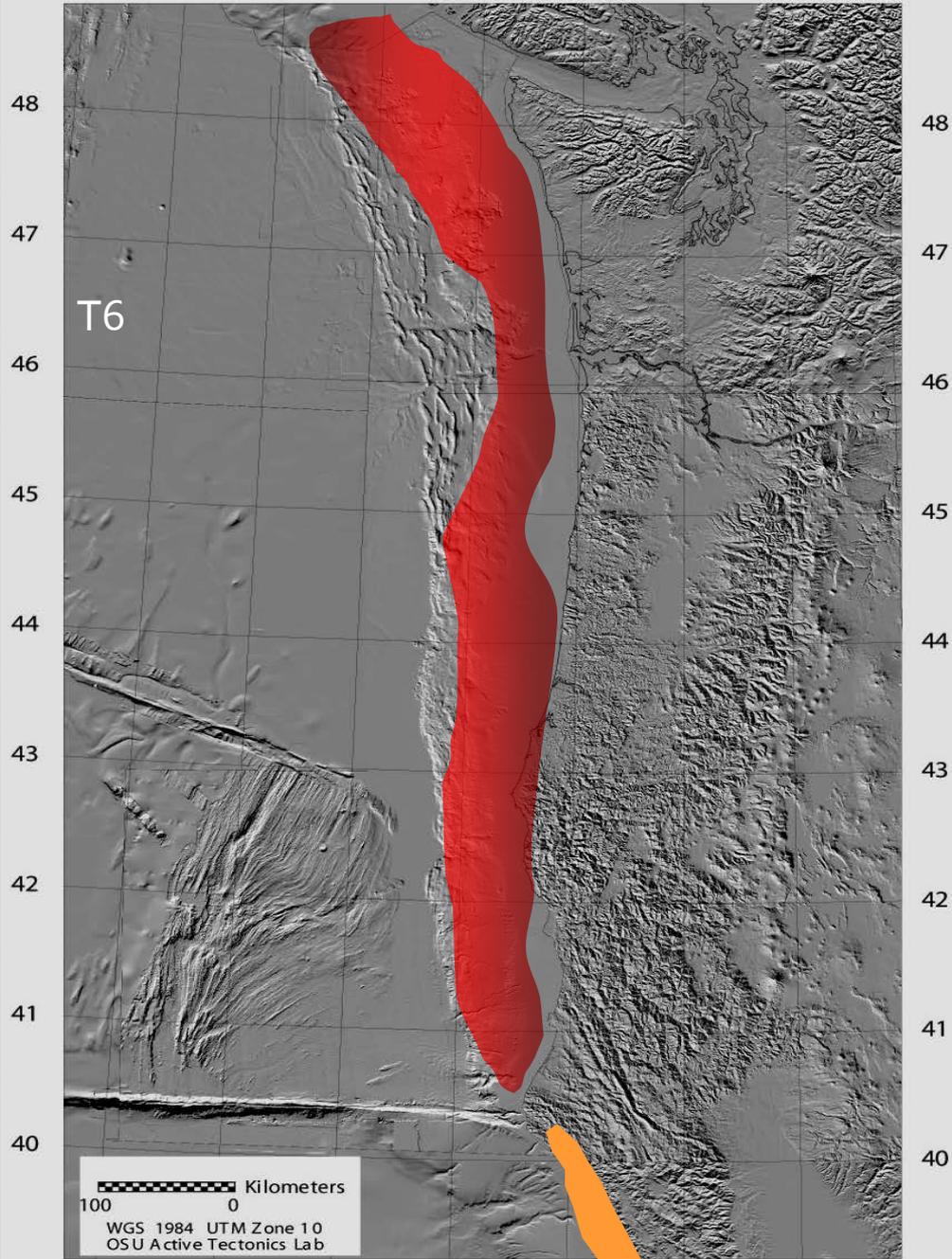


T6a

100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

-128 -127 -126 -125 -124 -123 -122

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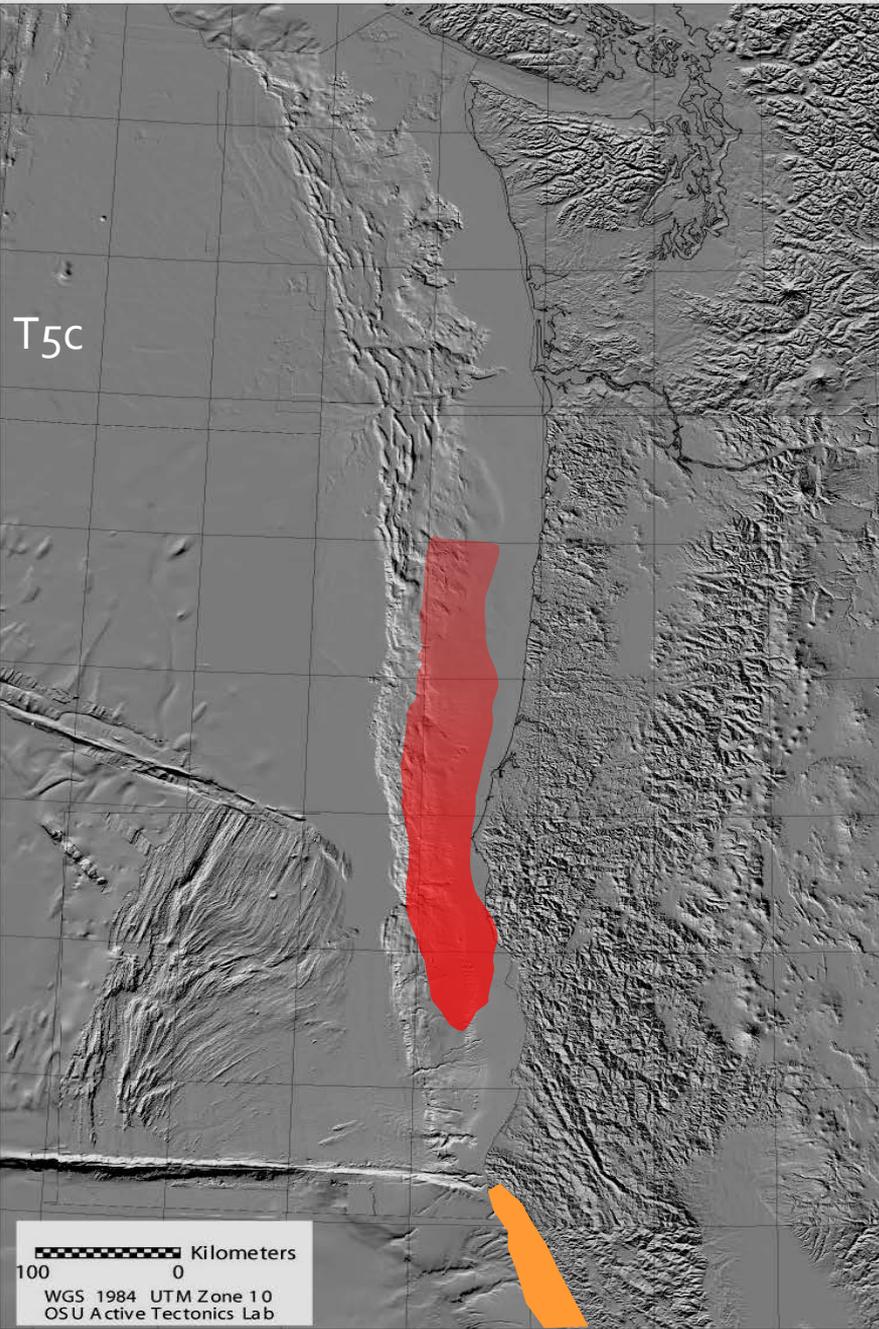
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T5c

100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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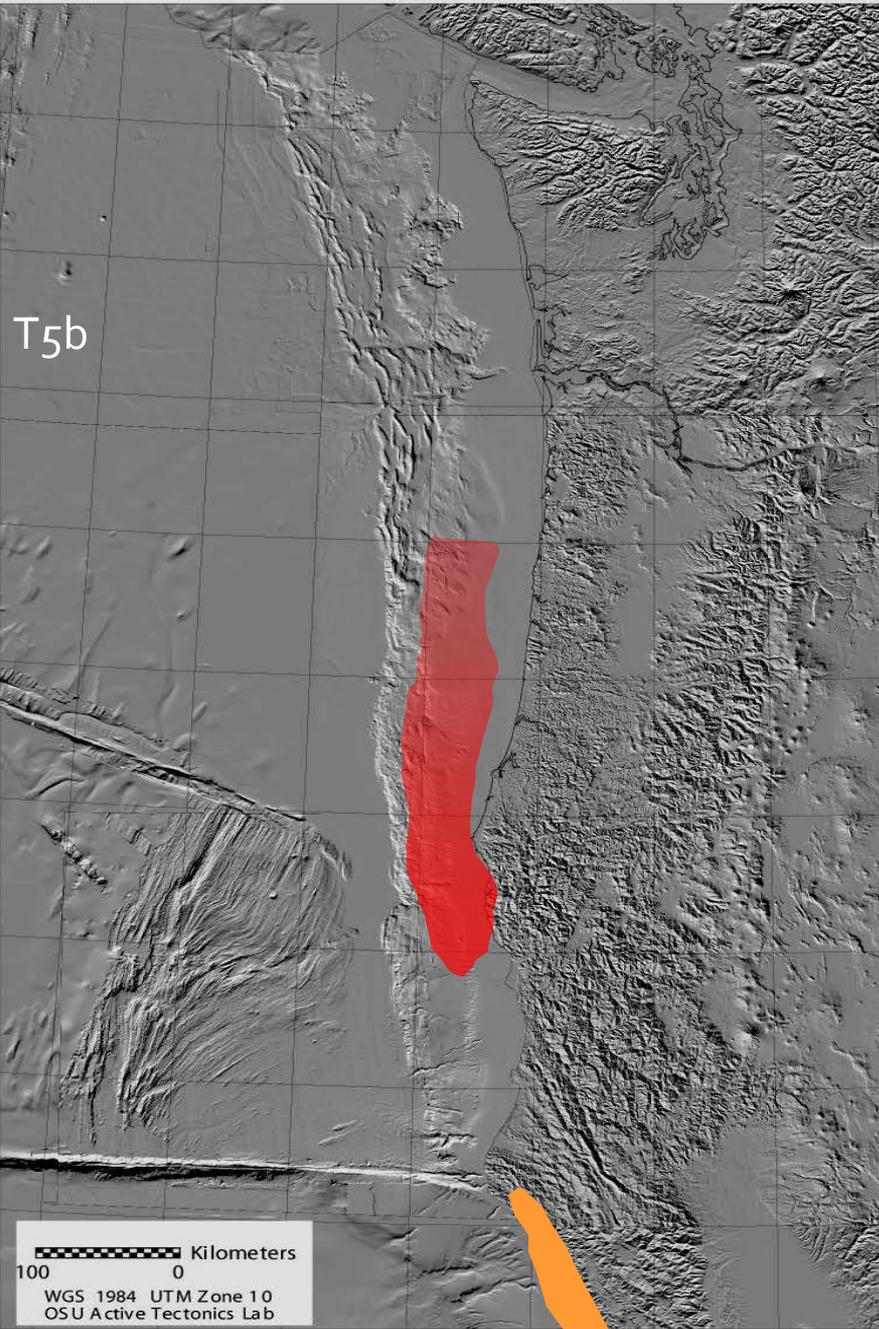
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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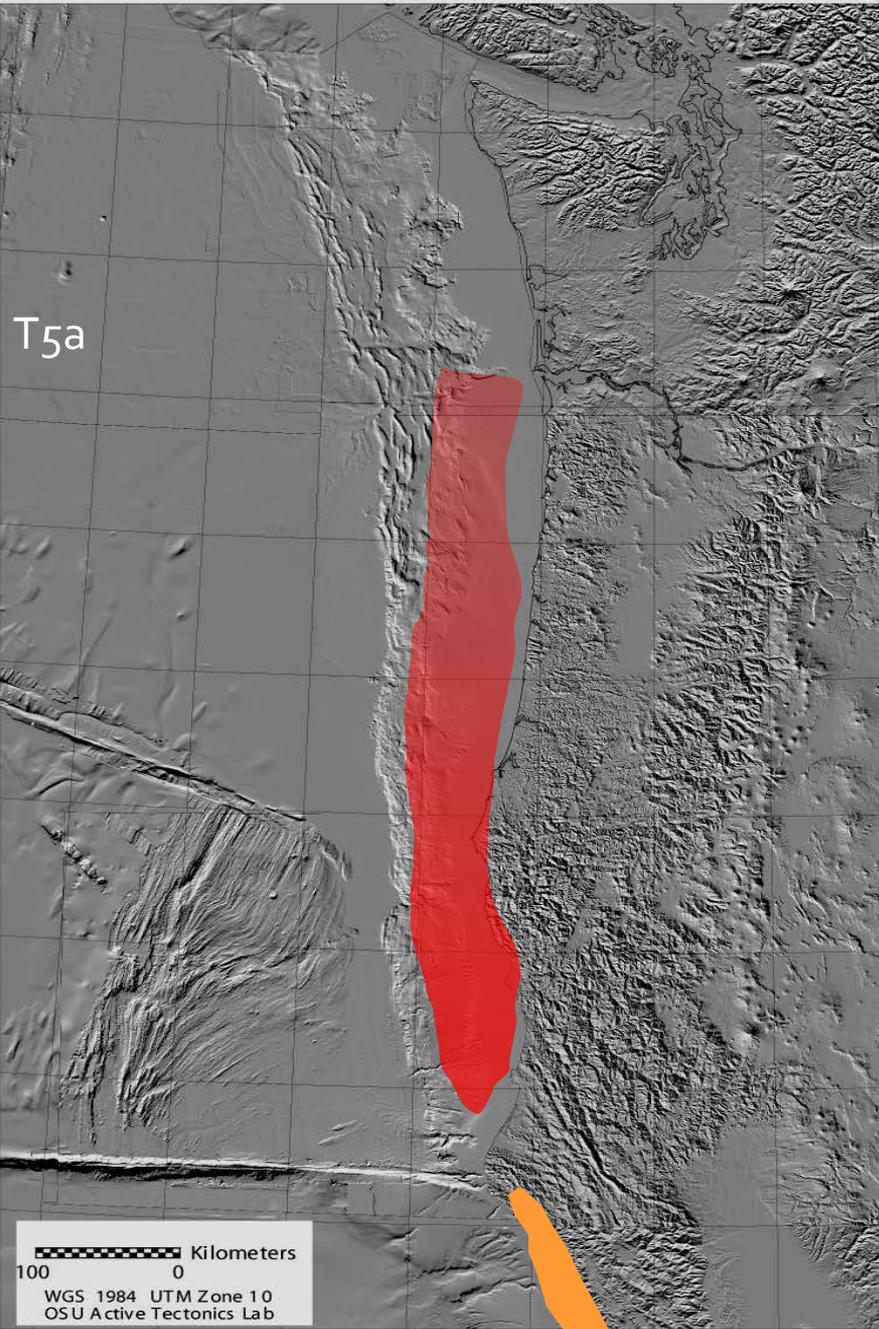
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T5a

100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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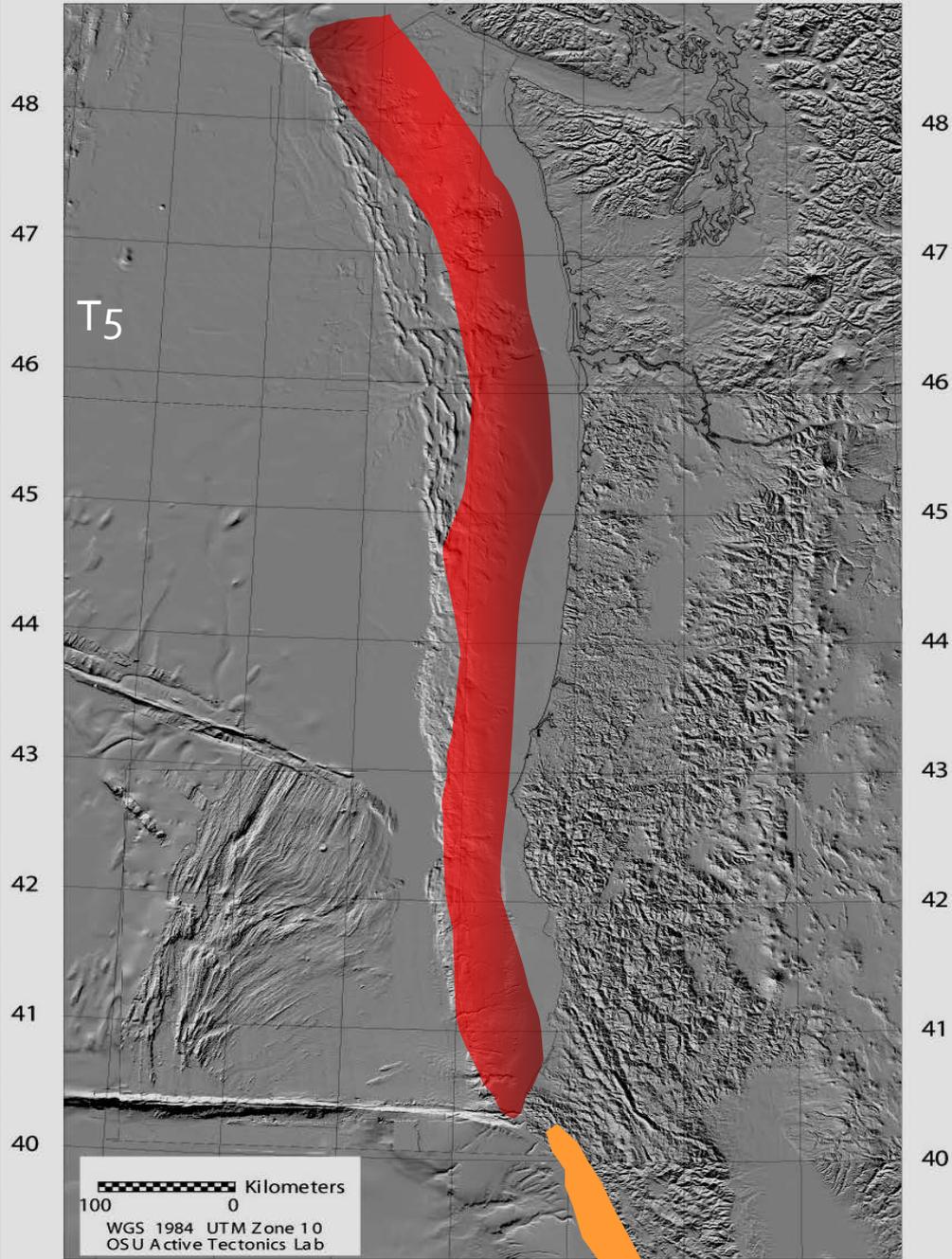
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100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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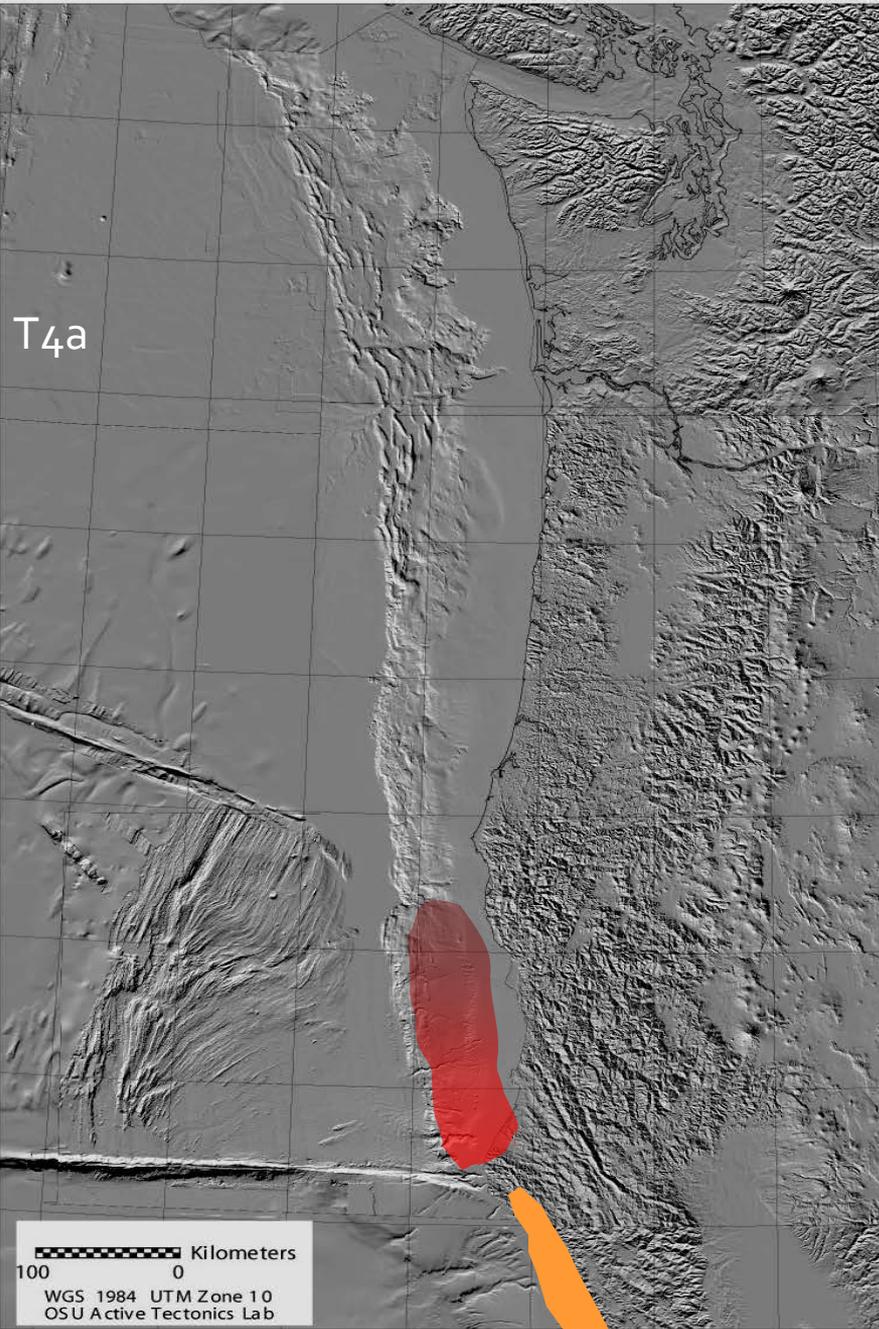
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T_{4a}

100 0 Kilometers
WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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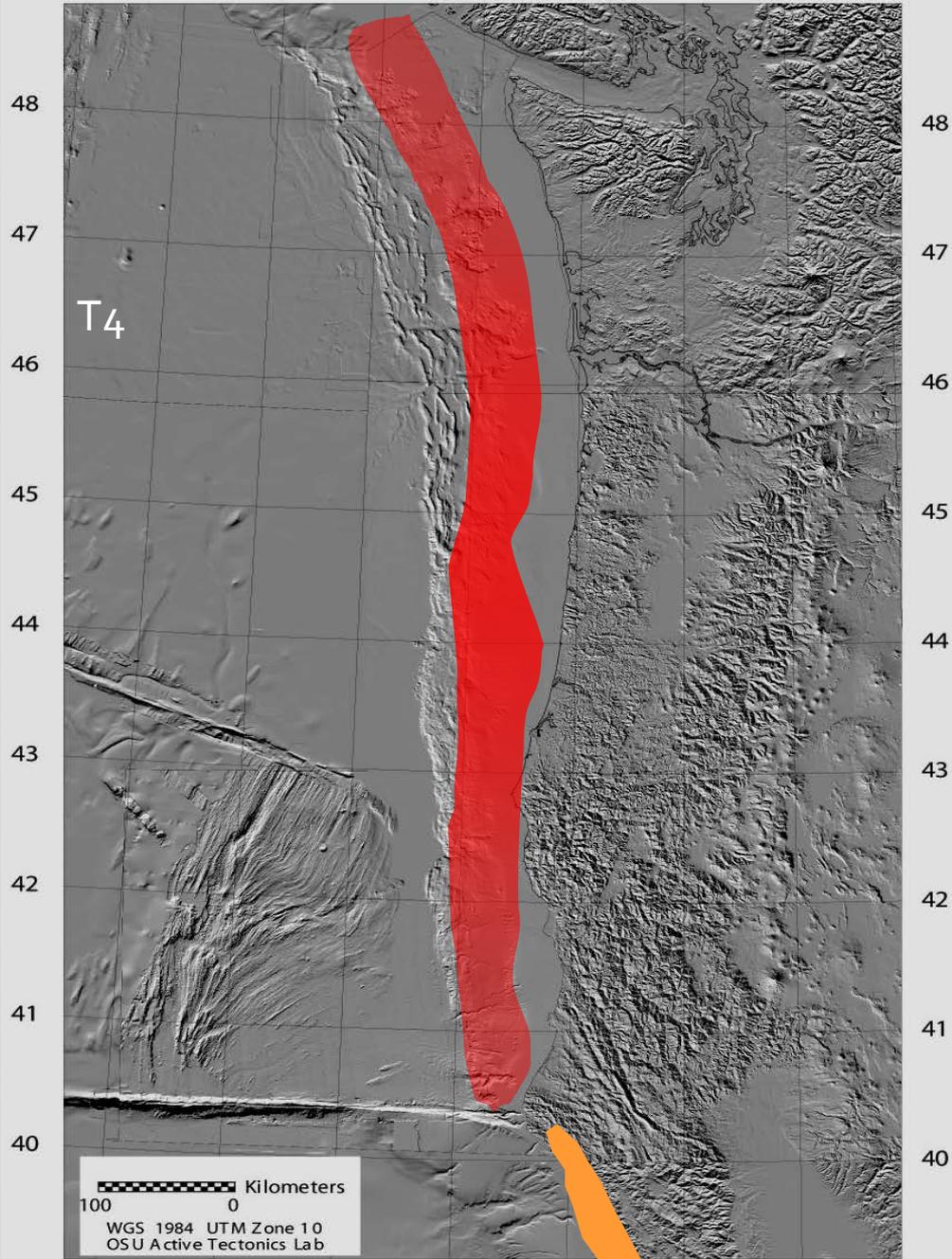
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Kilometers
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WGS 1984 UTM Zone 10
OSU Active Tectonics Lab

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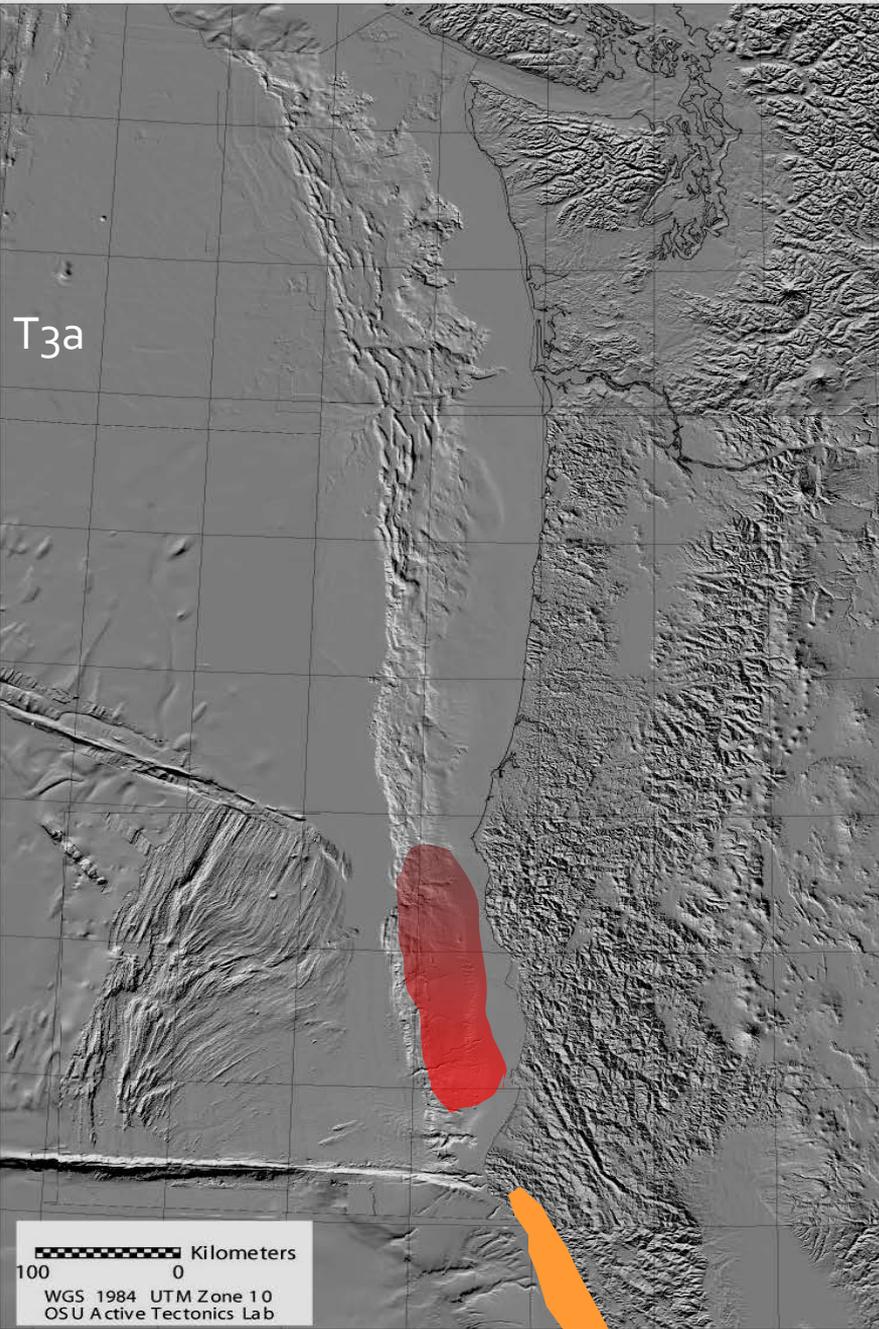
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OSU Active Tectonics Lab

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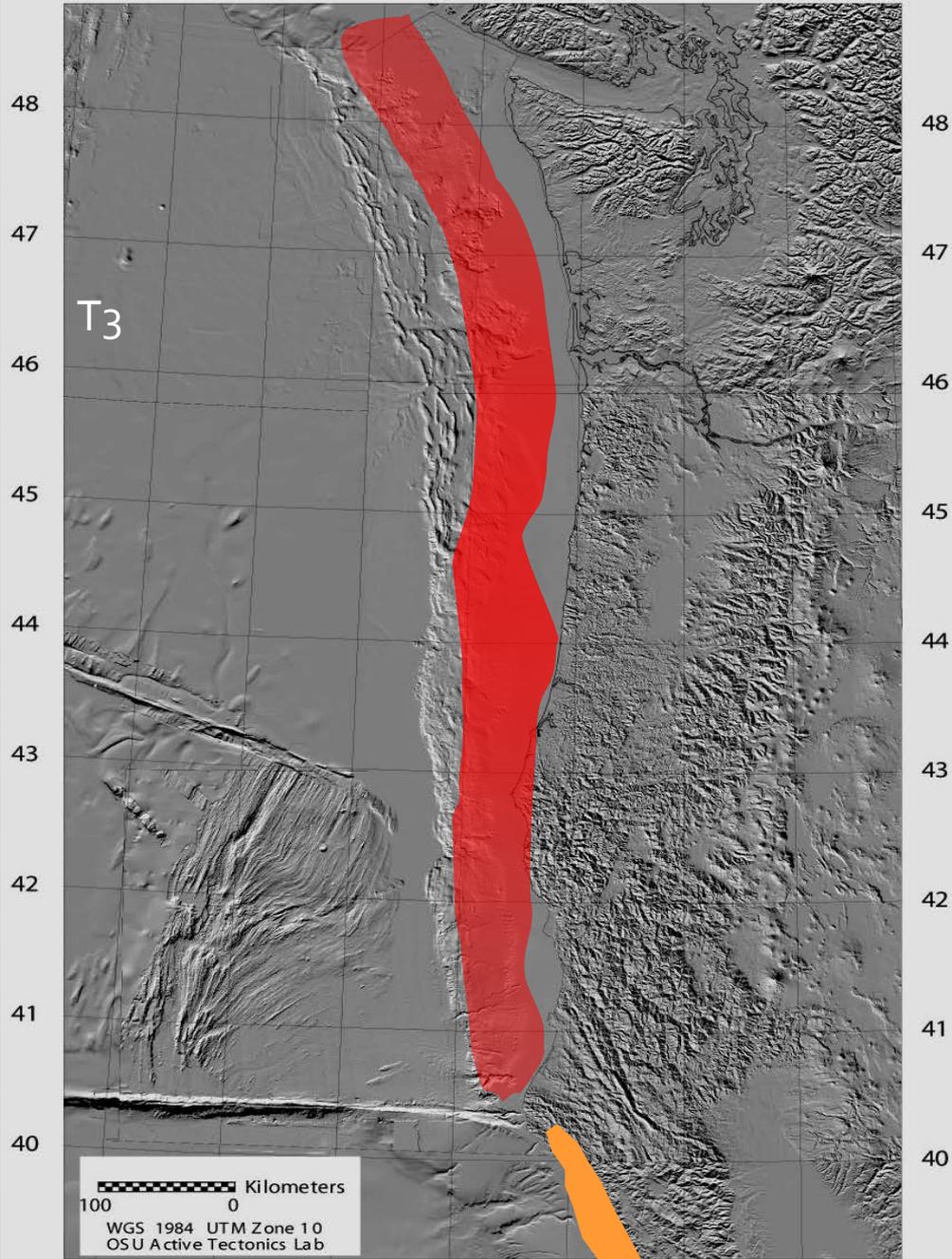
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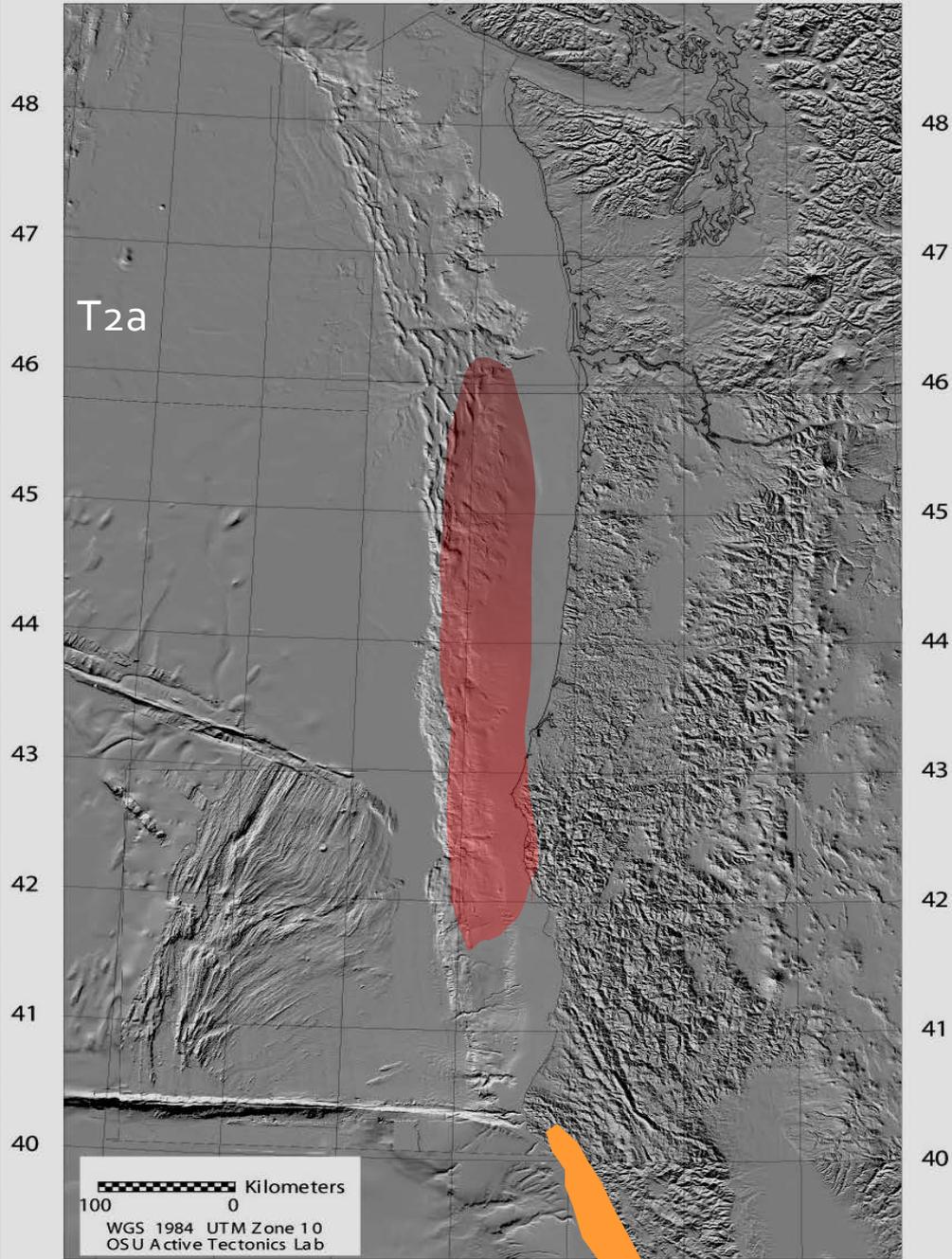
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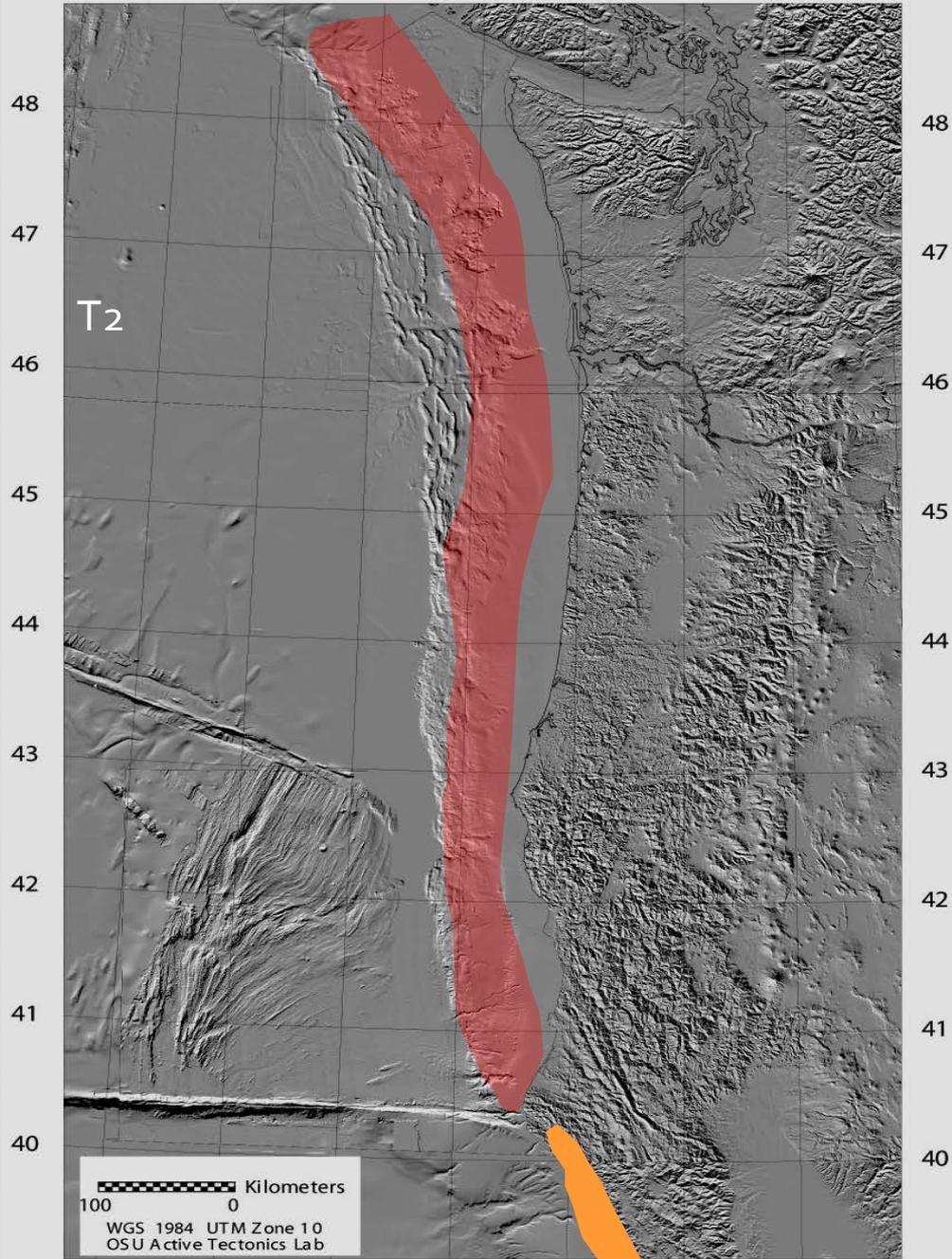
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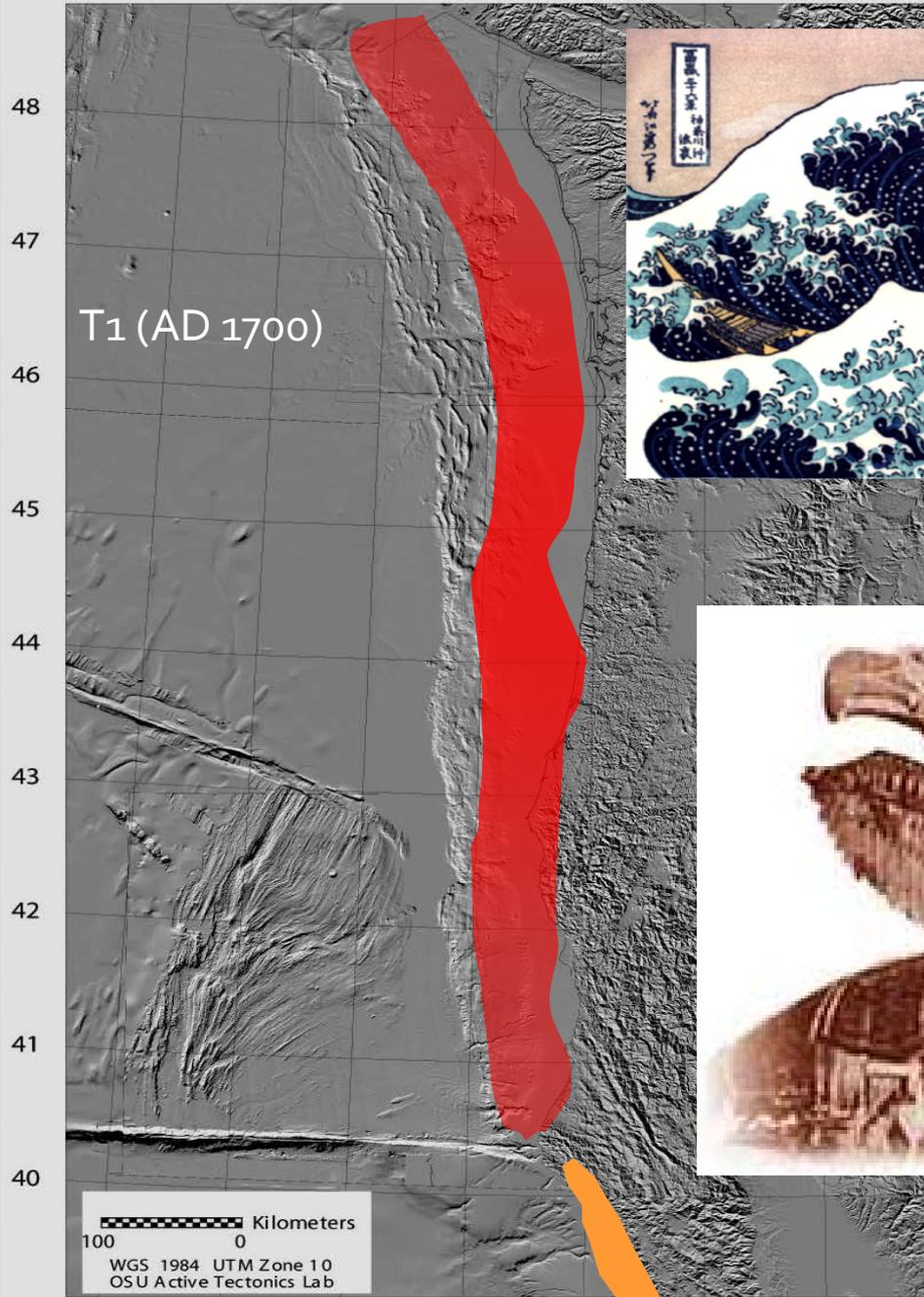
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OSU Active Tectonics Lab

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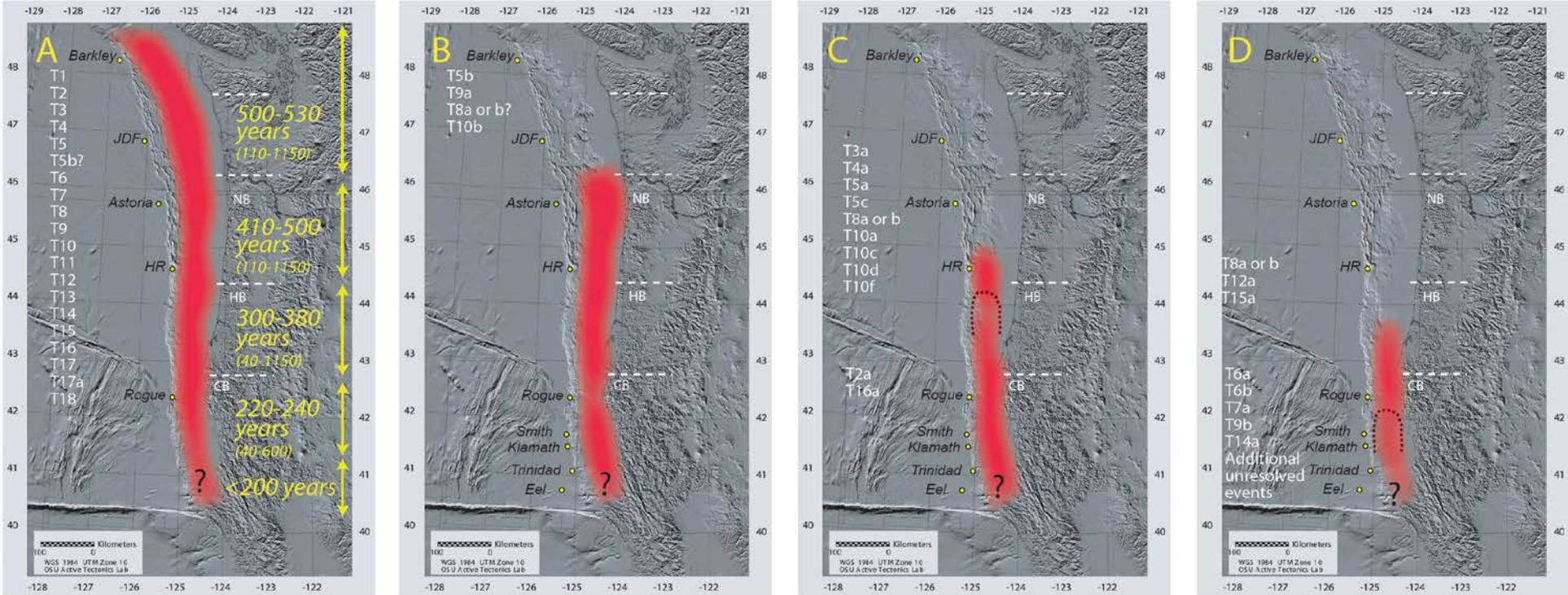


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Rupture lengths from Holocene Offshore/Onshore paleoseismic record.

While mean recurrence interval is ~ 500 years in northern Cascadia, it is only 240 years in the south. The NSAF recurrence during this time is similar, ~200 years. Smaller Cascadia events correlate with shorter rupture length and Bradley Lake tsunami deposits.

Segment boundaries are roughly compatible with ETS segment boundaries proposed by Brudzinski et al., 2007, though both sets of boundaries are quite crude.

Conclusions

- Understanding of channel pathways and their history and dynamics is critically important. Erroneous conclusions may be drawn from inappropriate sampling strategies.
- Many cores are required to unravel the unknowns. We may not be smarter than mud on the first try!
- Alternative pathways and turbidite counts are not supported by data or physics.
- Using appropriate cores from active channel/canyon systems, the turbidite stratigraphy is remarkably consistent. Linkages to Effingham Inlet, the Slipstream Slide, onshore lakes and marsh sites suggest consistent record of ~ 19-20 significant events in the Holocene.
- Bayesian models can help quantify and evaluate the numerous inputs to paleoseismic models (19 in this case). Improvements in the data result in improvements to the model.

Thanks for your attention!



A seal is swimming underwater, viewed from above. The seal's head is in the center, with its eyes and nostrils visible. Its flippers are extended outwards. The water is a clear, light blue-green color. There is a trail of white bubbles behind the seal's head, indicating its path. In the lower-left corner, the word "Questions?" is written in yellow text on a semi-transparent green rectangular background.

Questions?

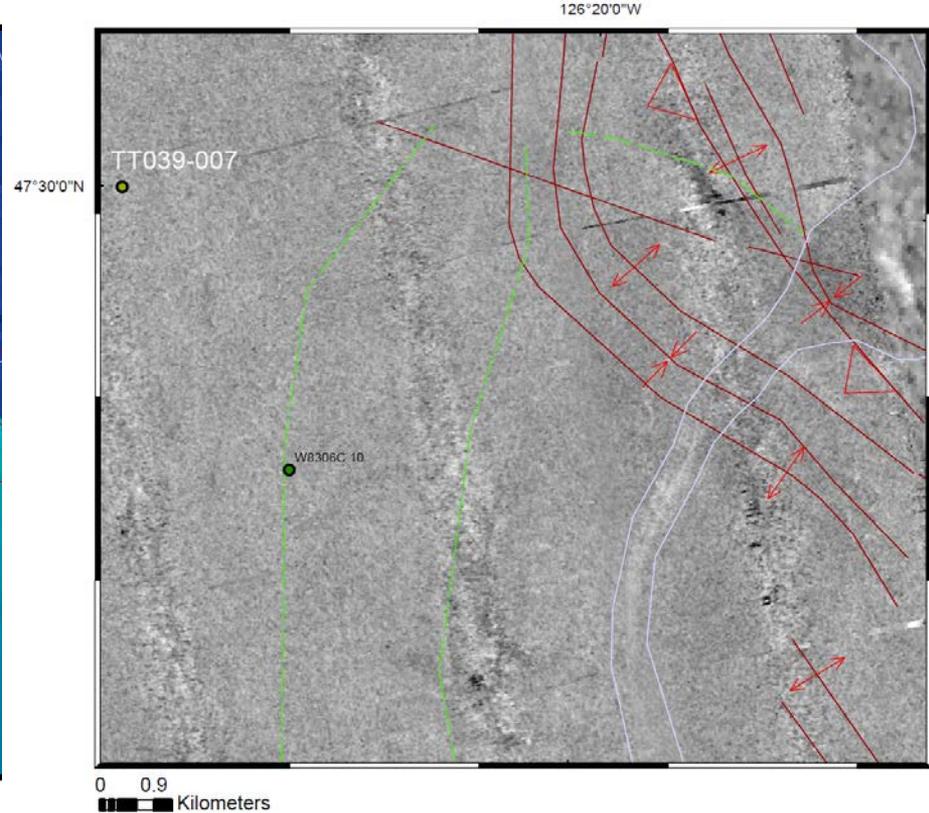
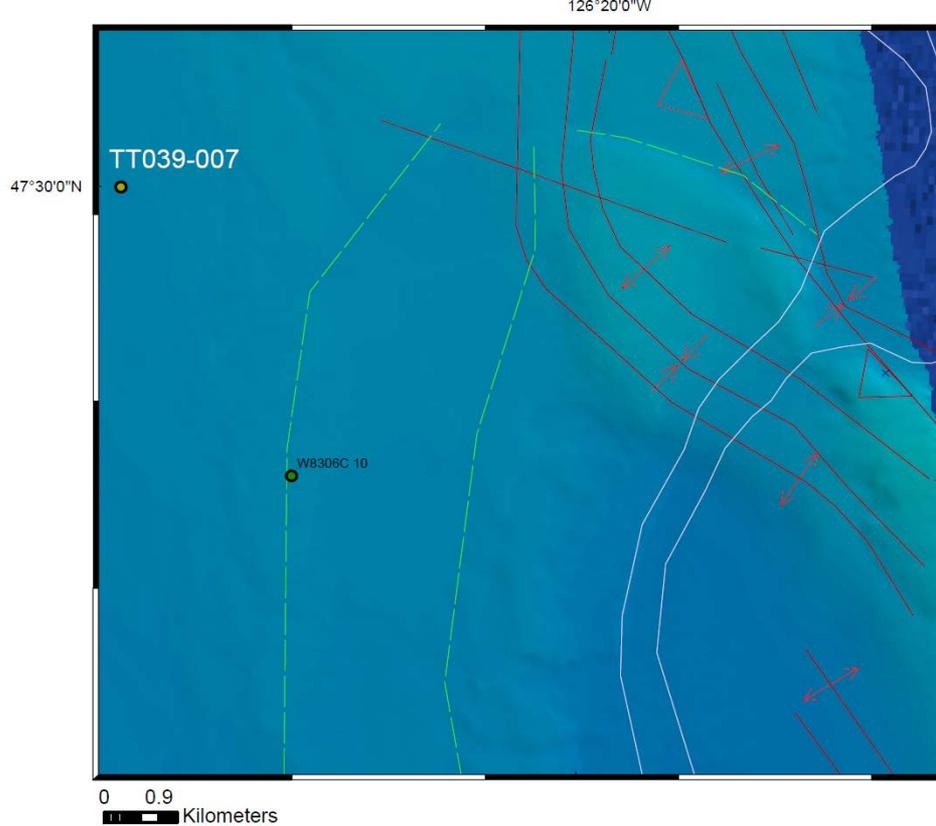
Willapa Wave field



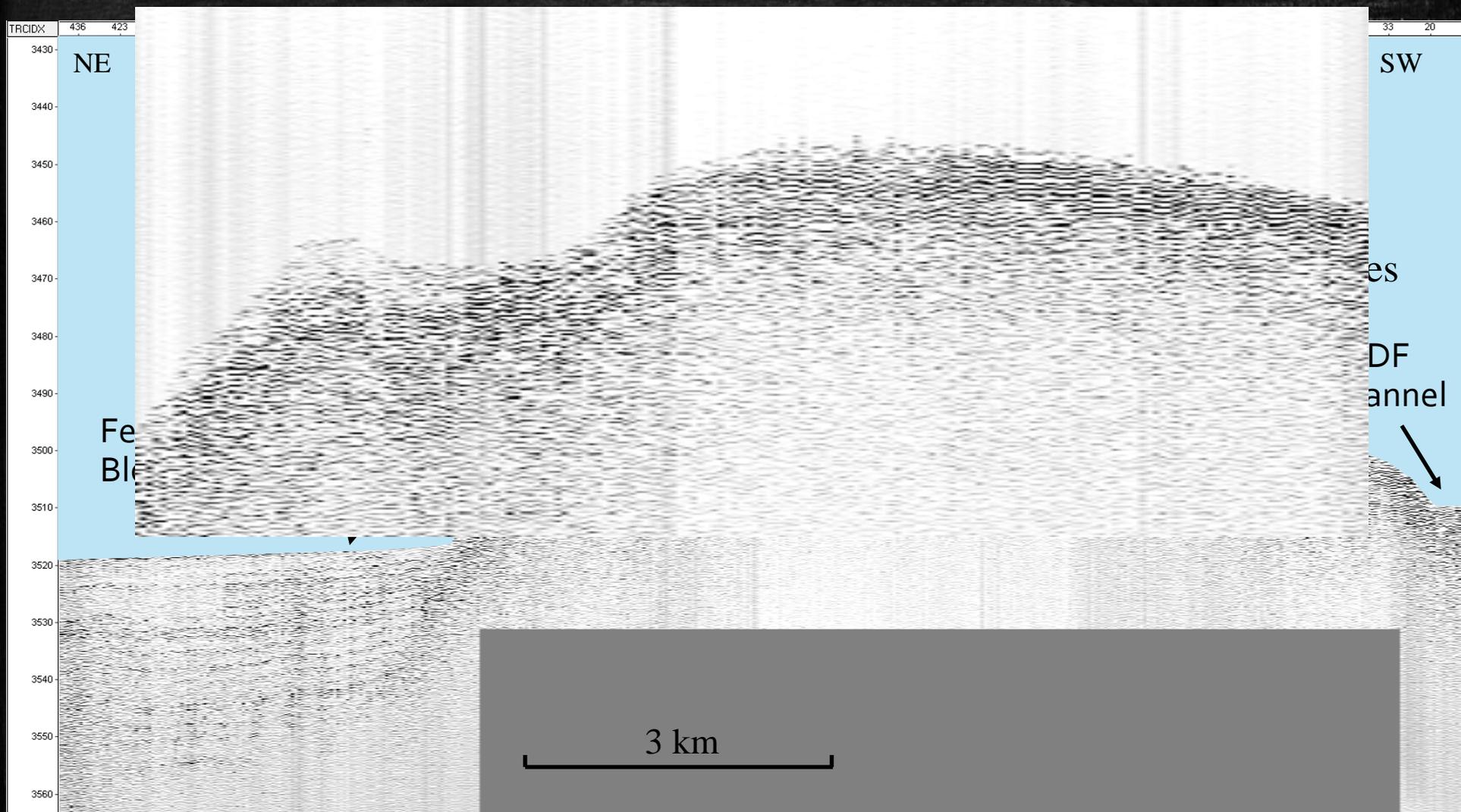
JDF Channel

Willapa Channel

The Holocene record in this core consists of ~ 50% hemipelagic mud, and 50% mud turbidites at a range of 12.5 km from the main Willapa bend. This area should be affected by the blowout events, but lack of Holocene record supports and older time frame.



The frontal thrust blocks both JDF and Quillayute, forcing JDF to go over the ridge, and redirecting Quillayute several km north where it joins JDF. A fine grained sediment pool results here as well. Walls of a paleo Quillayute Channel are visible, as is the sand tongue exiting this second bottleneck.

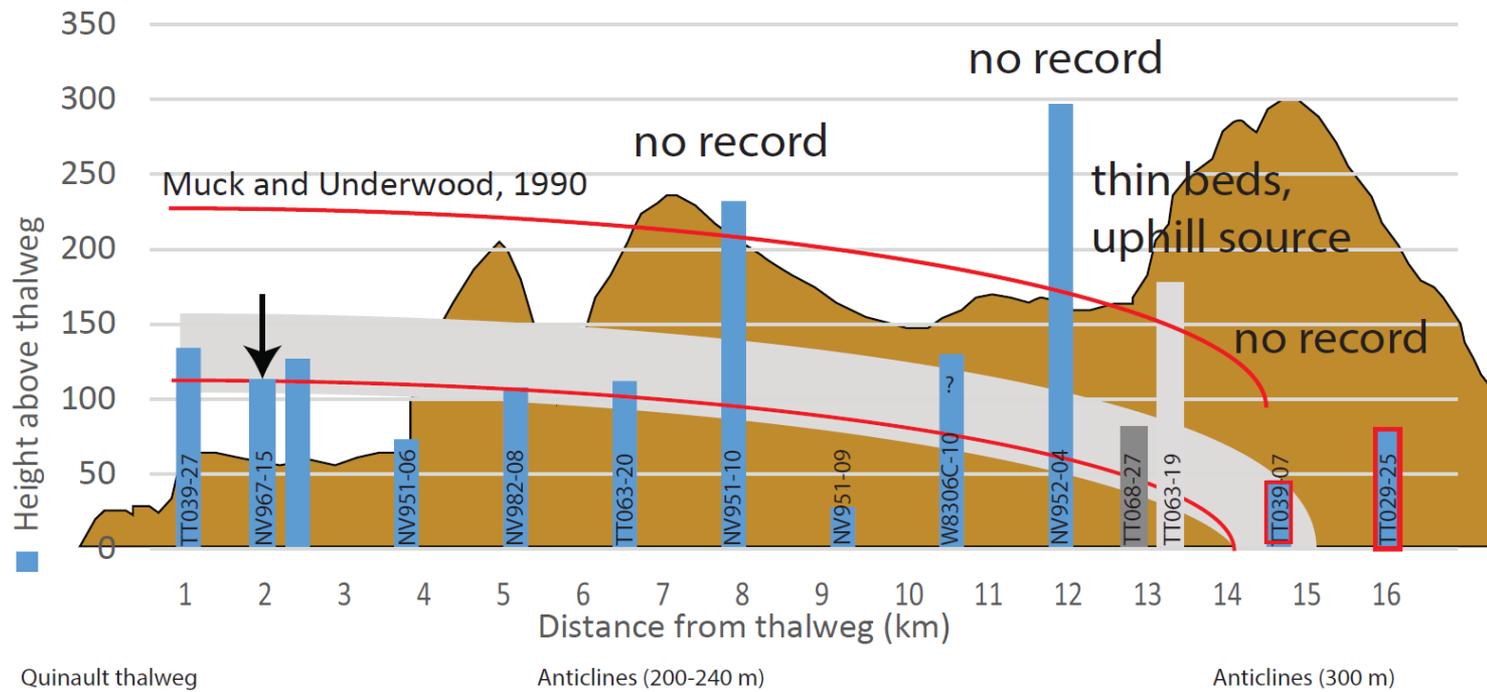


High-resolution chirp profile across the Feature A Blowout Basin to JDF Channel. Sandy turbidites sourced from JFD give way eastward to nearly transparent Holocene fill in the basin, consistent with backscatter data.

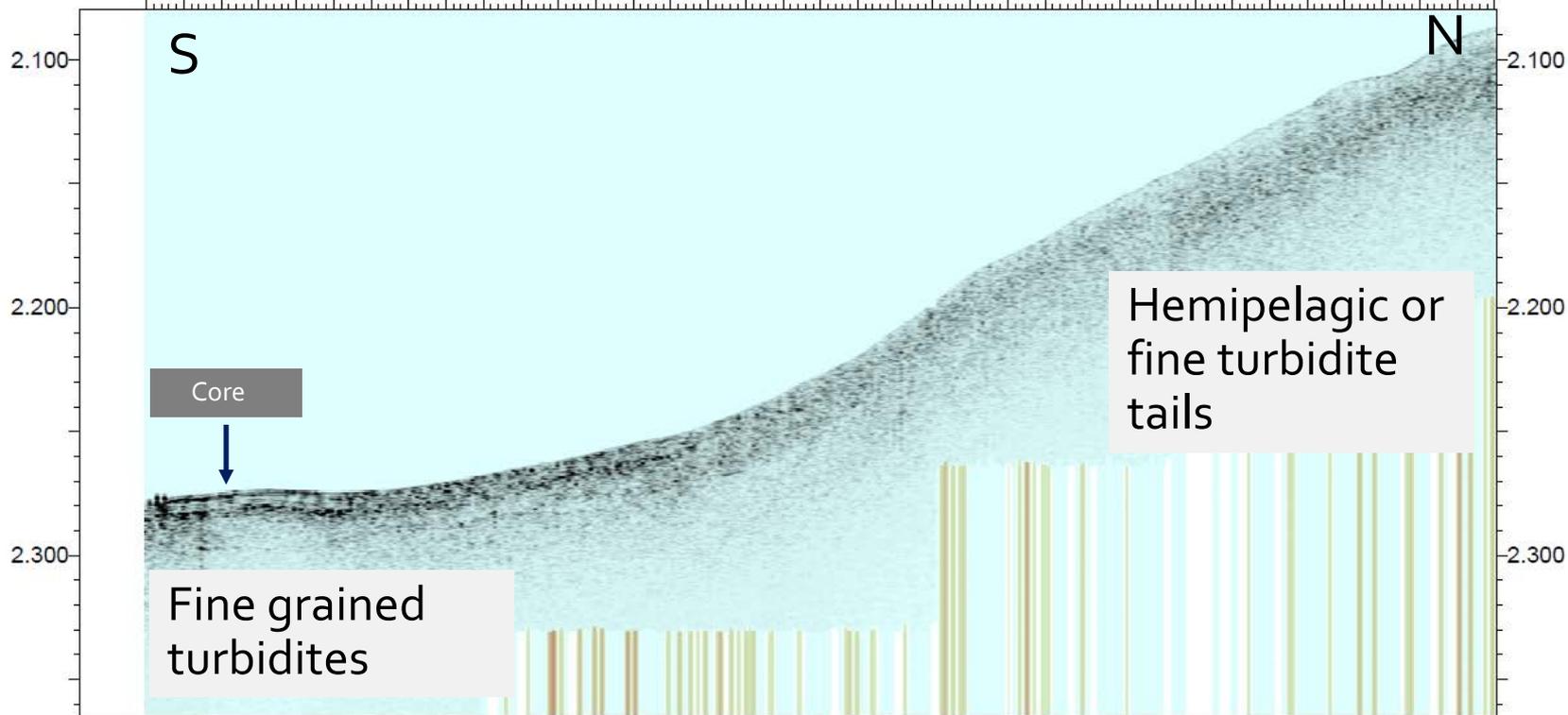
- A common misconception is that with evidence of M_g earthquakes, it should be simple to examine cores for this evidence which should be widespread.
- Surprisingly though, there are not as many sites with good paleoseismic records as one might think.
- Moreover, Cascadia is characterized by fine grained turbidite systems for the most part, particularly in the northern relict canyons. This means that techniques used must be capable of reliably detecting and characterizing fine grained deposits in detail.
- Lets look first at the aerial distribution of Holocene turbidite stratigraphy.

Model and empirical considerations

JDF and Quinault, Current Profile



SP: 23156.0 23200.0 23400.0
Offset: 0 500 1000 1500 2000 2500 3000 3500 4000 4500 5000 5500 6000 6500 7000 7500 8000 8500 9000



Seismic Micro-Technology, Inc.
Project: West_Coast_CHIRP
Project Location:
Line 0043_2011_151_2249_100212_CHP3.5_RAW_095_ED.SGY.NAV. Amplitudes [Filter ON 500, 1000, 2000, 2500]