

Bedrock detection using 2D electrical resistivity imaging along the Peikang River, central Taiwan

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ABSTRACT

We use the two-dimensional (2D) electrical resistivity imaging (ERI) method to detect the contact between sediment and bedrock in a modern fluvial system. We performed eight 2D electrical resistivity surveys along the Peikang River in Kuohsing area, central Taiwan, with a pole–pole array of 1-m electrode spacing to produce a spatial resolution of about 1 m. The resistivity data were inverted into subsurface electrical structures using the least-squares inversion techniques. Then, the contact between the bedrock and sediment was found by using an image processing technique called Laplacian edge detection (LED), which represents an objective approach in interpreting various geophysical images. Numerical modeling demonstrates that iso-resistivity lines are not in good agreement with the hypothetical interfaces between the bedrock and sediment, but the zero-lines of the Laplacian operation are strikingly accurate. Our results indicate a well-defined boundary in the resistivity structure that can be used to estimate the quantity of sediments covering bedrock, thus highlighting the utility of this technique in studies of landscape evolution, sediment transport, and sediment budgets.

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1. Introduction

The erosion of bedrock is an important mechanism that shapes topographic features and links climatic and tectonic processes. Previous research has shown that sediment in fluvial systems plays a dual role as it can both protect the bedrock from erosion and also abrade it (Sklar and Dietrich, 2001, 2004). The quantity of sediments covering bedrock is an important control on river incision (Korup and Montgomery, 2008). To fully understand how the transport and distribution of sediment affects the incision process, we need to evaluate the quantity of sediment covering bedrock. A reliable estimation of the contact between the sediments and the bedrock is thus crucial for exploring the volume of sediment in river bottoms. By using nondestructive geophysical investigation techniques, such as the two-dimensional (2D) electrical resistivity imaging (ERI) method, we can quantify the magnitude and variability of sediment covering bedrock faster, cheaper, and over a broader area than drilling a well or digging a trench (Hauck et al., 2003; Crook et al., 2008). The 2D ERI method has been a powerful technique to investigate shallow subsurface electrical structures in various environments (Cheng, 2000; Yang et al., 2002; Hauck et al., 2003; Sass, 2007; Cheng et al., 2008; Crook et al., 2008).

The 2D electrical resistivity imaging method, because of its high resolution, is commonly used in measuring the electrical structure of shallow subsurface within several tens to hundred meters deep. A 2D electrical resistivity imaging survey is produced by injecting current into the ground through two current electrodes and measuring the resulting voltage difference at two potential electrodes. This process is repeated for many current/potential electrode configurations to produce a pseudosection of apparent resistivity. The resultant data are then processed using a 2D inversion. Variations in the underlying sediment or rock units indicated by different electrical resistivities can then be observed as strong gradients in resistivity. Although the subsurface resistivity distribution could then be interpreted and mapped by eye, we will show that this subjective method should be done with caution because of the diffusive nature of the electrical field. This paper aims at defining a better location of the electrical interface between a high-resistivity layer and a low-resistivity layer, like the bedrock–sediment contact, by applying the Laplacian edge detection (LED) method (Marr and Hildreth, 1980; Vafidis et al., 2005) to an electrical resistivity image. Thus, we can find the contrasts at where resistivity changing from two main resistivity layers in the resistivity image. We will present our model verification of the LED method and also its application to detect the bedrock/sediment interface along the Peikang River, central Taiwan (Fig. 1).

The Peikang River drains the western flank of the Backbone Range, cuts through the Hsueshan Range across strike, and continues on through the western foothills (Fig. 1). The continuous succession of

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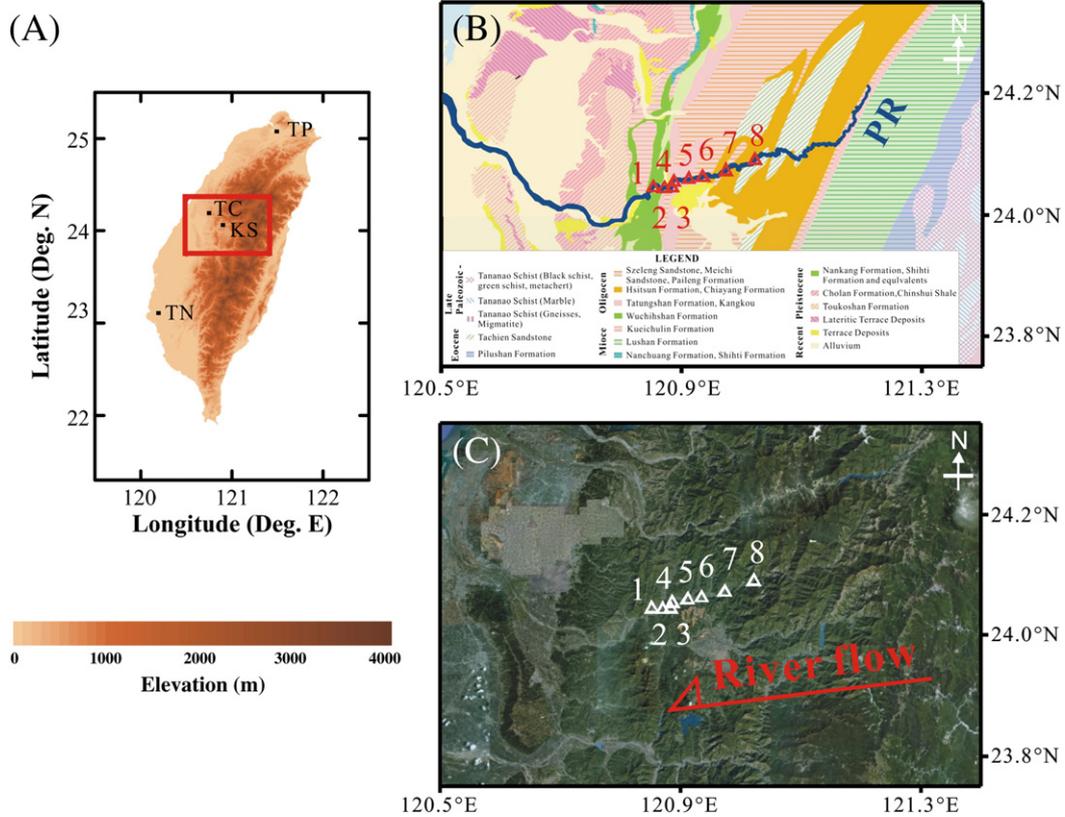


Fig. 1. (A) Digital elevation model for the island of Taiwan showing location of the study area (red box). TP = Taipei; TC = Taichung; TN = Tainan; and KS = Kuohsing. (B) Geological map of the Peikang River (PR, dark blue line), central Taiwan, as noted in (A). Triangles indicate site locations of the 2D ERI surveys. (C) Satellite image of (B). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Oligocene to Miocene strata in the Kuohsing area, central Taiwan, consists of the following formations in ascending order (Mao et al., 2002): Paileng Formation, Shuichangliu Formation, Shihszeku Formation, Tanliaoti Formation, Shihmentsun Formation, Changhukeng Formation, Talu Formation, Kuanyinshan Sandstone, and Kueichulin Formation. The river in the Kuohsing area is incising Miocene metasedimentary rocks over the Holocene (Yanites et al., 2008). Data from a local gauging station suggest annual flood discharges on the order of hundreds of m³/s with a maximum daily value of 1800 m³/s in 2004. Most importantly, this particular reach of the Peikang River is relatively unaltered by human disturbance. There are no major dams or other such structures throughout. Thus, we are seeing the sediment cover genuinely produced by natural processes.

2. 2D electrical resistivity imaging (ERI) method and data collection

Traditionally, the direct-current (DC) electrical survey is used to determine the subsurface resistivity distribution by measuring the electrical potential difference between a pair of potential electrodes (M and N in Fig. 2) on the ground surface with a current applied through a pair of current electrodes (A and B in Fig. 2) (Keller and Frischknecht, 1966; Telford et al., 1990). The apparent resistivity ρ_a in ohm-meter (Ω -m) is then computed from Ohm's law:

$$\rho_a = 2\pi n \Delta v / I \tag{1}$$

where n is distance in meters between the electrodes; Δv is the measured potential difference in mV; and I is the applied electric current in mA.

Recent developments in DC technology allow automatic measurements, by switching the current and potential electrodes between a

series of equally spaced electrodes laid out along a profile (Fig. 2B–D) (e.g., Yang et al., 2002; Hauck et al., 2003). This allows a dense sampling of subsurface resistivity variation at shallow depth within a short amount of time. We use the standard pole–pole array to map the near-surface electrical structures. This array is configured by setting one current (A) and one potential (N) electrode to be placed at a distance along the profile of “infinity.” In practice, this distance is often over 10 times the distance between the starting and ending electrodes used in the profile (B₁ and M_{end}^{end} in Fig. 2). This allows their influence to be ignored with respect to the close electrodes B and M (Keller and Frischknecht, 1966; Telford et al., 1990; Cheng et al., 2008). The ERI data are then collected by using a series of the equally spaced electrodes (B_{*i*} and M_{*i*}^{*i*} in Fig. 2B). The automatic survey is starting at one end of the profile, using the first central electrode as the current electrode B. The electrical potential is then measured sequentially along the profile using each following central electrode as the potential electrode M (Fig. 2B). Then, the current electrode is automatically switched to the second electrode along the profile and the potential measurement is conducted over the subsequent electrodes (Fig. 2C). The automatic switches are carried on from one end to the other until the survey along the profile is finished (Fig. 2D). This produces a subsurface map of the “apparent” resistivity distribution.

The apparent resistivity distribution of the subsurface structure is then inverted using the commercial RES2DINV[®] software to estimate the true resistivity structure. The algorithm uses a 2D smoothness-constrained, least-squares inversion with a Jacobian matrix calculation for the first iteration and then employs a quasi-Newtonian technique to reduce numerical calculations (Loke and Barker, 1995, 1996). The inversion is stopped once the difference of the root mean square (RMS) error between the current and previous iterations is

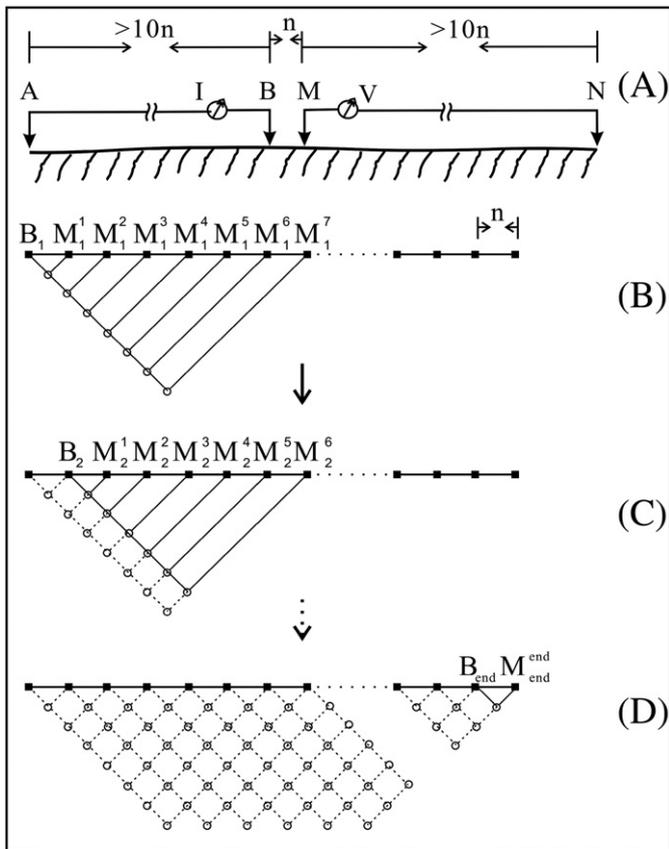


Fig. 2. Electrode configuration of the pole-pole array used in this study. (A) A sketch about the relationship of the current (A–B) and potential (M–N) electrodes. For the pole-pole array, the electrodes A and N should be located far away the central electrodes B and M at ~ 10 times the distance between B and M. (B)–(D) Automatic measurement of the pole-pole array. An automatic ERI survey begins at one end of the central group of B_i-M_i electrodes and works by switching sequentially the electrodes to the other end of the central electrode array. For details please refer to the text.

$< 0.1\%$. The inverted data produce the 2D resistivity distribution map, which can then be used for extracting information about the contact between sediment and bedrock.

The 2D ERI data in this study were collected at eight sites along the Peikang River in the Kuohsing area, central Taiwan (Fig. 1). At all sites, surveys were conducted mostly perpendicular to the river channel and located on the dry river beds, below bankfull flow, in a river dominated by summertime typhoon generated floods (Fig. 3). Then, to check for the reproducibility in the data, we conducted surveys parallel to the river channel at two sites, i.e., sites 2 and 7. We collected the high-resolution 2D ERI data by using an electrode spacing of 1 m, giving a vertical resolution of about 1 m. The infinity electrodes are located at distances of 10 times the profile length away from the center of each ERI profile. That is, for a 50-m-long ERI profile, the infinity electrodes were 500 m away; and for an ERI profile with 200 m, the infinity electrodes were 2 km away. Total length of the 2D ERI profile at each site depends on the width of the terrace in the river, which ranges from 47 to 191 m (Table 1).

3. Laplacian edge detection (LED) and its verification

There are many ways to objectively detect edges in an image (Marr and Hildreth, 1980; Vafidis et al., 2005). Two widespread categories of edge detection are the gradient- and Laplacian-based techniques. The difference between these techniques can be easily visualized with a 1D example as shown in Fig. 4. The gradient approach detects the edges of an image by finding the maximum of the first derivatives of

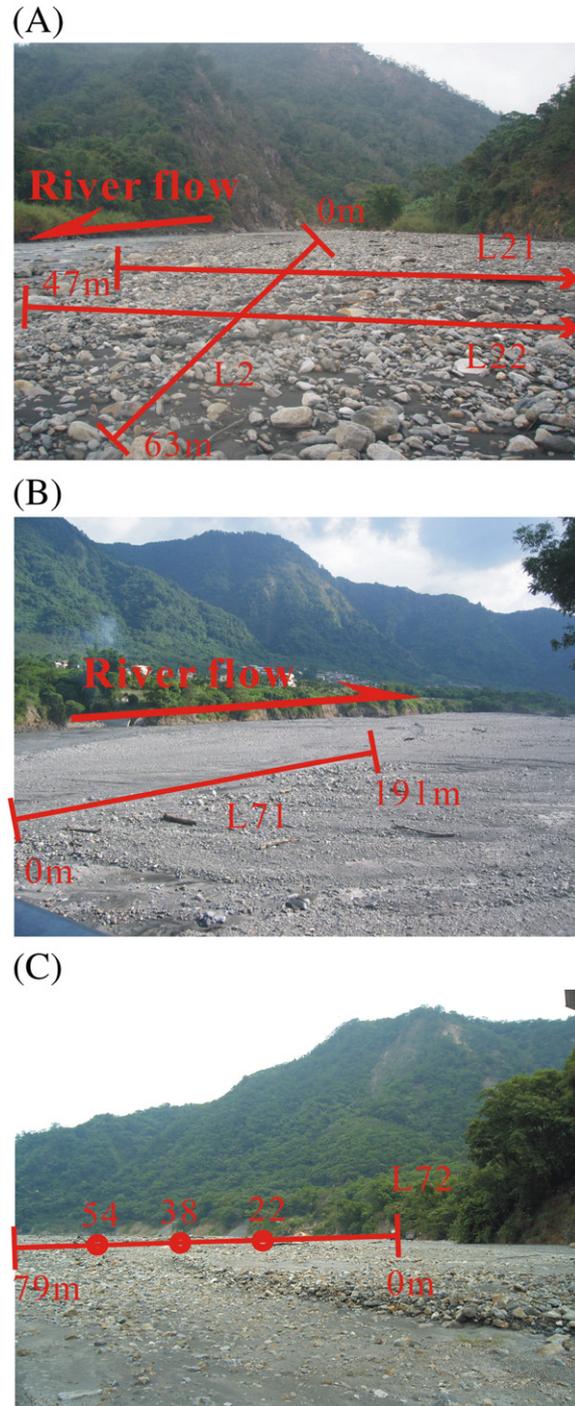


Fig. 3. Photographs of (A) site 2, (B) and (C) site 7. The solid lines mark the ERI profiles. The circles in (C) indicate the locations of the boreholes in Fig. 8D. As shown in these photographs, the ERI profiles were located on the dry river beds.

an image (Fig. 4B). The Laplacian method searches for values of zero in the second derivatives of the image (Fig. 4C). For a 2D image, where Z is the quantity of interest (e.g., resistivity) and x and y are coordinates, the Laplacian operator is defined as

$$\nabla^2 Z(x,y) = \frac{\partial^2 Z}{\partial x^2} + \frac{\partial^2 Z}{\partial y^2} \quad (2)$$

By calculating the Laplacian, we find where the Laplacian is zero (hereinafter we call it the Laplacian zero-line) it corresponds to the

Table 1
Information about eleven ERI lines located at eight sites (Fig. 1) along the Peikang River, central Taiwan.

ERI site/line #	Length of ERI profile (m)	Average and range of estimated depth of sediment-bedrock contact (m)	Perpendicular or parallel to the direction of river flow
Site 1/L1	79	4.0/2.4–6.9	Perpendicular
Site 2/L21	47	3.0/1.4–5.1	Perpendicular
Site 2/L22	47	3.0/1.3–5.0	Perpendicular
Site 2/L2	63	2.0/1.4–2.8	Parallel
Site 3/L3	95	5.2/1.1–9.9	Perpendicular
Site 4/L4	63	2.2/1.3–4.2	Perpendicular
Site 5/L5	47	2.6/1.5–3.5	Perpendicular
Site 6/L6	79	5.2/1.8–11.3	Perpendicular
Site 7/L71	191	11.8/6.2–15.3	Parallel
Site 7/L72	79	12.0/9.1–14.5	Perpendicular
Site 8/L8	79	8.3/1.6–13.4	Perpendicular

abrupt change (i.e., the edge) in the original data. This is appealing in electrical techniques because, as electric current flows through two materials with different electrical resistivities, the electrical charges are piled up at the electrical boundary. That is, electrons are discharged on one side of the boundary between two electrical media and recharged on the other side, producing abrupt changes in the inverted data. Thus, the Laplacian operator can, in a mathematical sense, provide a measure of *discharge* or *recharge* flux across a surface. We therefore expect that the LED method can be extremely valuable to the analysis of an electrical resistivity image.

To demonstrate the LED method, we modeled an arbitrary two-layered (high-resistance sediment over low-resistance bedrock) system with a saw-toothed interface (Fig. 5A). In addition, we also modeled anomalous high-resistivity blocks (e.g., gravel or sand lenses) embedded within the cover layer (Fig. 5B). The resistivity of the upper layer is set as 300 Ω m and the lower half-space layer 10 Ω m. For the second model in Fig. 5B, the high-resistivity anomalous blocks have a resistivity of 1000 Ω m and dimensions of 3 m in width and 2 m in thickness, comparable to our observations. Fig. 5A and B plot the inverted electrical images obtained from the forward-modeled resistivity data of the two models above. The gray solid lines show the hypothetical interface between the upper and lower layers (Fig. 5A), as well as the locations of the anomalous blocks (Fig. 5B). The black solid lines fairly coincident with the hypothetical interfaces are the Laplacian zero-lines obtained from the LED method. We also draw the iso-resistivity lines of 300 Ω m (the black dashed line above the interface) and 10 Ω m (the black dashed line below the interface). While those iso-resistivity lines are not in good agreement with the hypothetical interfaces, the zero-lines of the Laplacian operation are strikingly accurate. We have also tested the LED method for an electrical model with a low-resistivity layer over a high-resistivity layer (Fig. 5C). Consequently, the LED method can work as well to reveal the hypothetical boundary between electrical media.

Another concern associated with the fluvial system is the effect on the LED method of groundwater saturation. For the fluvial system in the dry season, the water-saturated fluvial sediment with relatively low resistivity is supposed to be beneath the unsaturated sediment. Without groundwater filling in the pore space, the upper dry unsaturated sediment is expected to take on high resistivity. We thus developed another three-layered model (Fig. 6) with a relatively high-resistivity layer (700 Ω m) on the top within a depth of 1.5 m. As shown in Fig. 6, the Laplacian zero-lines (black solid lines) together with large magnitude gradients (purple solid lines) can fit very well to the boundary (gray solid lines) between two main layers in the bottom. The dry unsaturated layer at the shallow depth looks like it is having little effect on the ability to reveal the bedrock–sediment contact of the LED method. The influence of groundwater table could thus be ignored in the analysis of field ERI data.

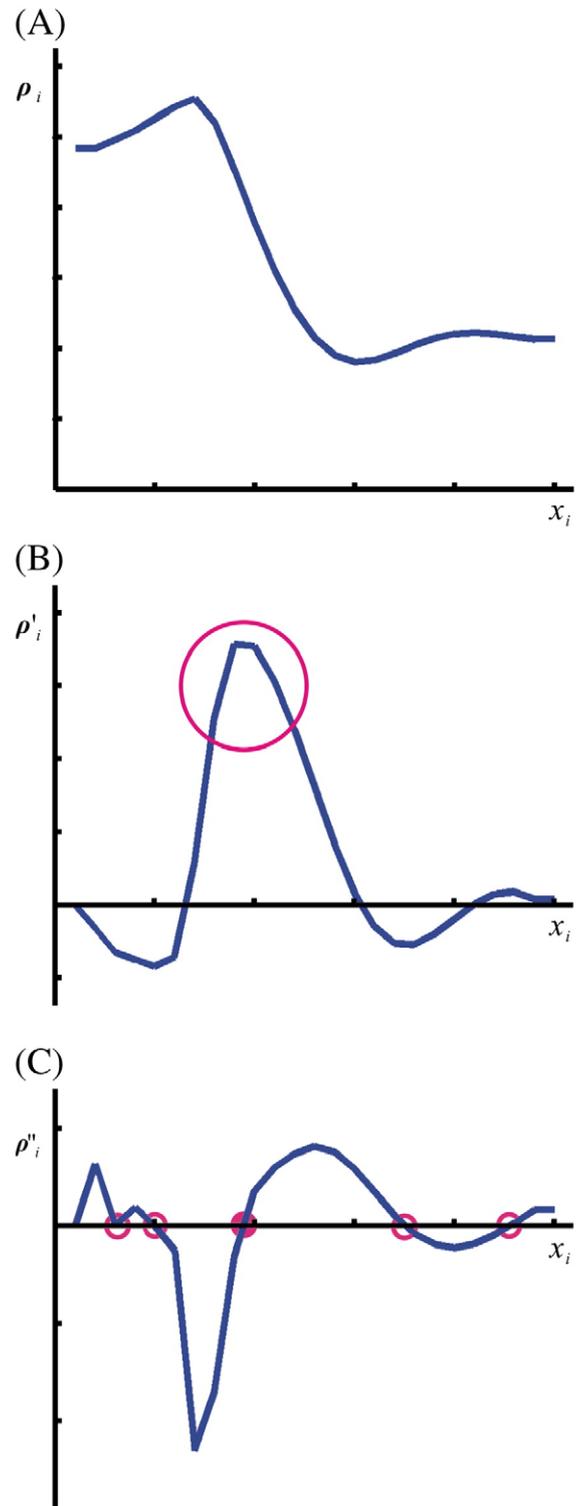


Fig. 4. One-dimensional example illustrating the LED concept: (A) original data. (B) The first derivative, gradient, of (A). Circle shows the areas with large gradient values indicating the possibly broad zone for the electrical boundary. (C) The second derivative, Laplacian, of (A). Circles show the locations with the zero Laplacian. One of the circles (solid one) can fairly locate the electrical boundary.

The above examples suggest that, when applying an external electrical current, the electric recharging/discharging process that occurs across the boundaries between electrical materials could be

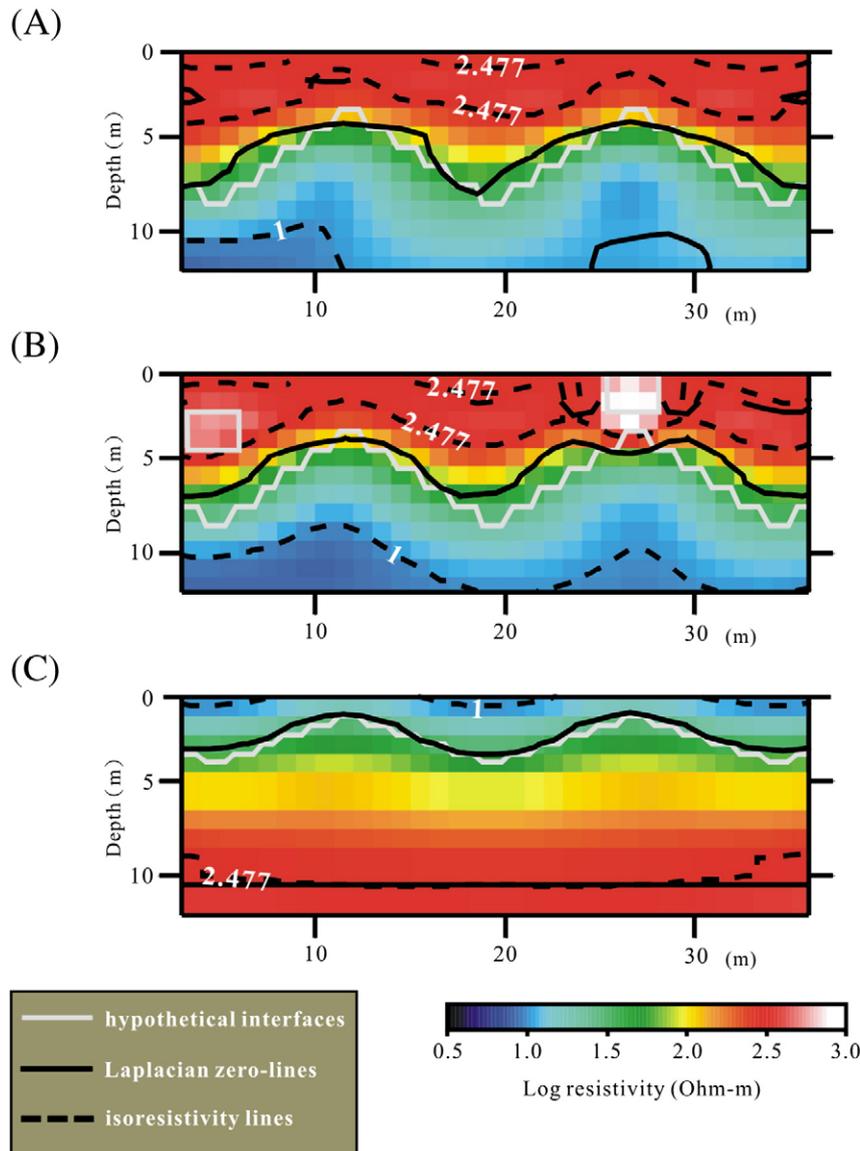


Fig. 5. Two-dimensional modeling results based on the LED method. The gray solid lines are the hypothetical boundaries between electrical media. The black solid lines are the Laplacian zero-lines. The black dashed lines in the red region are the isoresistivity lines of $300 (= 10^{2.477}) \Omega \text{ m}$, and in the blue region they are the isoresistivity lines of $10 (= 10^1) \Omega \text{ m}$. For the synthetic models in (A) and (B), the resistivity of the upper layer is set as $300 \Omega \text{ m}$ and the lower half-space layer $10 \Omega \text{ m}$. We also added small blocks with a very high resistivity of $1000 \Omega \text{ m}$ as the electrical anomalies embedded within the upper layer in model (B). (C) An inverse pattern of electrical model with low-resistivity ($10 \Omega \text{ m}$) over high-resistivity ($300 \Omega \text{ m}$) layers. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

characterized by the transition zone in an electrical image. We should note that we have chosen those Laplacian zero-lines with larger (top 20%) gradients as the estimated boundaries (solid circle in Fig. 4C), instead of those Laplacian zero-lines with smaller gradients (open circles in Fig. 4C). We did not merely consider the transition zone (orange to green areas in Figs. 5 and 6) as the electrical boundary because of the wide depth range of that zone. The drawback of the LED method is that it also shows some local Laplacian zero-lines around the shallow anomalous blocks as shown in Fig. 5B. Theoretically speaking, those small-scale zero-lines represent the real boundaries between the high-resistivity anomalies and the host matrix. Practically speaking, they represent noise induced by those shallow anomalous bodies that has nothing to do with the bedrock–sediment contact we are targeting. Therefore, special attention is required if such noise structures are present in the interpreted electrical images. For a better estimation of the bedrock–sediment contact, we suggest ignoring the small-scale distorted Laplacian zero-lines not associated with large magnitude gradients in the resistivity values.

4. Field example of ERI and LED: the Peikang River, central Taiwan

Shown in Fig. 7 are raw (upper panels in Fig. 7A–C) and inverted (middle panels in Fig. 7A–C) pseudosections for the ERI surveys at sites 1, 8, and 7 along the Peikang River, together with the corresponding best-fit electrical models (lower panels in Fig. 7A–C). Three sections of best-fit electrical models at site 2 are also plotted in Fig. 7D. A similar pattern of subsurface resistivity distributions that depicts a high-resistivity zone above a low-resistivity layer could be found in all of the ERI profiles. The sediment cover has a high resistance because the fluvial deposits produce a relatively high porosity covering the bedrock. The resistivity of the small apertures of pore spaces between these deposits is usually of lower conductance than the consolidated bedrock. This produces the distinct changes in the resistivity distributions, indicating the contact between unconsolidated sediments and consolidated bedrock. Thus, we expect to see the transition in the resistivity distribution changing from an upper resistive covering layer to a lower conductive underlying layer.

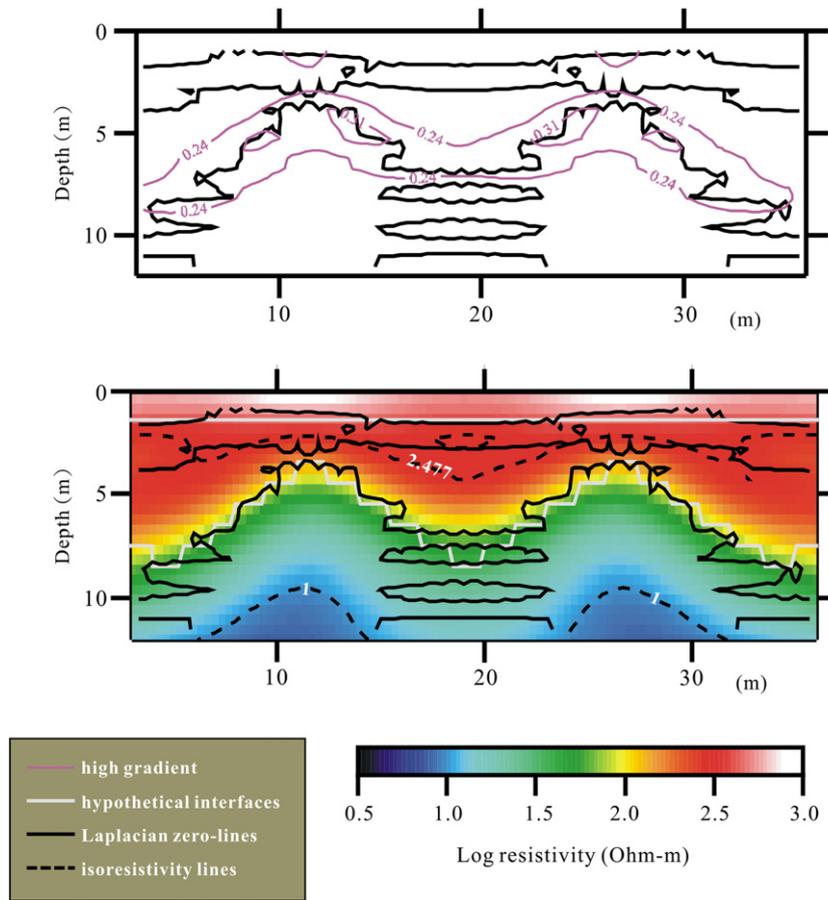


Fig. 6. The LED result for a hypothetical three-layered model with a very high resistivity ($700 \Omega \text{ m}$) layer indicating thin unsaturated sediment on the top within a depth of 1.5 m. The gray solid lines are the hypothetical boundaries between electrical media. The black solid lines are the Laplacian zero-lines. The black dashed lines are the iso-resistivity lines. The purple solid lines are the isogradient lines. Note that the Laplacian zero-lines bounded by large magnitude gradients are in very good agreement to the hypothetical boundary between the two main layers in the bottom, which suggests that the LED method is not much affected by the top dry layer. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

By applying the LED technique, we then estimate the contacts between the fluvial sediments and underlying bedrock from the inverted resistivity distribution maps. As examples, we show the contact obtained by LED at sites 1, 7, and 8 (Fig. 8). After taking the logarithm of the resistivity data and reimagining the inverted structures, we find that the resistivity distribution could be divided into three general parts: the top layer (red) with high resistivity of about $300 \Omega \text{ m}$, the bottom layer (green or blue) with low resistivity of about $30 \Omega \text{ m}$, and a transition layer between these end members. We could *subjectively* define the thin transition layer as the contact between the upper resistive sediments and the lower conductive bedrock; however, based on our simple model demonstrated above, we could also *objectively* define the contacts as the Laplacian zero-lines (white lines in Fig. 8) with large magnitude gradients (blue lines in Fig. 8). The modeling results shown in Section 3 suggest that the Laplacian zero-line can define the contact better than choosing an iso-resistivity line as the contact. Therefore, following the LED analysis, we have in Fig. 8 drawn our estimated bedrock–sediment contact (red lines) for each ERI profile.

Digging the trenches/boreholes in some areas is sometimes difficult (e.g., the conservation area for water resources), and is usually more expansive than the nondestructive geophysical techniques. However, after getting some extra budget and complicated official approval, we were allowed to dig some boreholes (Fig. 3C), which constitutes a very important independent validation of the sediment thickness. Although only three boreholes were dug at site 7

because of a limited budget, they provide a valuable check on the ERI and LED results mentioned above. As shown in Fig. 8D, our estimation based on the ERI-LED analysis could be fairly close to the real location of the sediment–bedrock contact.

Yet another way of verification could be based on the establishment of a reliable statistical analysis. Fig. 9 shows the resistivity distribution for all electrical images acquired along the Peikang River. At the high-resistivity end, those resistivities show a well-defined log-normal distribution, i.e., the Gaussian distribution of the logarithm of resistivity data (red curve in Fig. 9). Because of the symmetry of the normal distribution in logarithmic resistivity, we can easily find an approximate log-normal distribution with the logarithmic mean of 2.452 ($\sim 300 \Omega \text{ m}$) and the standard deviation of 0.215 to fit the data at the high-resistivity end. A heavy tail in the low-resistivity end can also be seen in the resistivity distribution. Probably, the electrical resistivity of sediments around the whole study area could be considered as the random variable because of the log-normal distribution of grain size of fluvial sediments (Kodama, 1994; Rice and Church, 1996). Note that the resistivity of fluvial sediment is basically dominated by two factors: porosity and water saturation of pore spaces (Archie, 1942). Therefore, the normal bell shape in the high-resistivity end suggests the resistivity distribution of covering sediments, while the skew to the low-resistivity end indicates the gradual change in the electrical property of the rock as a consequence of their lithological property. We thus expect that the contact between the sediments and bedrock should be located near the margin of the

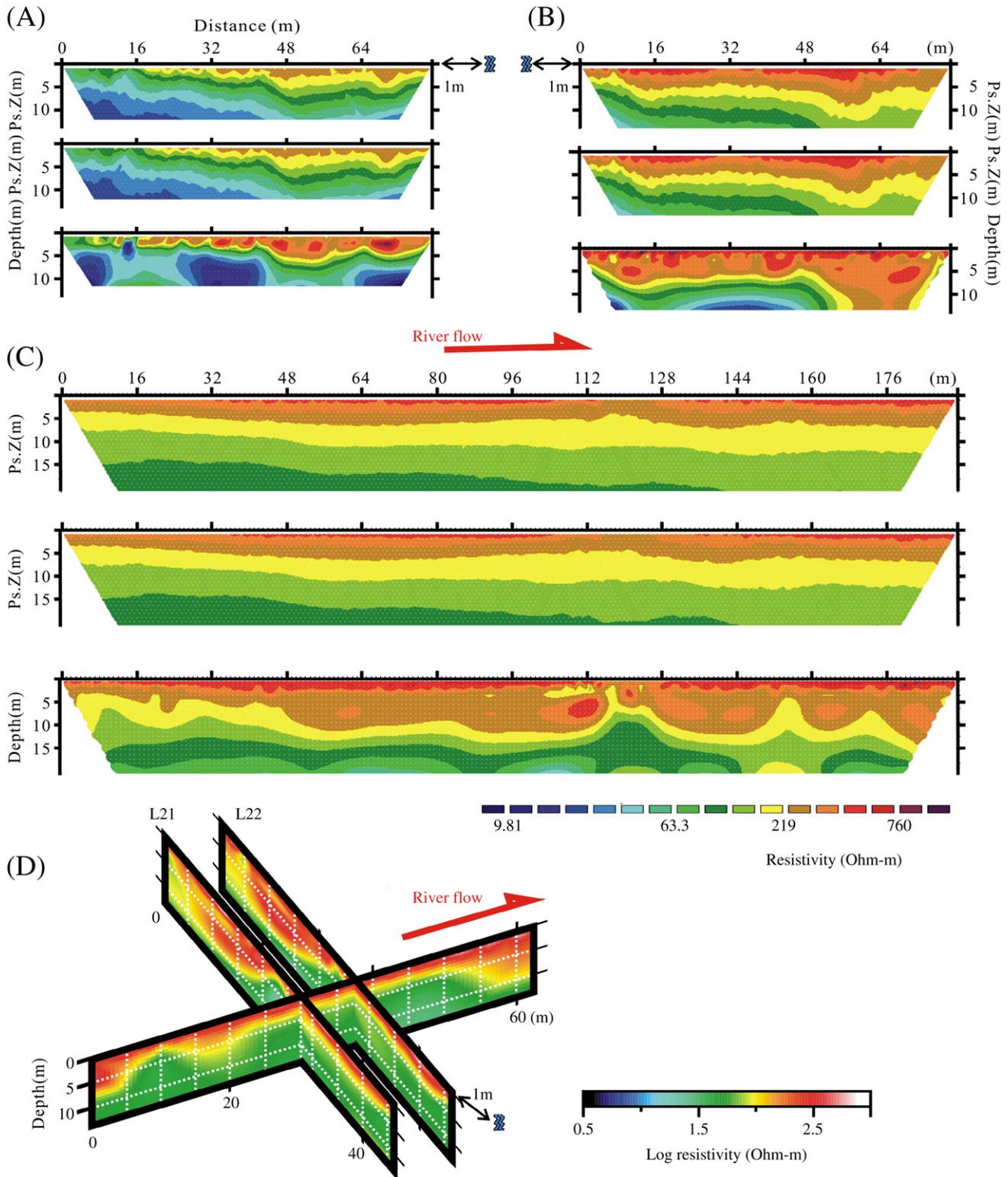


Fig. 7. ERI inversion results at (A) site 1 (ERI profile L1), (B) site 8 (ERI profile L8), and (C) site 7 (ERI profile L71), respectively. Each result at a site includes three panels: the observed apparent resistivity pseudosection (upper), the theoretical pseudosection of the best-fit inversion model (middle), and the best-fit electrical model (lower). Note that the vertical axis is the pseudo-depth (Ps.Z) for the upper and middle plots. (D) The best-fit electrical models of the ERI profiles L2, L21, and L22 located at site 2. The symbol of water flow indicates the relative location to each ERI profile of the river channel.

high-resistivity Gaussian bell shape. As shown in Fig. 9, the white cross denotes the range of resistivities of those areas where the Laplacian are zero for the electrical profiles. The horizontal elongation indicates one standard deviation for the resistivity values of those

where Laplacian is zero. Such a result of statistical verification suggests that the LED method can give a fairly consistent, reliable definition of the boundary between the fluvial sediments and bedrock from the resistivity image.

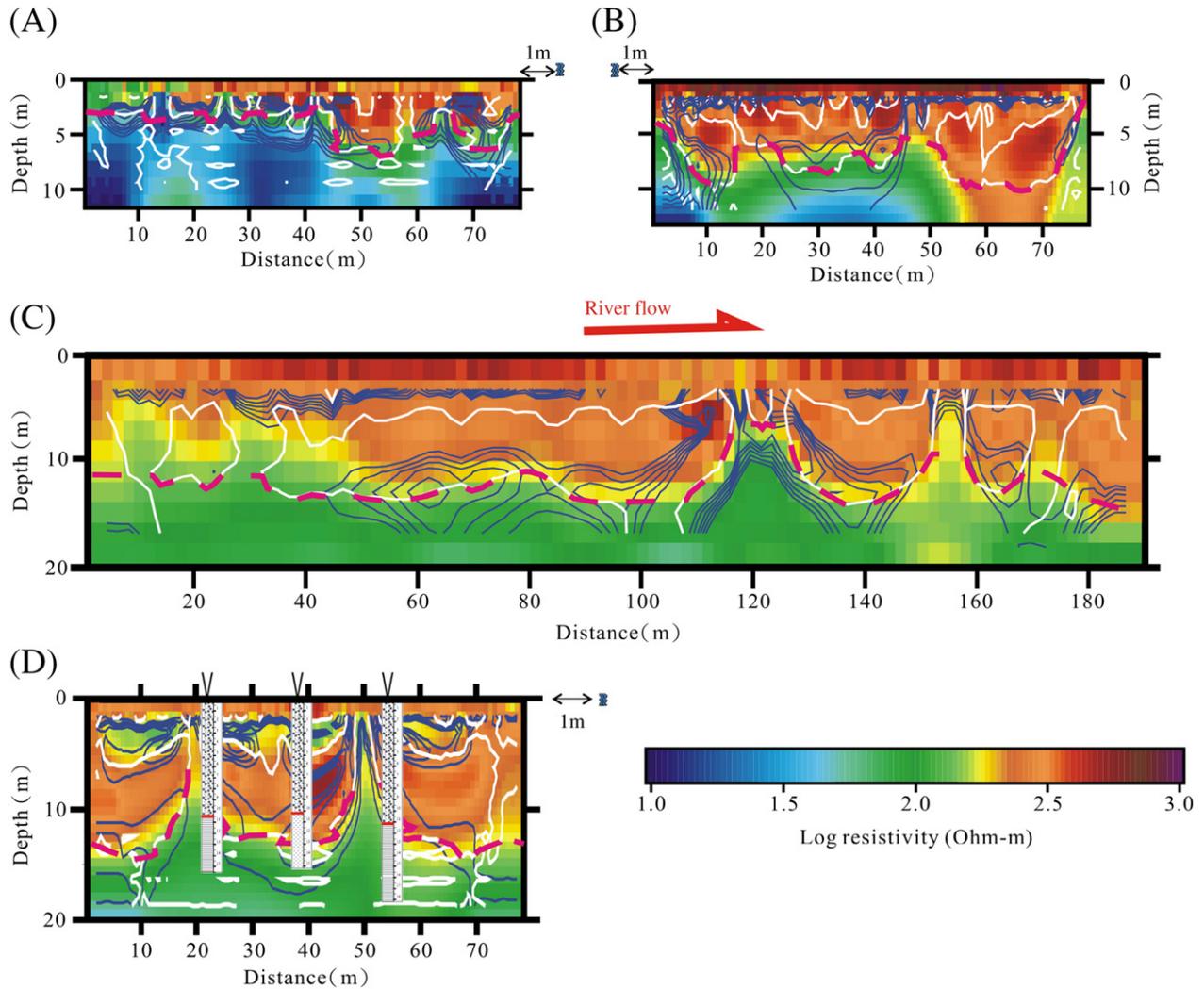


Fig. 8. LED results for the ERI profiles (A) L1 at site 1, (B) L8 at site 8, (C) L71, and (D) L72 at site 7. The symbol of water flow indicates the relative location to each ERI profile of the river channel. Blue lines are isograd lines with large magnitude. White lines are the Laplacian zero-lines. We interpret the Laplacian zero-line as having large magnitude gradient at the sediment–bedrock contact, as indicated by red dashed line. For independent validation of the bedrock–sediment contact depth estimated from the LED analysis, three boreholes were dug at 22, 38, and 54 m along the ERI profile L72 (Fig. 3C). The bedrock–sediment contact for these boreholes is at 10.5, 10.2, and 11.2 m deep, respectively. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Over the whole study area we can thus estimate the depth of sediment–bedrock contact, which ranges from several to ten meters deep. Table 1 summarizes the length of the ERI profile and the average and range of the depth of sediment–bedrock contact at each site. Accordingly, we took the mean cross-sectional area of sediments from the surveys orientated perpendicular to the river flow direction and then multiplied it by $\sim 30,000$ m, i.e., the length of the study reach. We can thus estimate that roughly 12 million m^3 of sediment is stored in the modern floodplain above the bedrock of the Peikang River.

5. Conclusion

In this study, the 2D ERI method was used to detect the boundary of sediments and bedrock along the Peikang River, central Taiwan. As shown in Fig. 7D, comparison between one river-parallel and two river-perpendicular profiles at site 2 suggests that the electrical models obtained in this study are very much reproducible and in good agreement with the resistivity pattern. We find a consistent distribution in all the 2D ERI profiles (Fig. 7), where a shallow high-resistivity zone (interpreted to be fluvial sediment) covers a low-resistivity zone (interpreted to be bedrock). Apparently, resistivity of

the sediment follows a log-normal distribution with an average of about $300 \Omega \text{ m}$. The broad distribution of resistivity in these sediments is likely influenced by randomness associated with grain size, pore space, and level of water saturation.

We used the LED method to locate the boundaries in the 2D electrical resistivity images. The utility of the LED method was demonstrated by numerical modeling of a hypothetical two-layered structure (Fig. 5), which revealed that the Laplacian zero-line is coincident with the hypothetical boundary while the selection as the boundary of the iso-resistivity lines misses the contact by many meters. Aided by the LED method, this study found a good estimation of the contact between sediments and bedrock along the Peikang River, which is useful for understanding the influence of fluvial sediment cover on incision process (Yanites et al., 2008). Statistical analysis of the resistivity distributions (Fig. 9) and borehole data (Fig. 8D) suggest that the LED method gives reliable estimates of the electrical boundary between the sediment and bedrock. We have estimated that roughly 12 million m^3 of sediment is stored in the modern floodplain above the bedrock of the Peikang River. Data such as this is necessary to close a sediment budget and calculate residence times in a fluvial system (Malmon et al., 2003). Our methodology

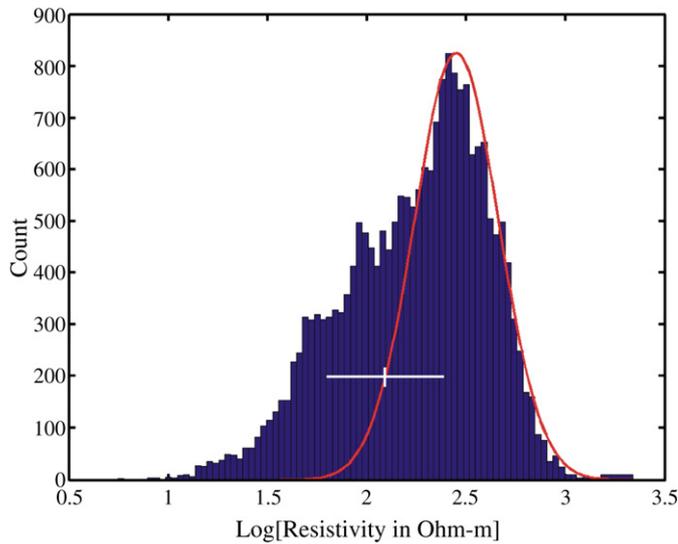


Fig. 9. Histogram of log-resistivity from all measured ERI profiles along the Peikang River. The red curve is an approximate log-normal distribution with the mean of 2.452 and the standard deviation of 0.215, indicating the resistivity distribution of fluvial sediments. The white cross marks the range of resistivity values at the likely boundaries between the sediment and bedrock estimated from the LED method. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

presented in this paper therefore offers an exciting new technique to study sediment systems in bedrock landscapes.

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